

# A numerical study on the winter mixed layer on the shelf-slope region south of Japan

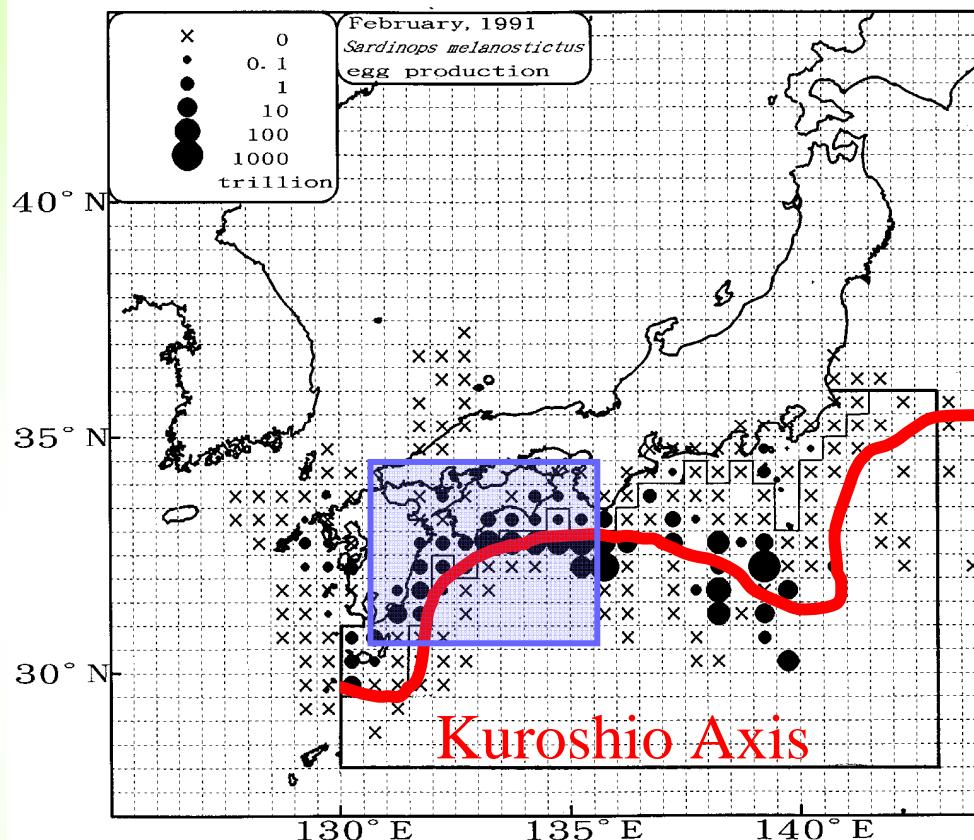
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Yuichi Hirota\*  
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# Motivation of this study



Distribution of sardine eggs in February 1991  
(Watanabe et al., 1997)

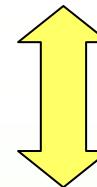
+January to March

Spawning season of sardine

+Main spawning ground

southwest of Japan

Inshore side of the Kuroshio



Missing Link

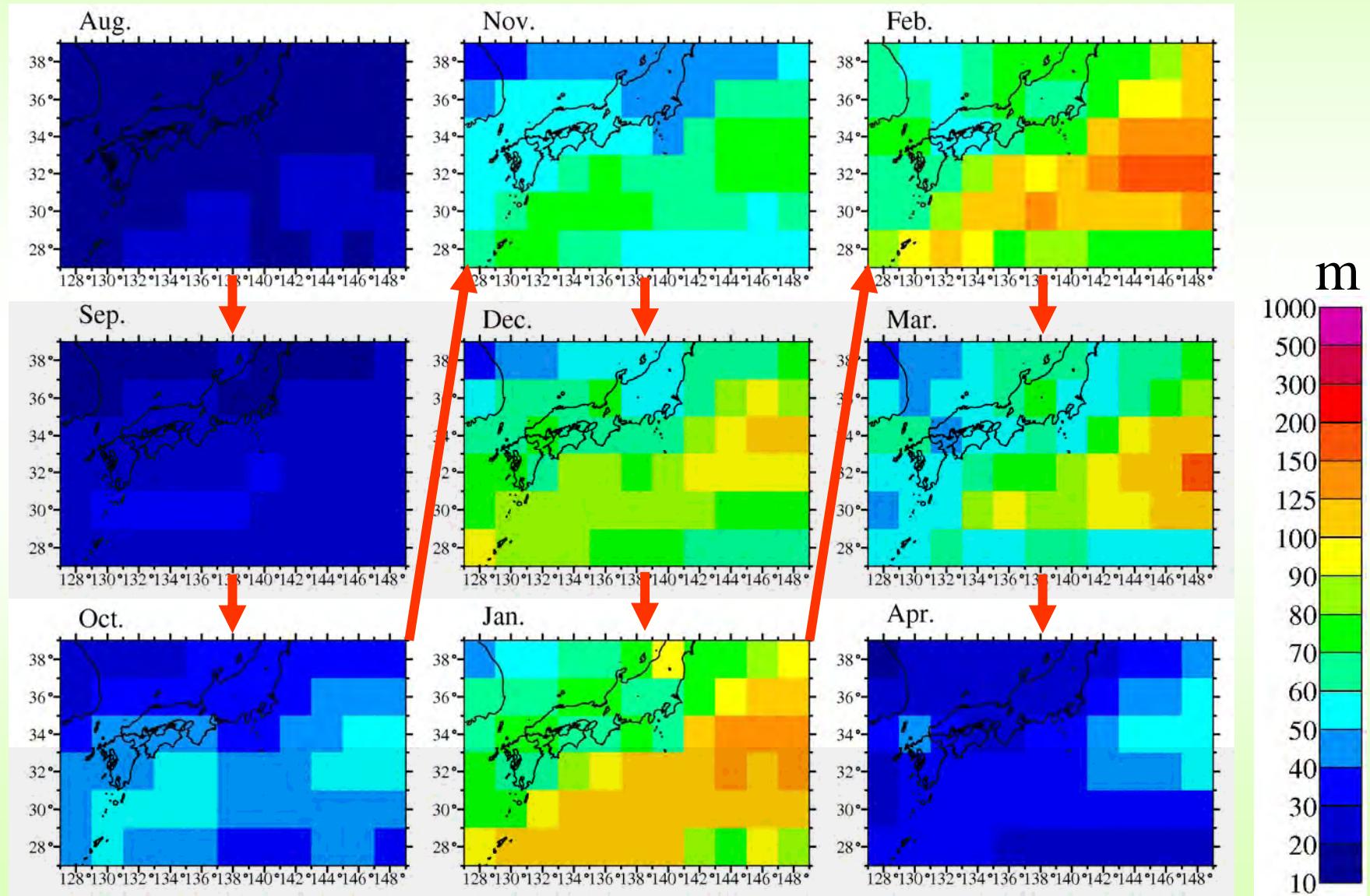
+Oceanographic condition

Developed winter mixed layer

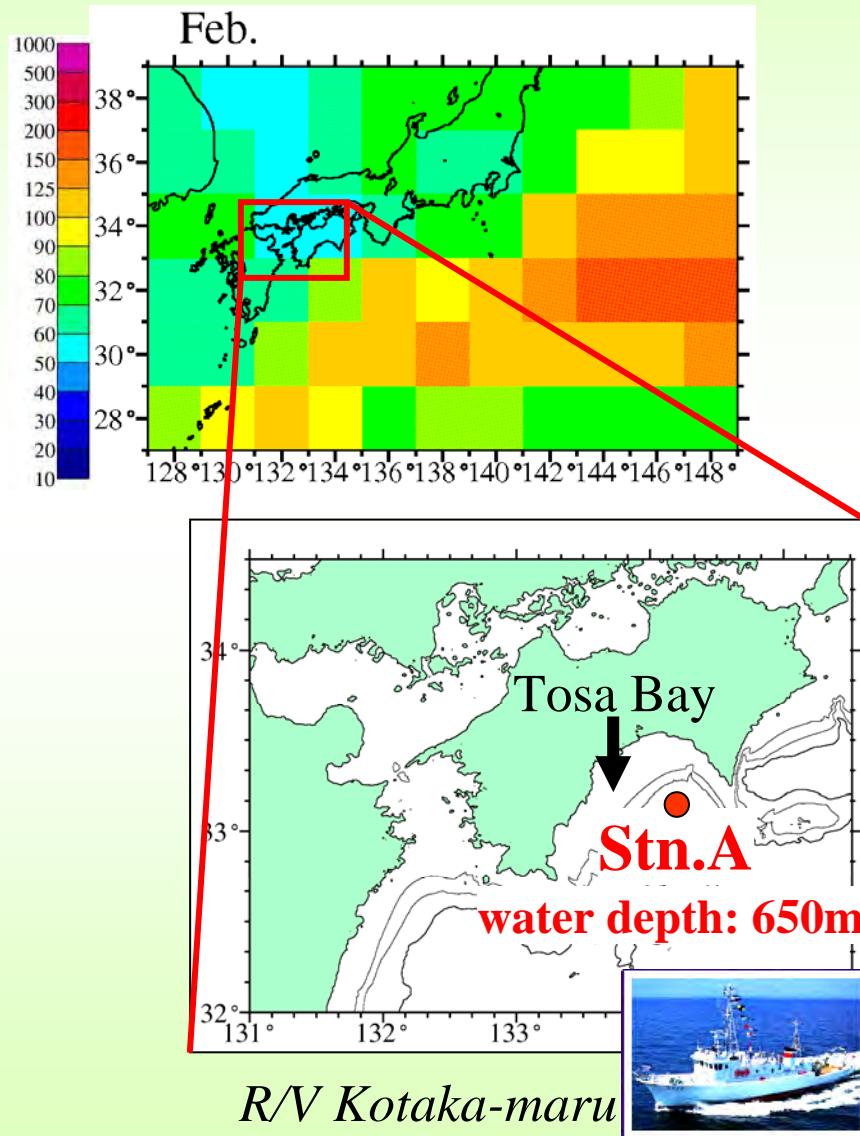
The aim of this study is  
to examine properties and dynamics of mixed layer variation  
on the basis of high-resolution model

# Climatological seasonal cycle of mixed layer depth ( $2^\circ \times 2^\circ$ )

de Boyer Montégut et al. (2004, JGR)



# Data Analysis of monthly temperature profiles on the region southwest of Japan



## Analysis Data

Station : Stn.A in Tosa Bay  
Period : 1991~2004  
Frequency : monthly after 1995  
Instrument : CTD&STD&XBT  
Vertical resolution : 1m

## Definition of Mixed Layer Depth

$$\Delta T = 0.2 \text{ } ^\circ\text{C}$$

The same criterion is used in  
de Boyer Montégut et al. (2004)

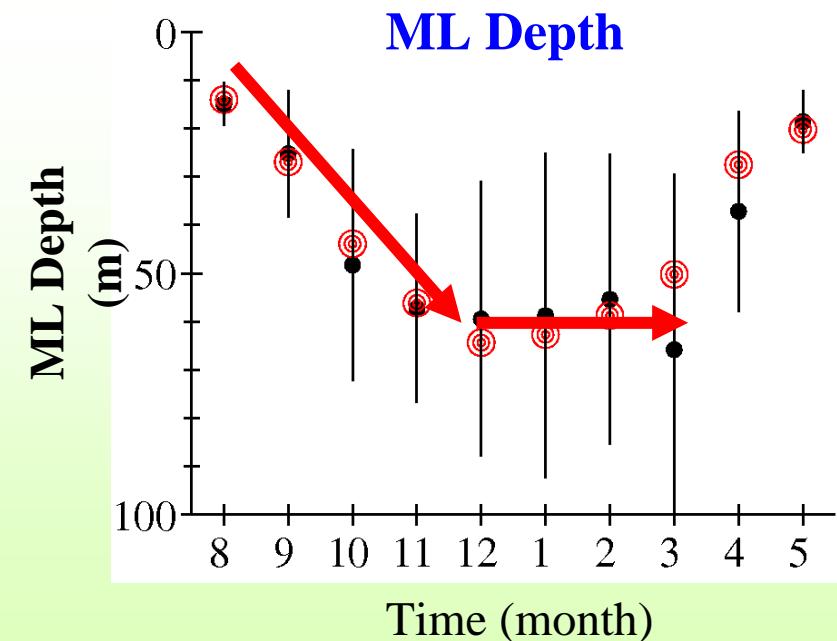
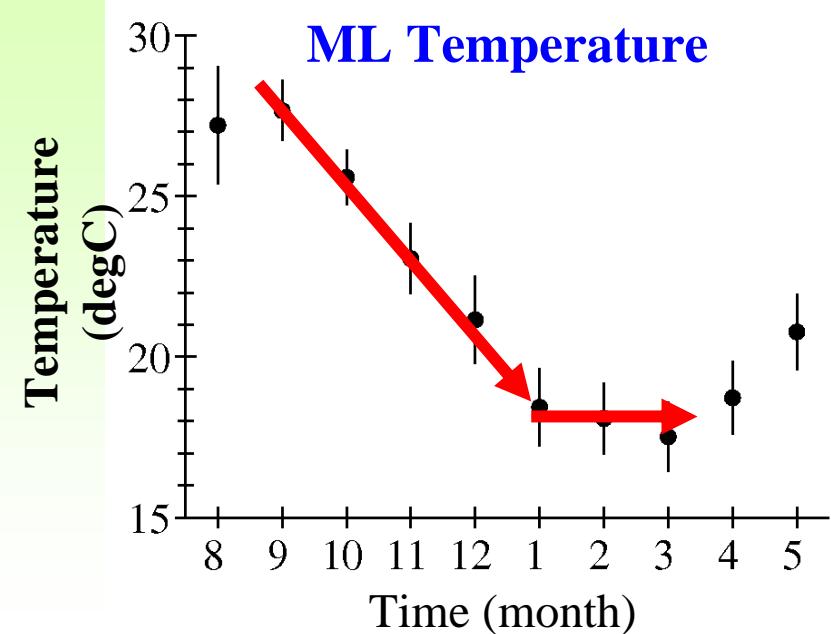
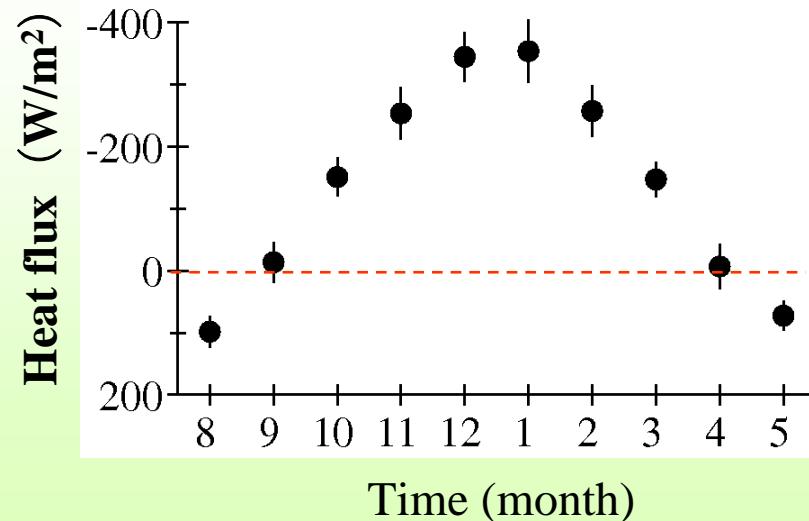
## Monthly ML temperature & depth at Stn.A in Tosa Bay

● : monthly mean  
vertical bar :  $\pm 1$  S.D.

○ : monthly climatology  
de Boyer Montégut *et al.* (2004, JGR)

## Net Heat Flux at the sea surface

estimated via bulk formulation (COARE)  
using JRA25+GHRSSST



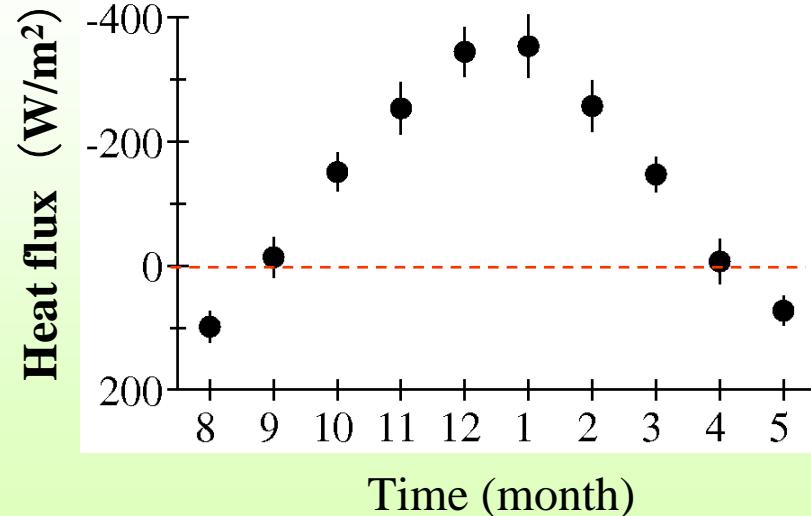
## Monthly ML temperature & depth at Stn.A in Tosa Bay

● : monthly mean  
vertical bar : 1 S.D.

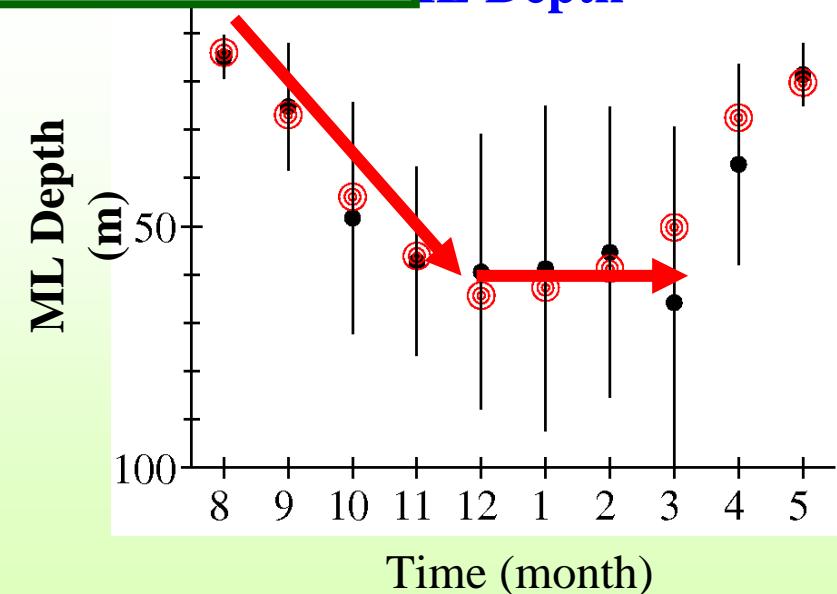
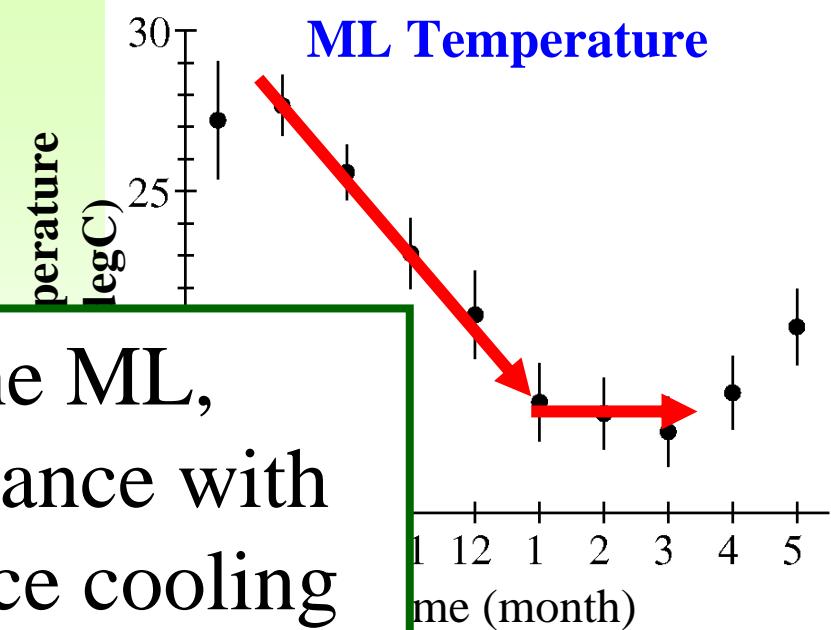
○ : clim.  
de Boyer Monté

### Net Heat Flux

estimated via bulk  
using JR



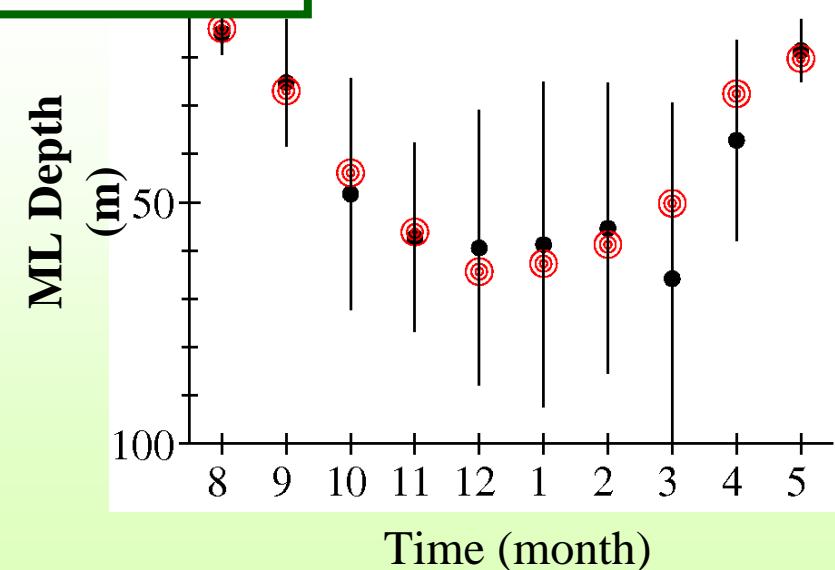
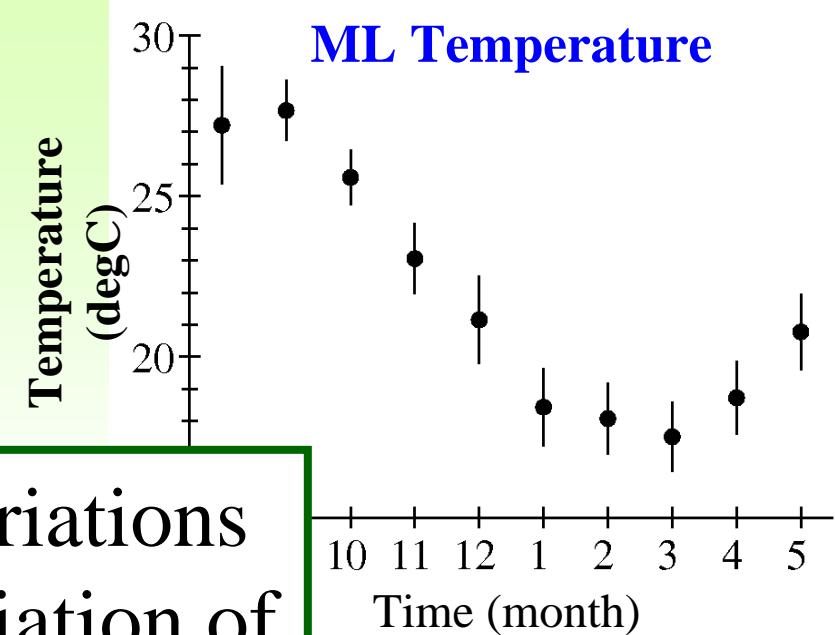
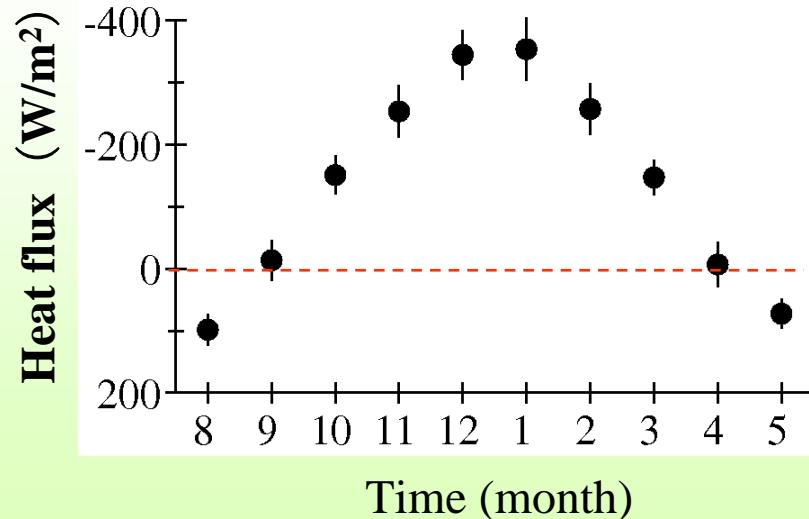
Within the ML,  
what can balance with  
the sea surface cooling  
from Jan. to Mar.?



## Monthly ML temperature & depth at Stn.A in Tosa Bay

● : monthly mean  
vertical bar :  $\pm 1$  S.D.  
○ : climatology

Can intra-monthly variations  
affect the monthly variation of  
ML from Jan. to Mar.?

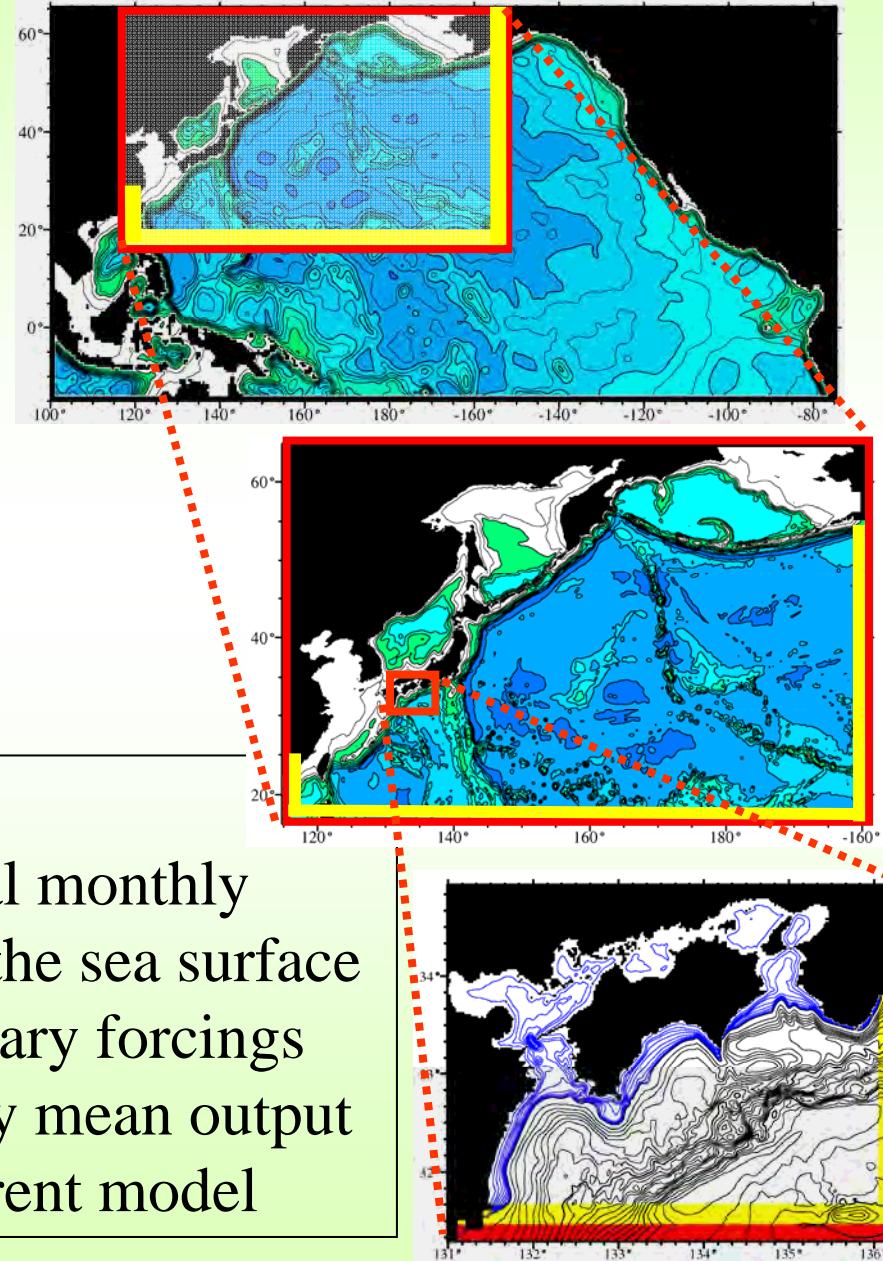


# Dynamical downscaling system

ROMS  
Regional  
Ocean  
Modeling  
System

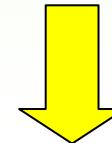
## Forcings

1. Climatological monthly mean flux at the sea surface
2. Lateral boundary forcings based on daily mean output from each parent model

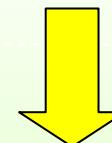


one-way  
nesting system

1/2 degree  
Basin-scale  
( $O(10^3)$  km)

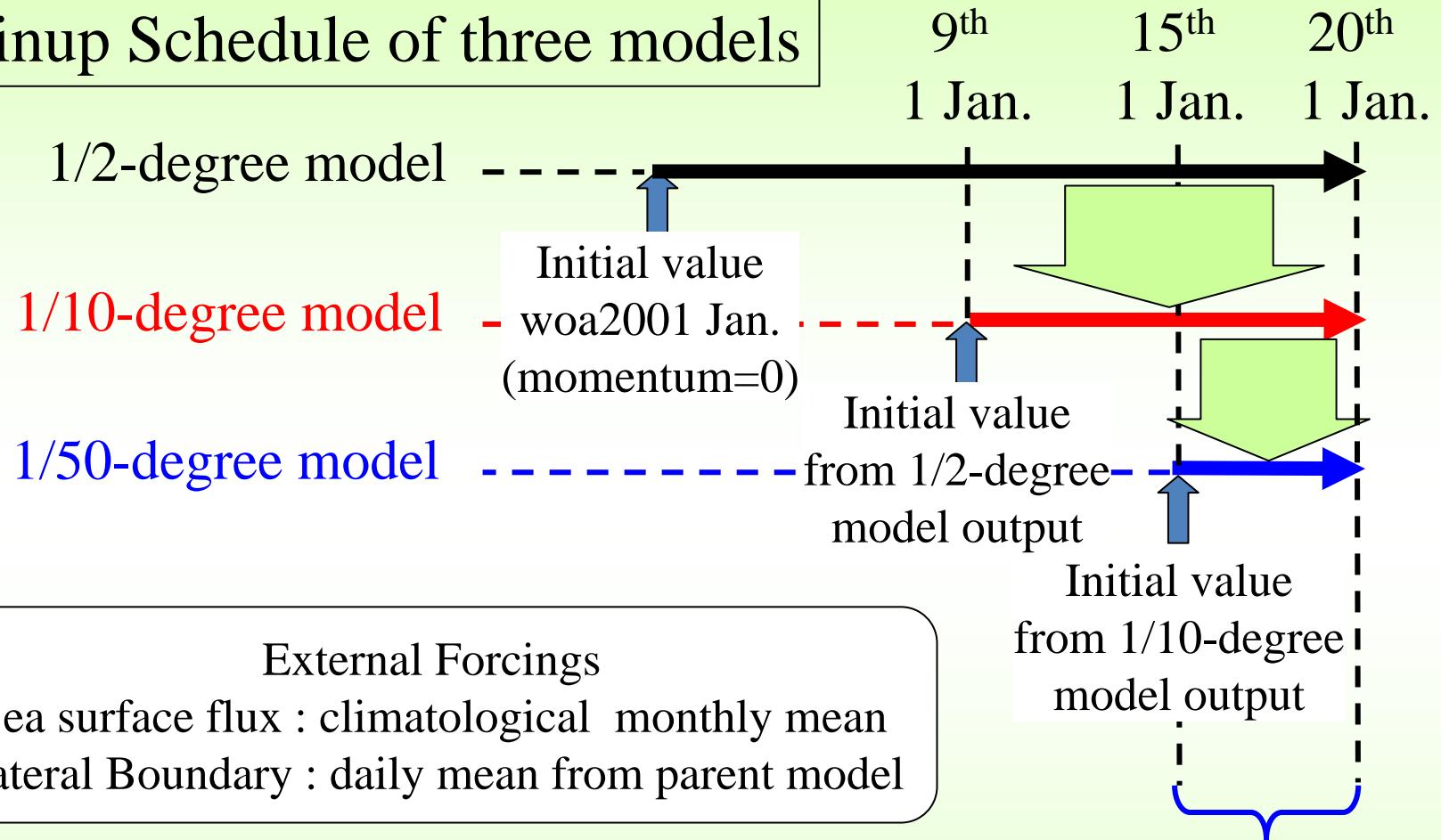


1/10 degree  
Mesoscale  
( $O(10^2)$  km)



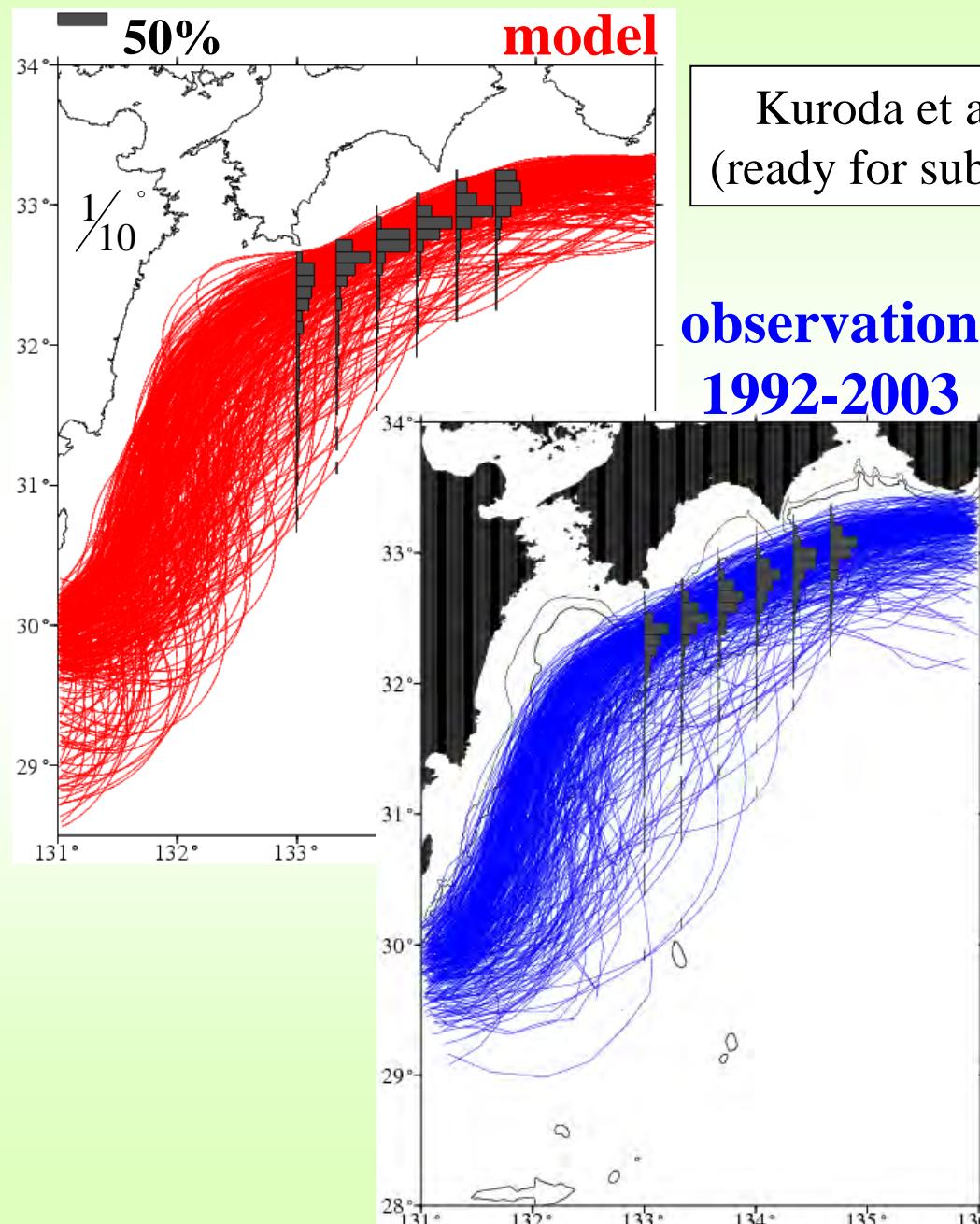
1/50 degree  
Submesoscale  
( $O(10^1)$  km)

## Spinup Schedule of three models

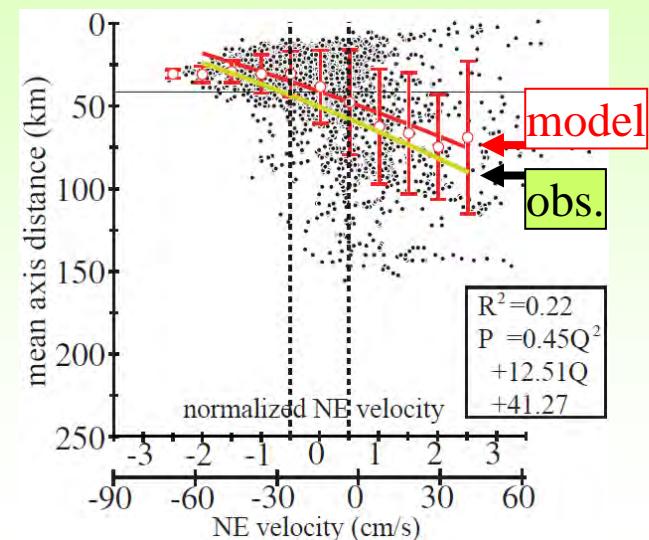


Total: 5-year simulation  
Analysis: 2<sup>nd</sup> year to 5<sup>th</sup> year

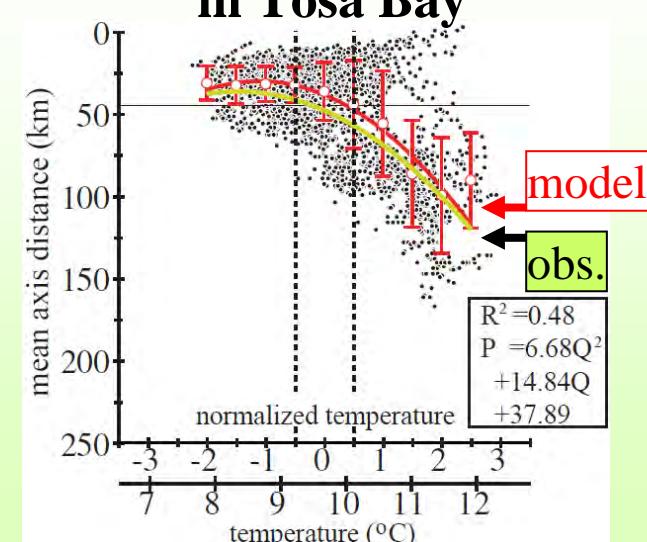
## Kuroshio axis position



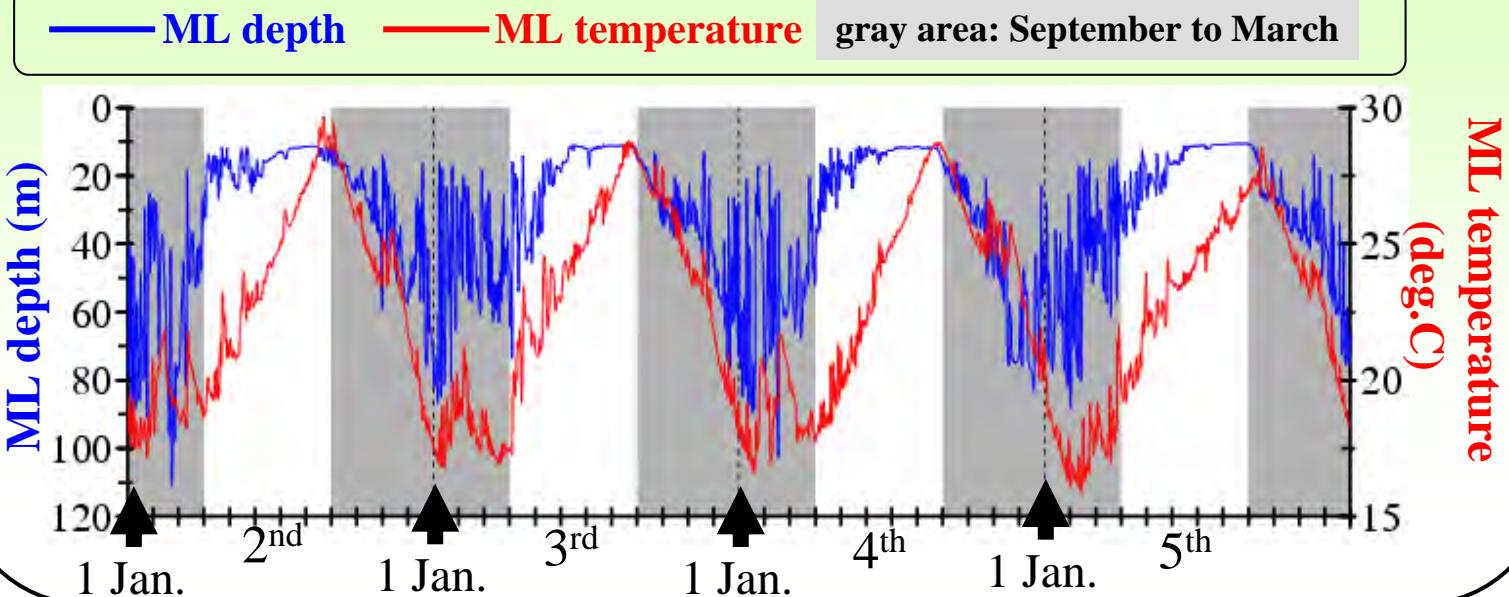
## 1/50 Kuroshio axis distance vs. Surface current in Tosa Bay



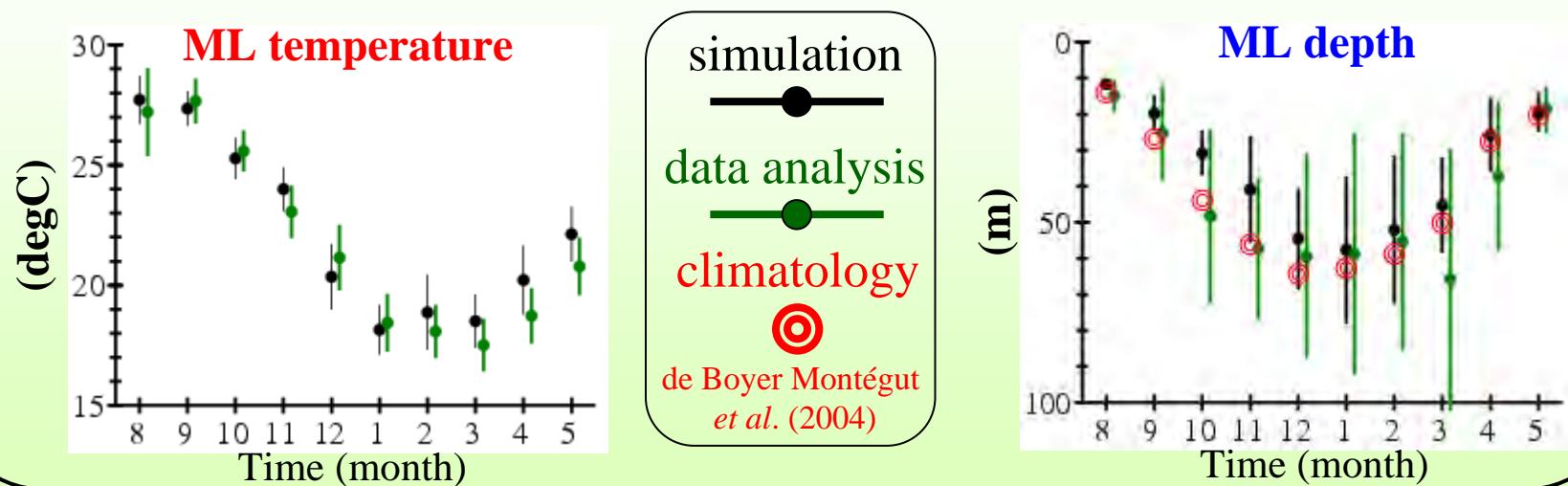
## 1/50 Kuroshio axis distance vs. Subsurface temperature in Tosa Bay



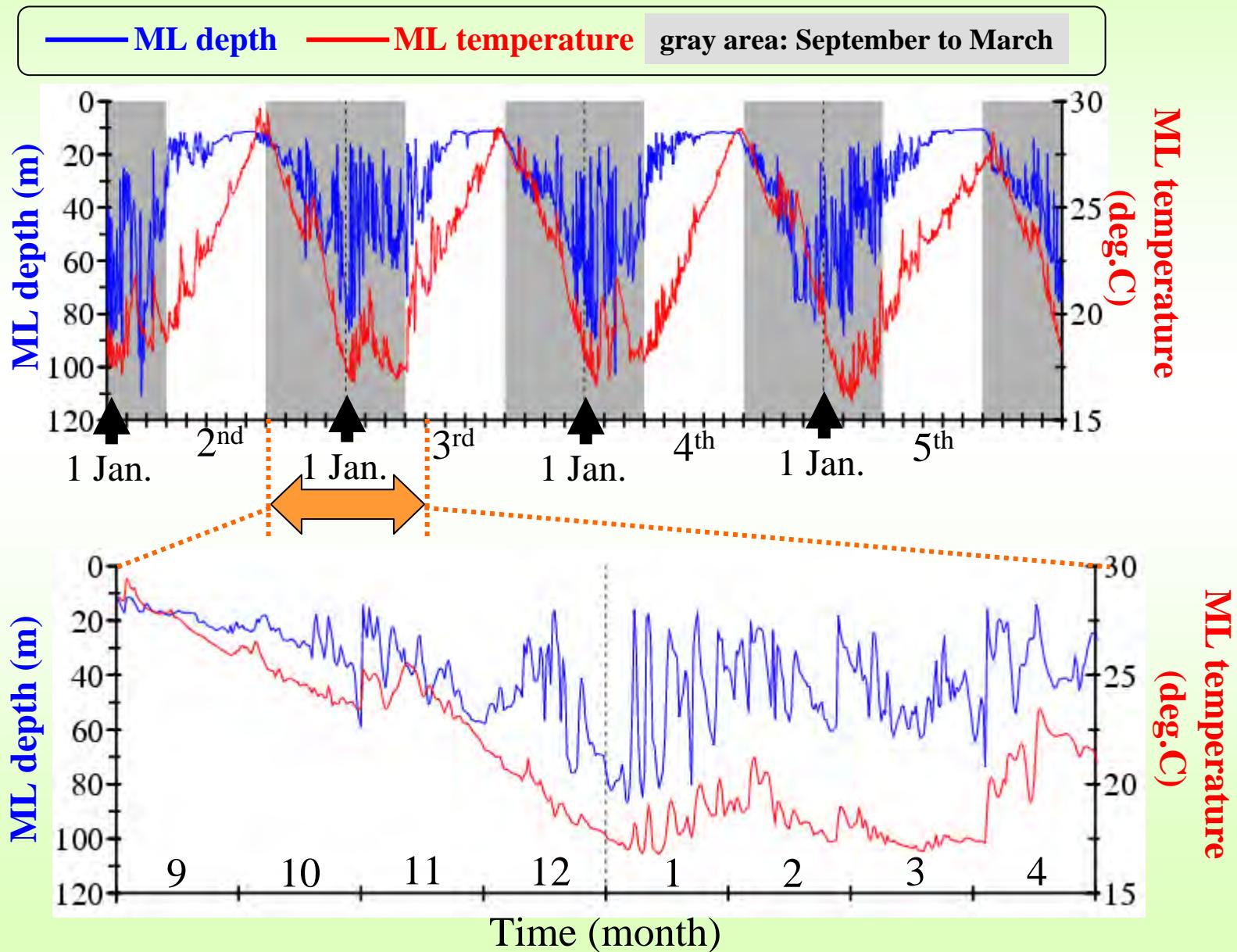
**12-hourly** time series of mixed layer at Stn.A in Tosa Bay  
from the 2<sup>nd</sup> to 5<sup>th</sup> year

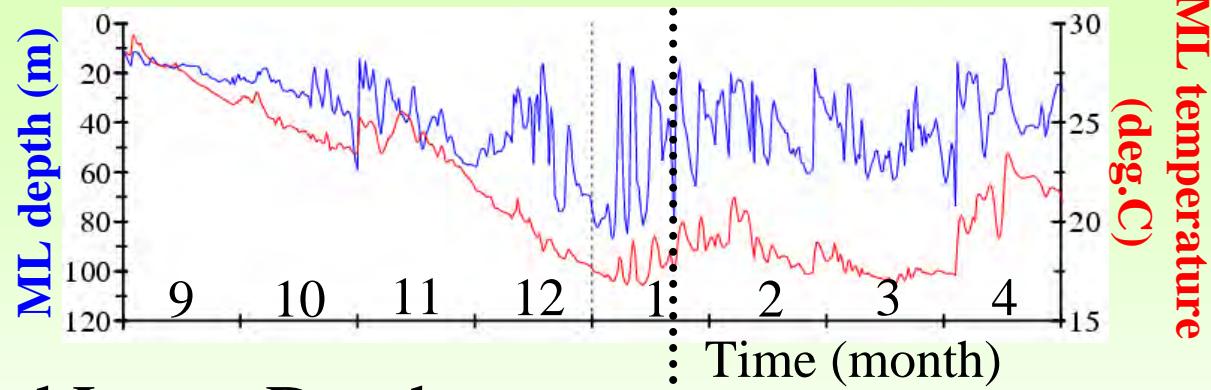


**Monthly mean** time series of mixed layer at Stn.A in Tosa Bay

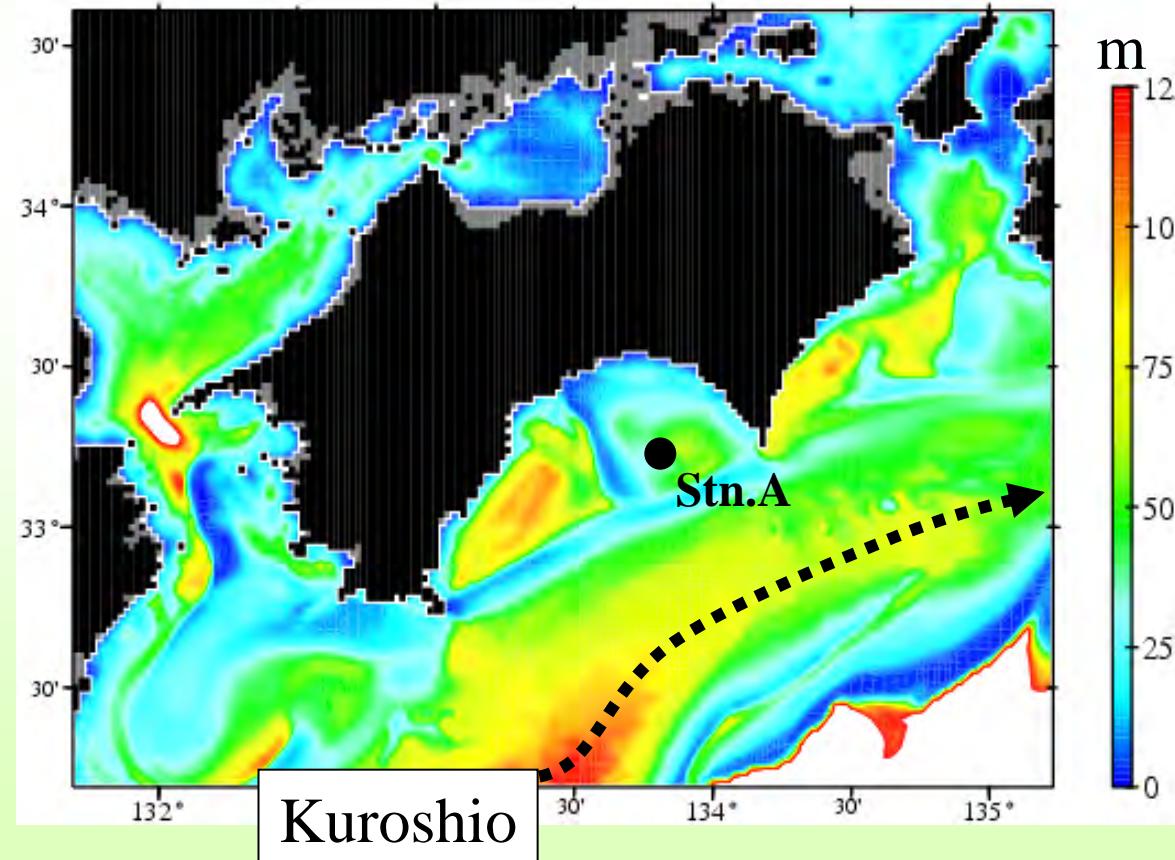


**12-hourly time series of mixed layer at Stn.A in Tosa Bay  
from the 2<sup>nd</sup> to 5<sup>th</sup> year**





Snapshot of Mixed Layer Depth  
20 January in the 3<sup>rd</sup> year



Spatial features

Patch-like deeper ML  
Streak-like shallower ML  
along the Kuroshio front

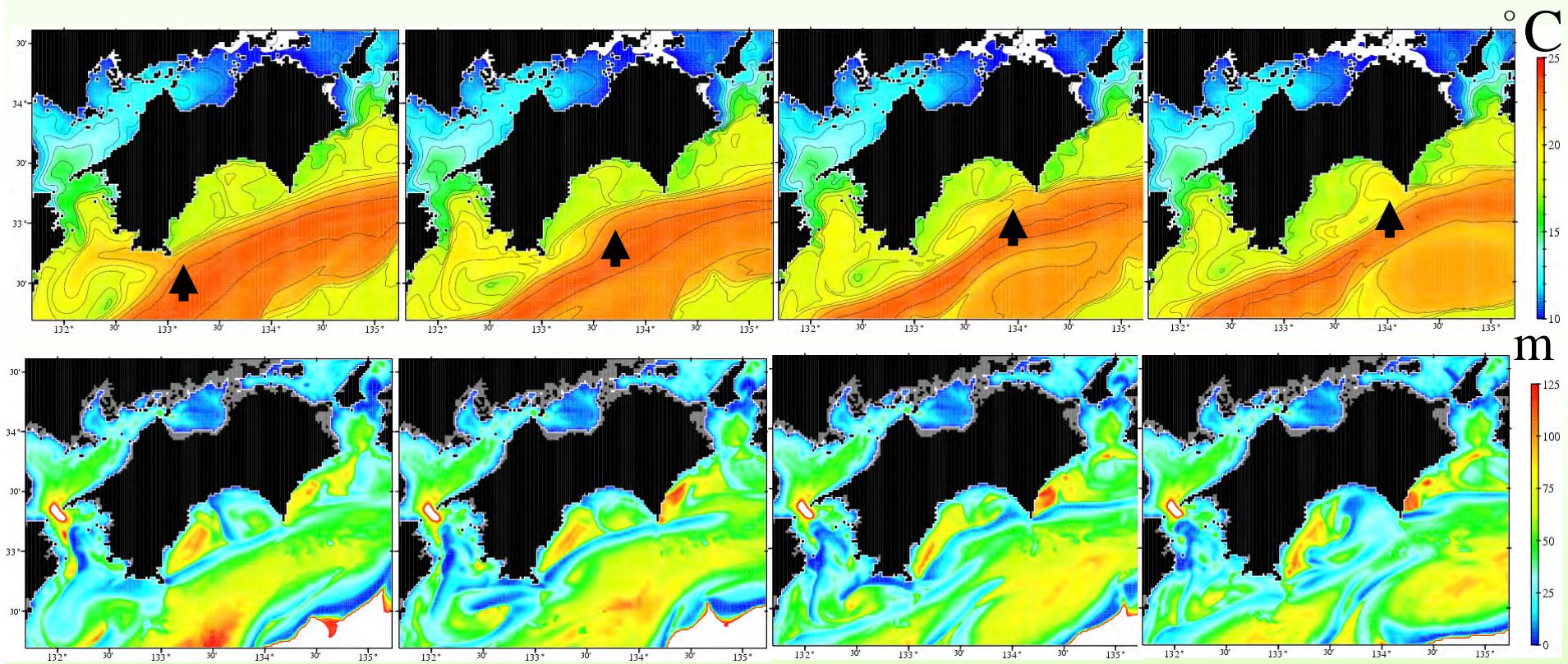
Daily snapshots of  
SST (upper panels) and ML depth (lower panels)

20 Jan.

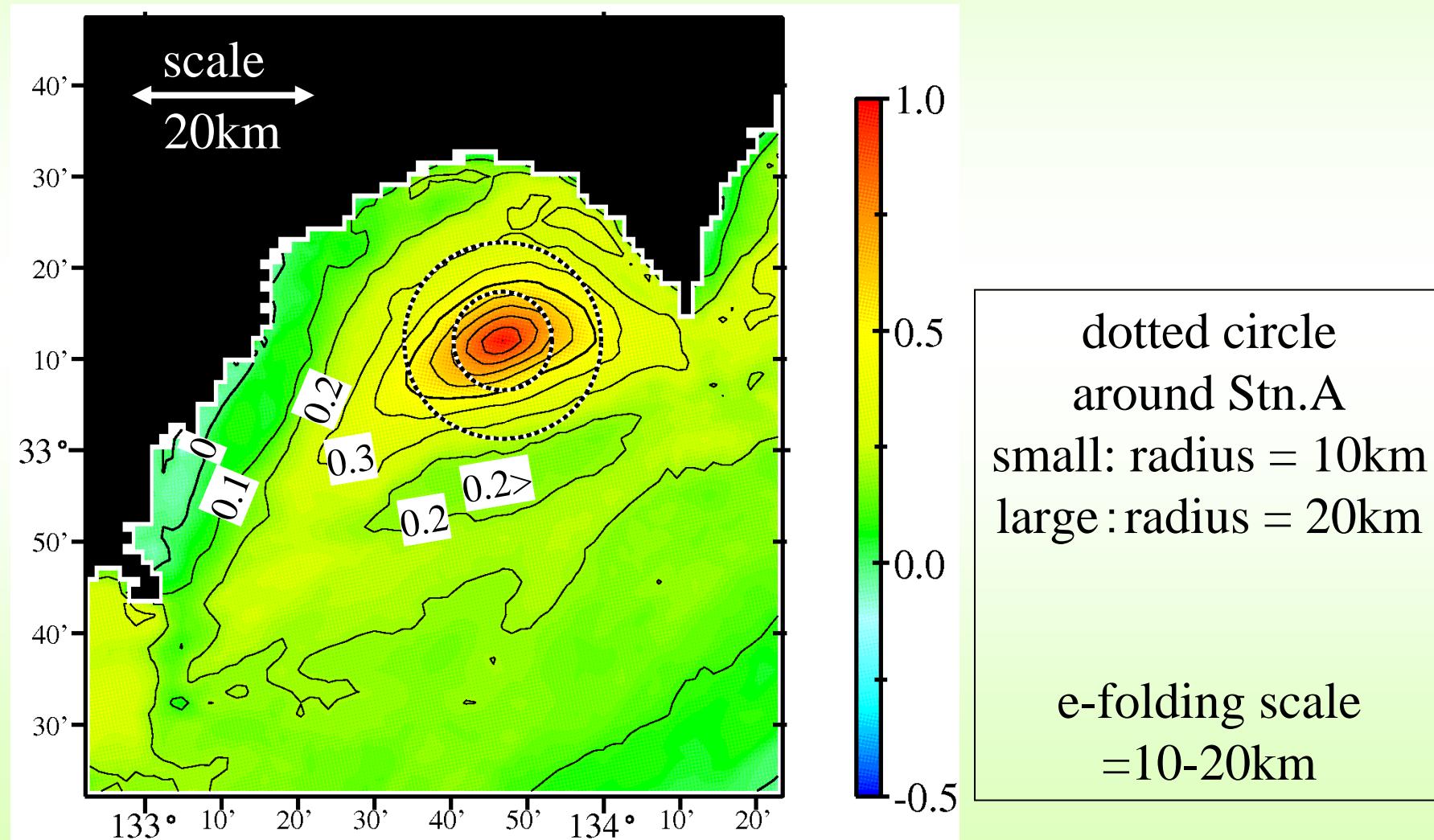
21 Jan.

22 Jan.

23 Jan.



# Correlation coefficient of mixed layer depth for Stn.A from January to March



# Heat balance of mixed layer temperature at a fixed station (Stn.A)

$$\frac{\partial T_{\text{ML}}}{\partial t} = \left[ \frac{1}{\rho C_p} \frac{\partial q}{\partial z} \right] + \frac{1}{h} (\kappa_z \nabla_z T)_{z=0} - \frac{1}{h} (\kappa_z \nabla_z T)_{z=-h}$$

shortwave  
penetration
longwave  
sensible and  
latent heat fluxes
vertical  
diffusion  
at ML bottom

**Heat flux through the sea surface**

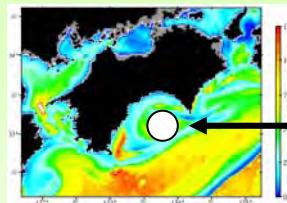
$$-\frac{1}{h} \int_{-h}^0 (\nabla \cdot \mathbf{u} T) dz + \frac{1}{h} (\nabla_H (\kappa_H \nabla_H T)) - \frac{1}{h} \frac{\partial h}{\partial t} \Delta T$$

advection
horizontal  
diffusion
entrainment

$$T_{\text{ML}} = \frac{1}{h} \int_{-h}^0 T dz$$

**Each term is diagnostically estimated from model output  
by a method of Kim et al. (2006, JPO)**

# Term estimation of the mixed-layer-temperature equation on the slope of Tosa Bay



## Part 1: Monthly Mean

$$\frac{\partial T_{\text{ML}}}{\partial t} = \left[ \frac{1}{\rho C_p} \frac{\partial q}{\partial z} \right] + \frac{1}{h} (\kappa_z \nabla_z T)_{z=0} - \frac{1}{h} (\kappa_z \nabla_z T)_{z=-h} - \frac{1}{h} \int_{-h}^0 (\nabla \cdot \mathbf{u} T) dz + \frac{1}{h} (\nabla_h (\kappa_h \nabla_h T)) - \frac{1}{h} \frac{\partial h}{\partial t} \Delta T$$

1.            2.            3            4.            5.            6.

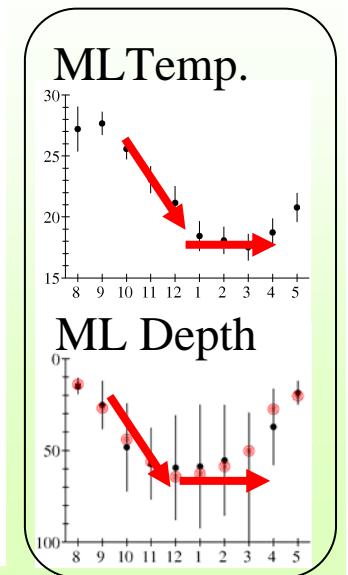
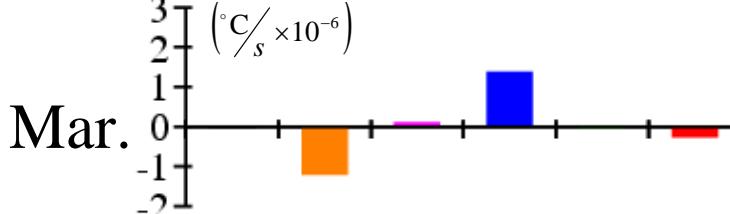
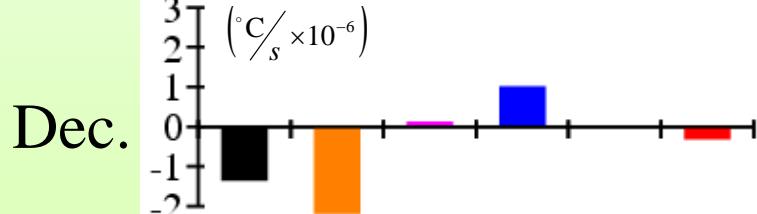
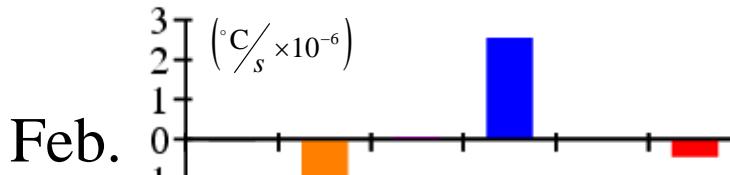
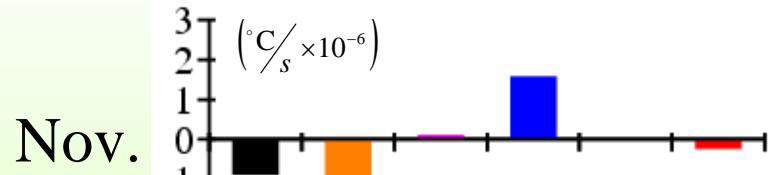
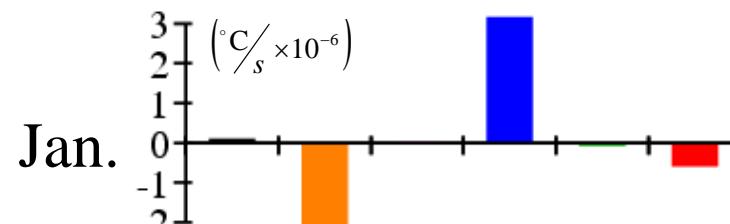
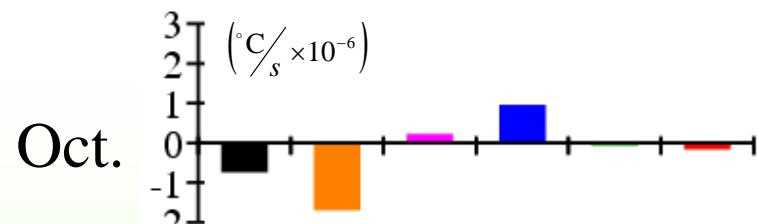
Tendency      Heat flux through  
the sea surface

Vertical  
diffusion

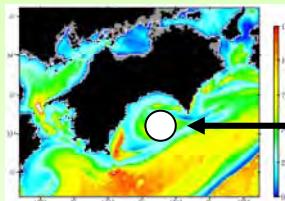
Advection

Horizontal  
diffusion

Entrainment



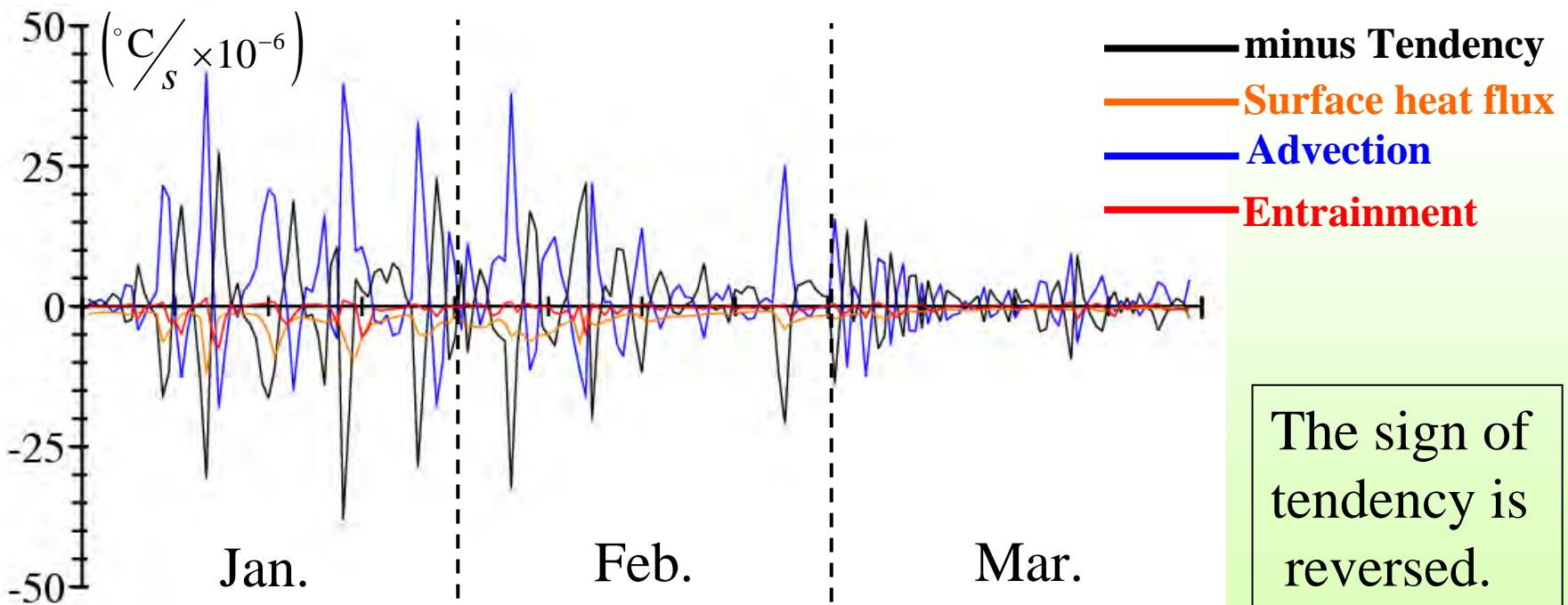
# Term estimation of the mixed-layer-temperature equation on the slope of Tosa Bay



## Part 2: 12-hour mean

$$\frac{\partial T_{ML}}{\partial t} \approx \underbrace{\left[ \frac{1}{\rho C_p} \frac{\partial q}{\partial z} \right]}_{\text{Tendency}} + \underbrace{\frac{1}{h} (\kappa_z \nabla_z T)_{z=0}}_{\text{Heat flux through the sea surface}} - \underbrace{\frac{1}{h} \int_{-h}^0 (\nabla \cdot \mathbf{u} T) dz}_{\text{Advection}} - \underbrace{\frac{1}{h} \frac{\partial h}{\partial t} \Delta T}_{\text{Entrainment}}$$

12-hour time series from January to March in the 3rd year



# Decomposition of monthly mean advection

Original term

$$-\frac{1}{h} \int_{-h}^0 (\nabla \cdot \mathbf{u} T) dz$$

$\mathbf{u} = (u, v, w)$

$$\mathbf{u} = \bar{\mathbf{u}} + \mathbf{u}'_L + \mathbf{u}'_H$$

$$T = \bar{T} + T'_L + T'_H$$

Average

High Frequency

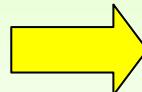
Low Frequency (Intra-monthly)  
(>30-day period)

$$\begin{aligned}
 & -\frac{1}{h} \int_{-h}^0 (\nabla \cdot \mathbf{u} T) dz \\
 &= -\frac{1}{h} \int_{-h}^0 (\nabla \cdot \bar{\mathbf{u}} \bar{T}) dz - \frac{1}{h} \int_{-h}^0 (\nabla \cdot \bar{\mathbf{u}} T'_L) dz - \frac{1}{h} \int_{-h}^0 (\nabla \cdot \bar{\mathbf{u}} T'_H) dz \\
 &\quad - \frac{1}{h} \int_{-h}^0 (\nabla \cdot \mathbf{u}'_L \bar{T}) dz - \frac{1}{h} \int_{-h}^0 (\nabla \cdot \mathbf{u}'_L T'_L) dz - \frac{1}{h} \int_{-h}^0 (\nabla \cdot \mathbf{u}'_L T'_H) dz \\
 &\quad - \frac{1}{h} \int_{-h}^0 (\nabla \cdot \mathbf{u}'_H \bar{T}) dz - \frac{1}{h} \int_{-h}^0 (\nabla \cdot \mathbf{u}'_H T'_L) dz - \frac{1}{h} \int_{-h}^0 (\nabla \cdot \mathbf{u}'_H T'_H) dz
 \end{aligned}$$

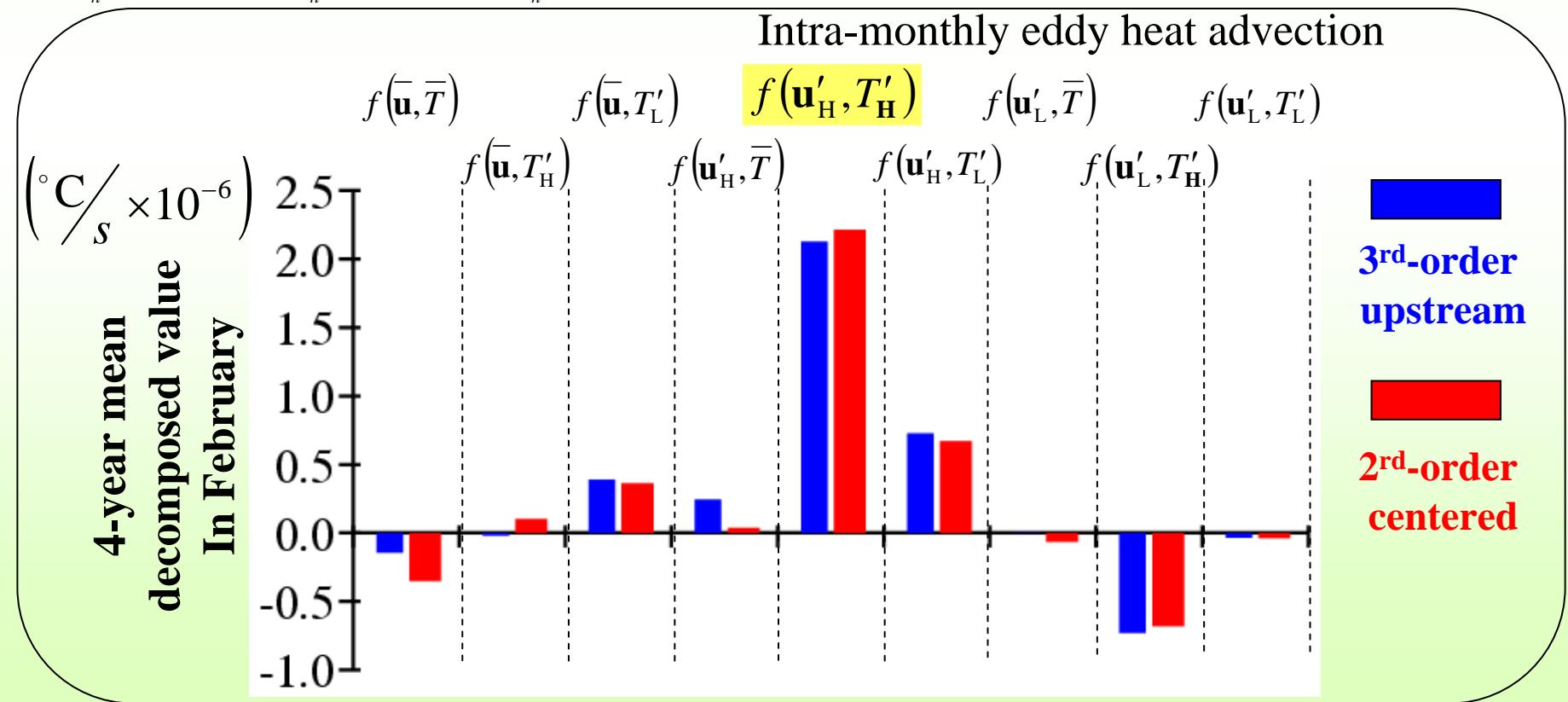
If  $h$  is constant,  
Orange Terms=0

# Contribution of decomposed components (February)

$$\begin{aligned}
 & -\frac{1}{h} \int_{-h}^0 (\nabla \cdot \mathbf{u} T) dz \\
 &= -\frac{1}{h} \int_{-h}^0 (\nabla \cdot \bar{\mathbf{u}} \bar{T}) dz - \frac{1}{h} \int_{-h}^0 (\nabla \cdot \bar{\mathbf{u}} T'_L) dz - \frac{1}{h} \int_{-h}^0 (\nabla \cdot \bar{\mathbf{u}} T'_H) dz \\
 &\quad - \frac{1}{h} \int_{-h}^0 (\nabla \cdot \mathbf{u}'_L \bar{T}) dz - \frac{1}{h} \int_{-h}^0 (\nabla \cdot \mathbf{u}'_L T'_L) dz - \frac{1}{h} \int_{-h}^0 (\nabla \cdot \mathbf{u}'_L T'_H) dz \\
 &\quad - \frac{1}{h} \int_{-h}^0 (\nabla \cdot \mathbf{u}'_H \bar{T}) dz - \frac{1}{h} \int_{-h}^0 (\nabla \cdot \mathbf{u}'_H T'_L) dz - \frac{1}{h} \int_{-h}^0 (\nabla \cdot \mathbf{u}'_H T'_H) dz
 \end{aligned}$$



$$\begin{aligned}
 & f(\mathbf{u}, T) \\
 &= f(\bar{\mathbf{u}}, \bar{T}) + f(\bar{\mathbf{u}}, T'_L) + f(\bar{\mathbf{u}}, T'_H) \\
 &\quad + f(\mathbf{u}'_L, \bar{T}) + f(\mathbf{u}'_L, T'_L) + f(\mathbf{u}'_L, T'_H) \\
 &\quad + f(\mathbf{u}'_H, \bar{T}) + f(\mathbf{u}'_H, T'_L) + f(\mathbf{u}'_H, T'_H)
 \end{aligned}$$



# Conclusions

Seasonal cycle of ML depth and temperature exhibits two regimes

1. September to November

Monthly mean ML depth increases & ML temperature decreases

2. January to March

Monthly mean ML depth and temperature become constant

For the monthly mean heat balance in the ML  
Heat flux through the sea surface + entrainment ~  
advection



eddy heat advection

There is dominant intra-monthly submesoscale variability  
associated with the Kuroshio or its frontal disturbances

Intra-monthly variation contributes significantly to the monthly/seasonal  
variation of heat balance within the ML

# Thank you for your attention

