Anomalous oceanic conditions along the US West Coast in 2014: inferences from a high-resolution regional ocean model

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/ support: NOAA WCOFS project

West Coast Ocean Forecast System (WCOFS): based on Regional Ocean Modeling System (ROMS, www. myroms.org)

Horizontal resolution: 2-km

Vertical resolution: 40 terrain-following layers

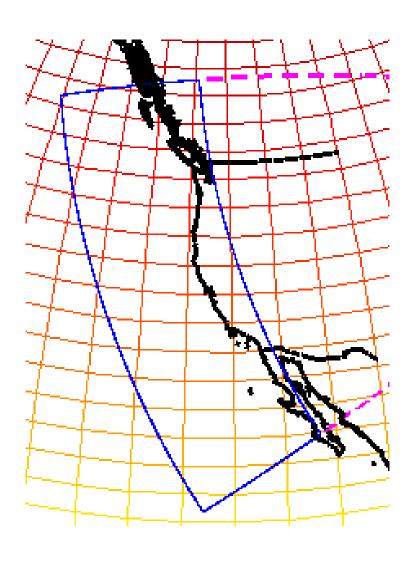
Forcing:

- Surface winds and heat flux (12-km NOAA NAM)
- @open boundary: global model (HYCOM/RTOFS)
 - + tides (Oregon State Tidal Inverse Soft.)
- River inputs: Columbia R., Fraser R., Puget Sound

Skill assessments: 6+ -year model run without data assimilation, Oct 2008 thru Dec 2014 (daily-averaged outputs)

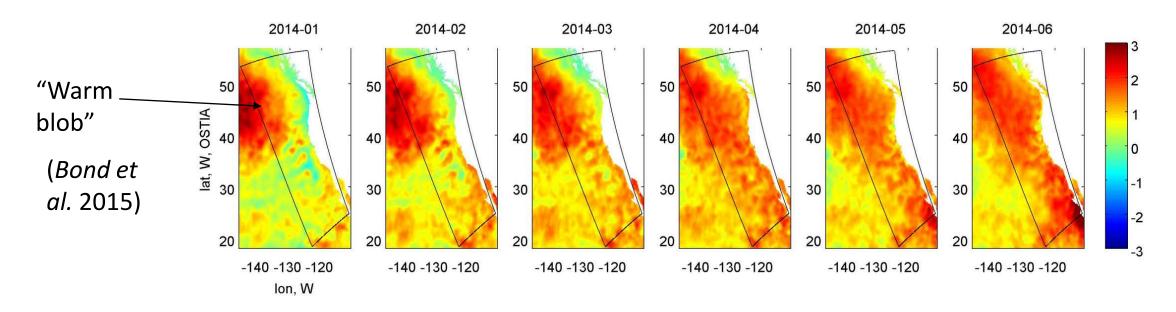
This talk:

- Anomalous warming in NEP in the beginning of 2014
- Variability along the continental slope, on isopycnal surface σ_{θ} =26.5 kg/m 3

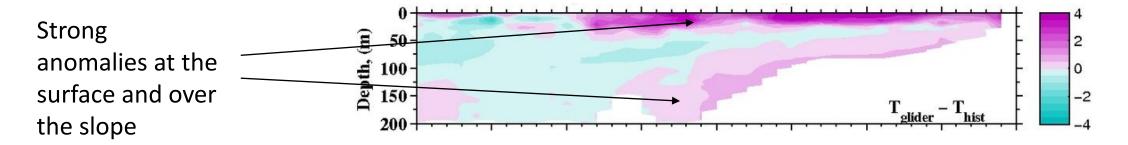


Strong temperature anomalies emerge in NEP 2014 (and these would persist until the end of 2015)

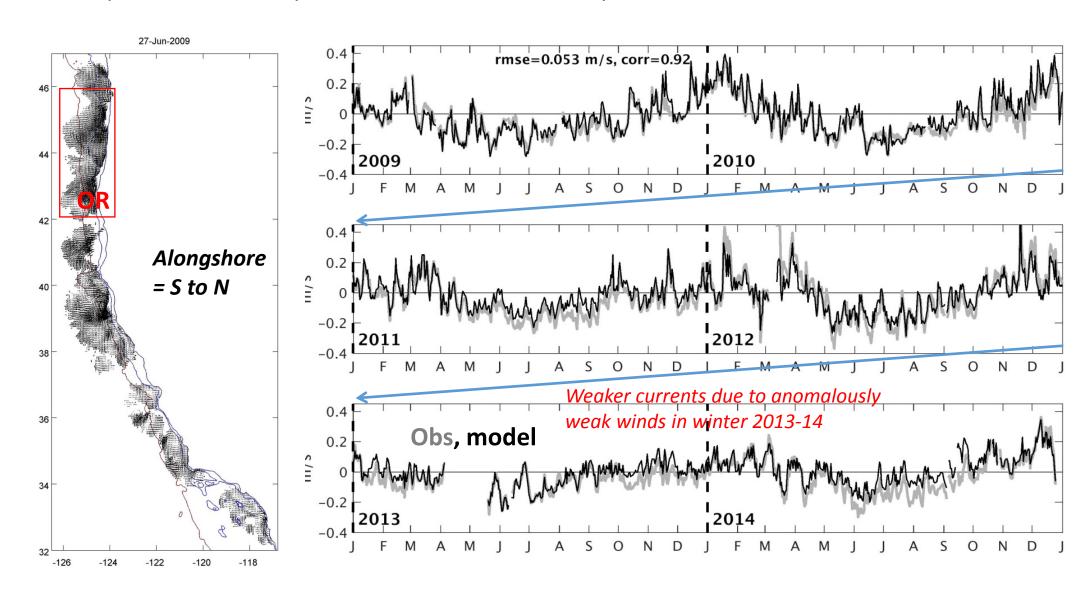
Satellite OSTIA SST anomaly, w /respect to 2009-2013 climatology (OSTIA)



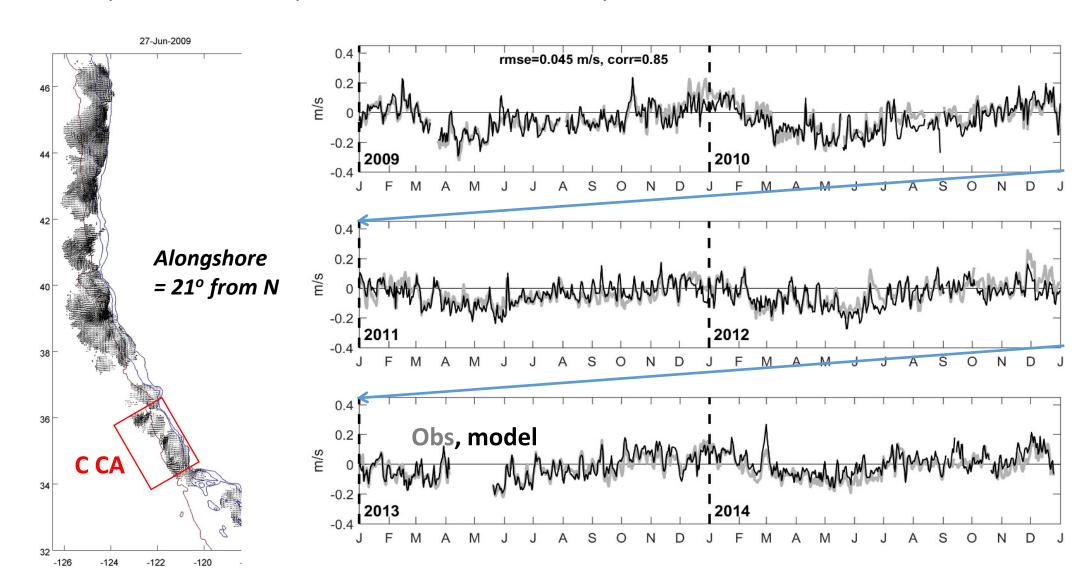
Glider temperature anomaly off Oregon, 44.6N (September 2014) [Barth, OSU]



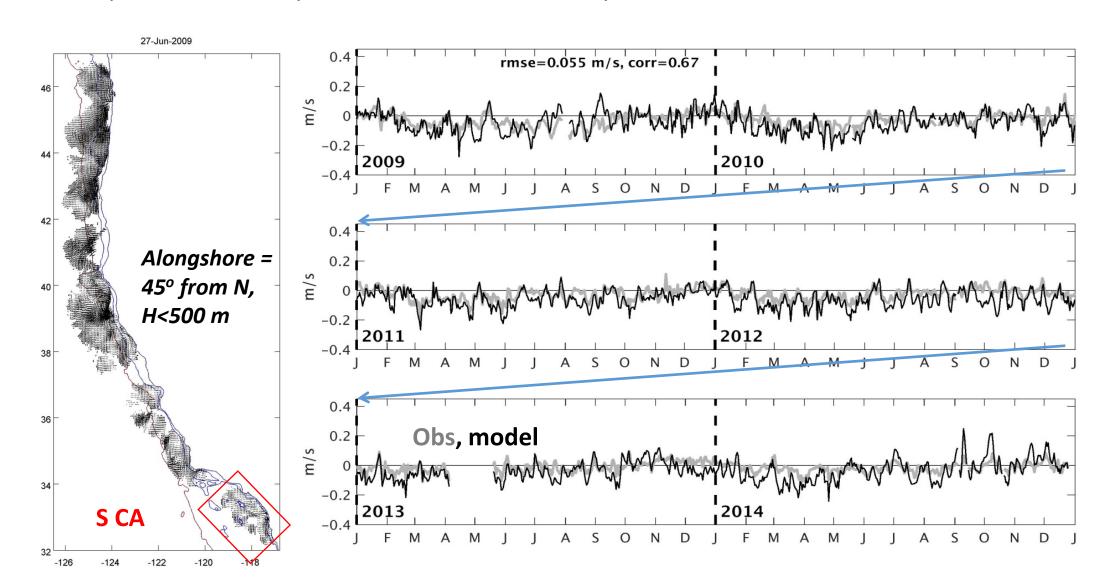
HF radar vs. WCOFS surface currents (area-averaged, daily-averaged alongshore currents... sim. to Durski et al. Oc. Dyn. 2015): variability is predicted on temporal scales from several days to seasonal and interannual



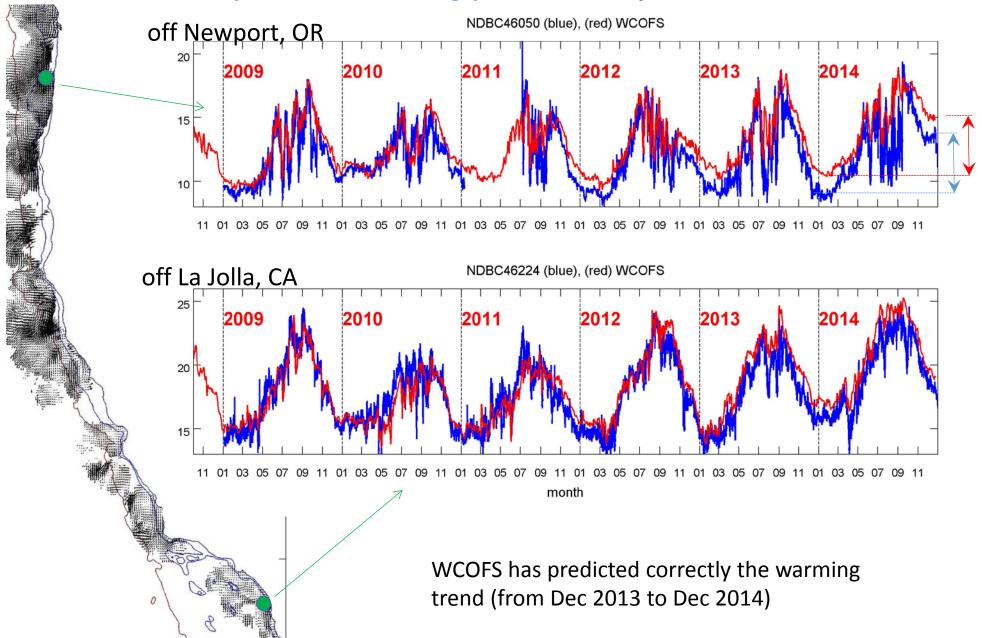
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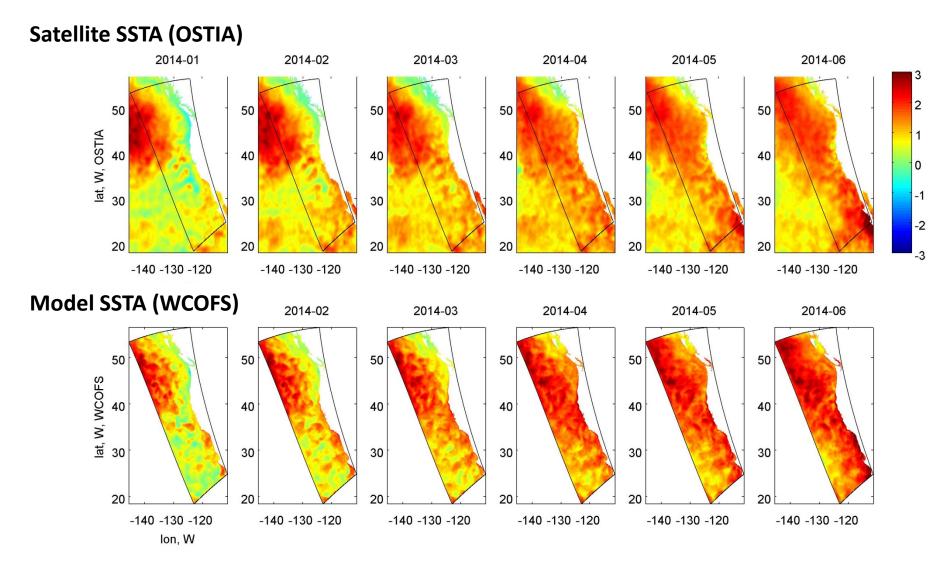
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Near-surface T (NDBC shelf moorings) / WCOFS comparison:

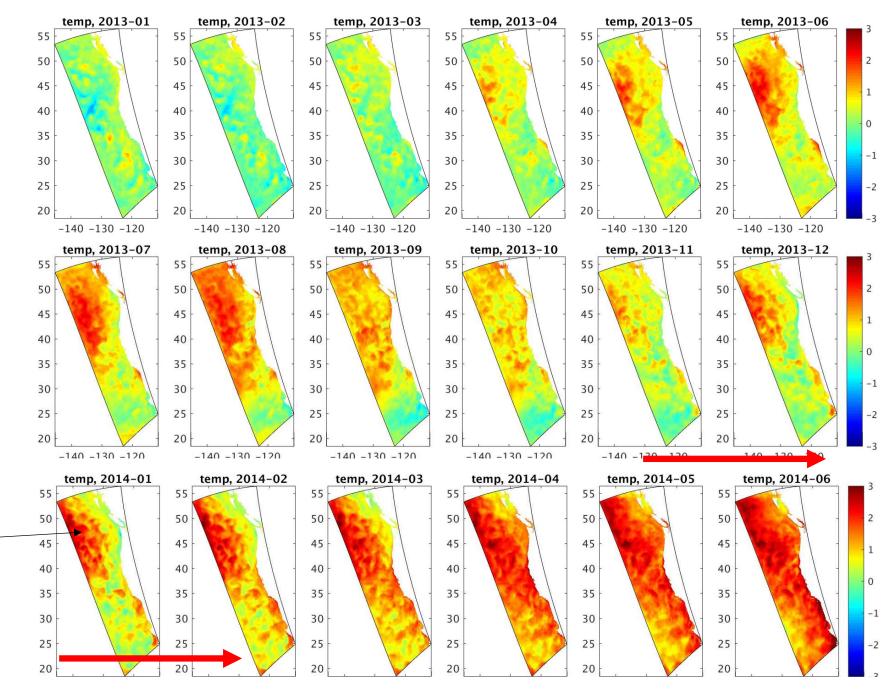


Compared to sat. SST anomaly, WCOFS predicts the appearance of the warm blob by Jan 2014, and wide-spread warming along the US Coast by summer 2014



Anomaly w/ respect to 2009-2013 climatology (computed similarly for sat. and model)

ROMS SST anomaly, Jan 2013 through Jun 2014 (w resp to 2009-2013 clim, 3-mo running ave)



Warming during Dec 2013 – Jan 2014 – Feb 2014

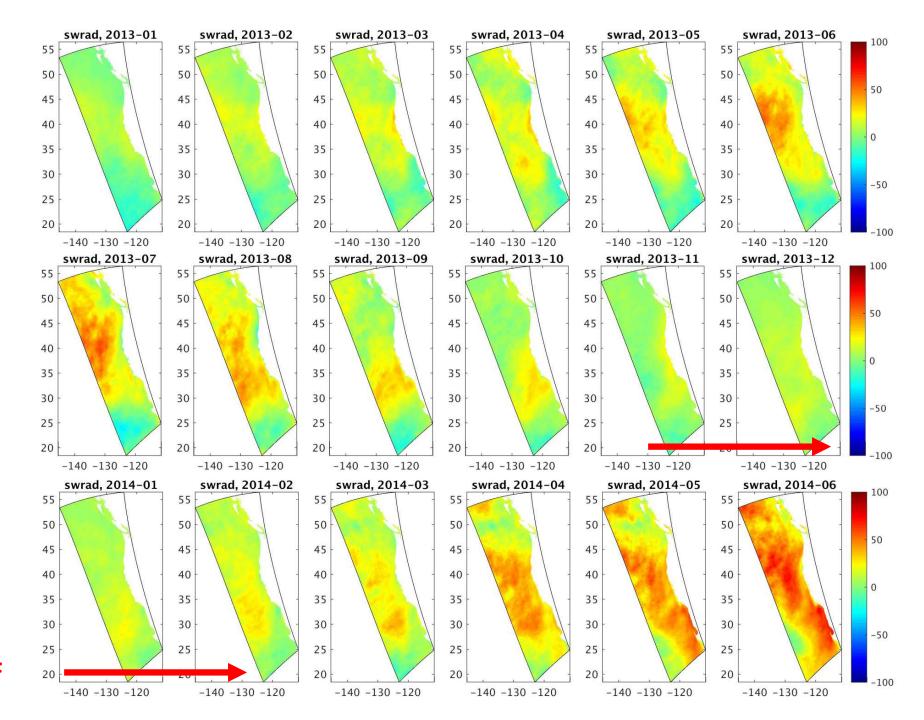
13ND-14JF

Anomaly in net shortwave radiation (SWRAD), w /respect to 2009-2013:

Fall 2013-winter 2014 does not receive anomalous SWRAD

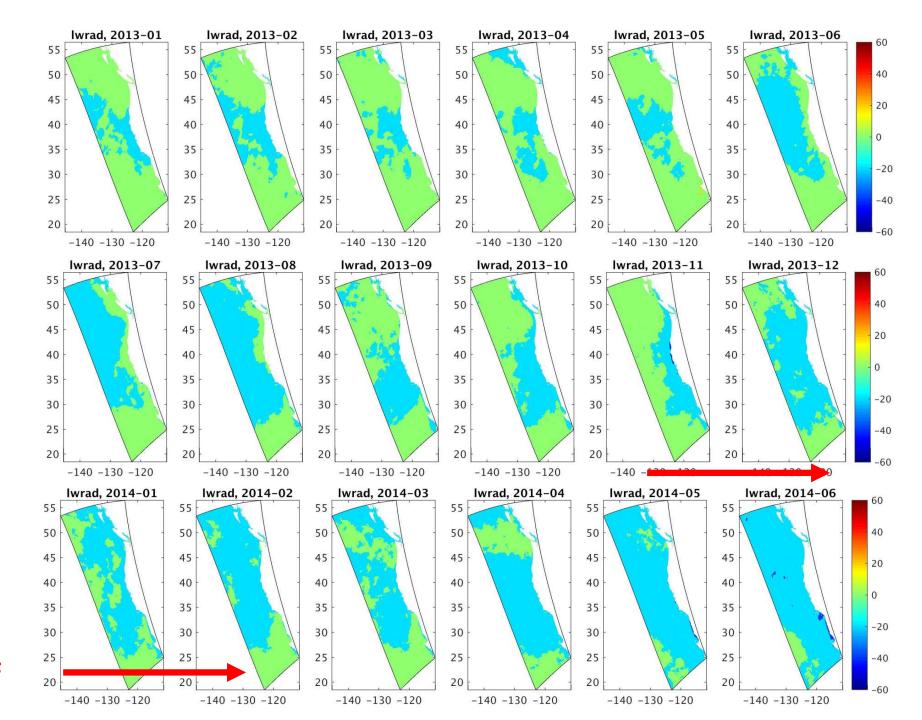
Summer 2013: 30-50 W/m2 extra

Summer 2014: SWRAD anomaly certainly contributed to warming all along the US West Coast and Mexico

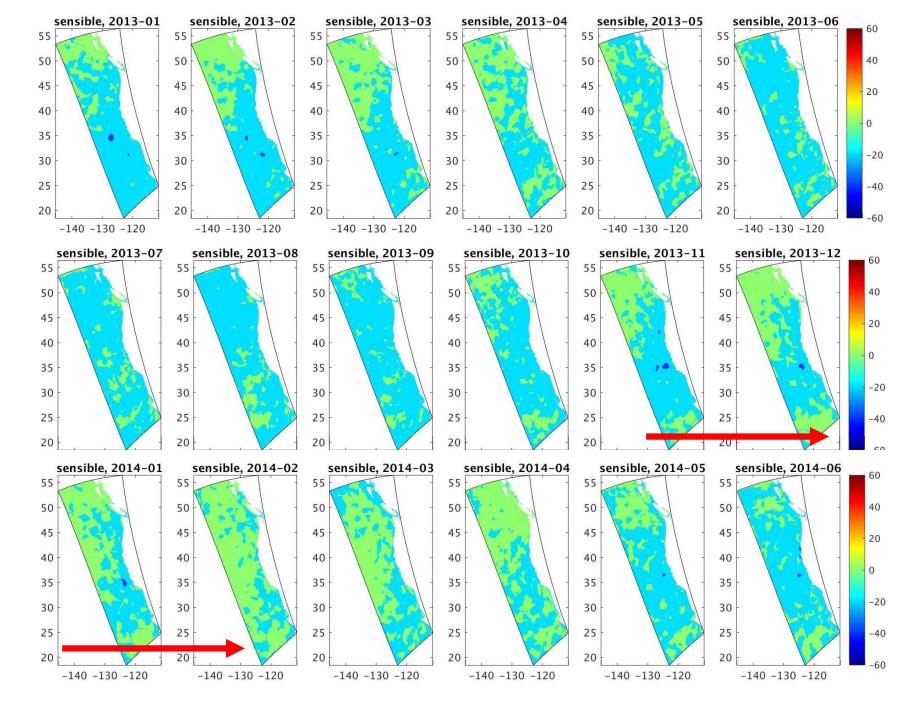


13ND-14JF

Anomaly in net longwave radiation (LWRAD), w /respect to 2009-2013:



Anomaly in sensible heat flux, w /respect 2009-2013:



13ND-14JF

Anomaly in latent he flux, w /respect to 2009-2013:

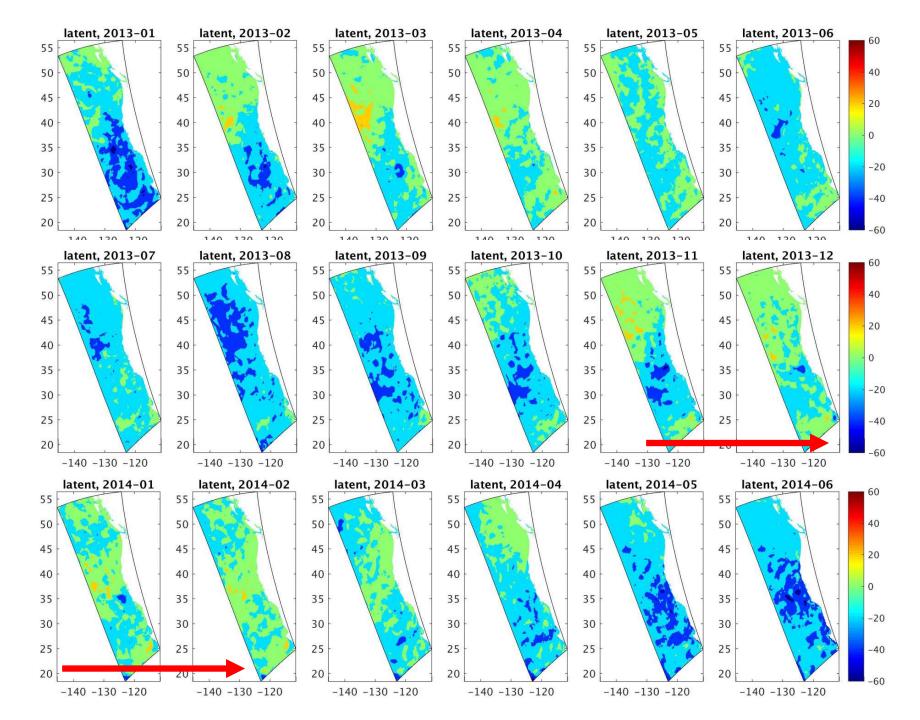
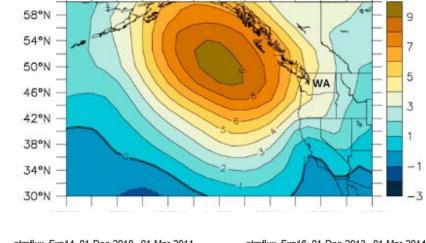
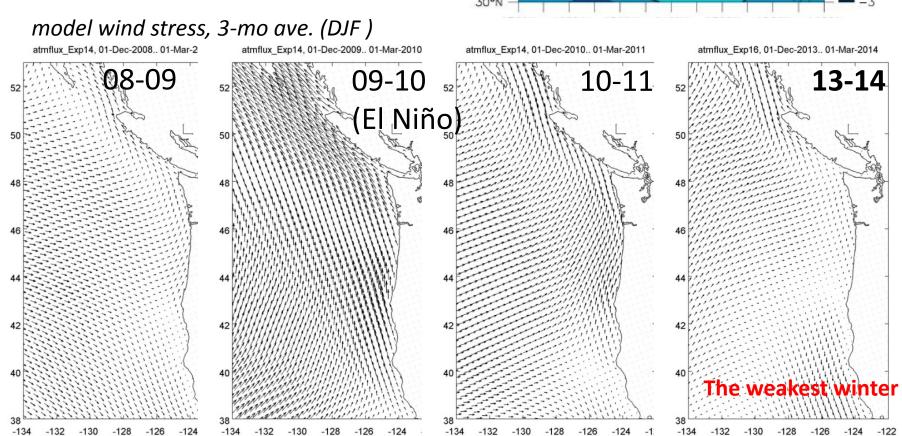


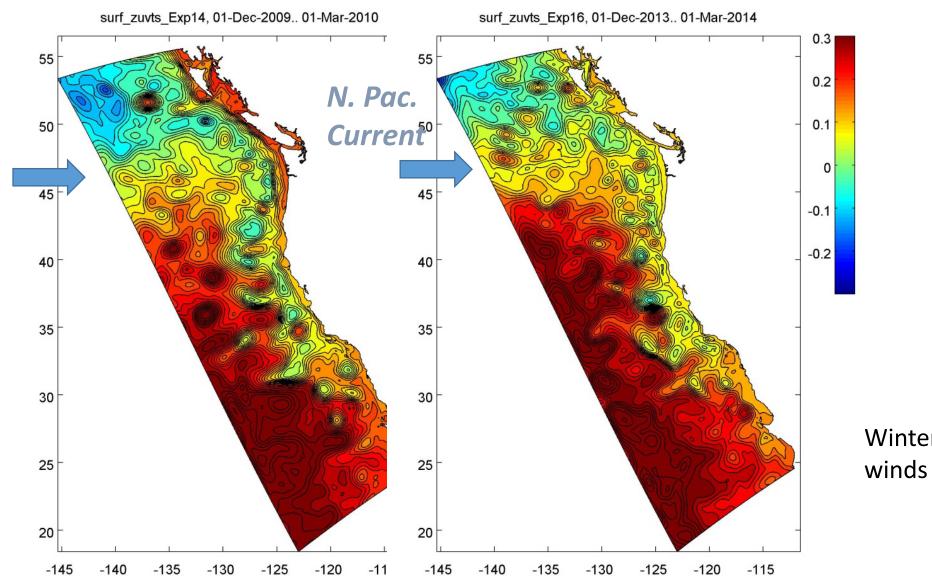
Figure 2 from Bond, Cronin,
Freeland & Mantua, GRL, 2015:
Mean Sea Level P anomaly (hPa) Oct
2013-Jan 2014, relative to 19812011 mean





DJF ave SSH: 2009-10 and 2013-14:

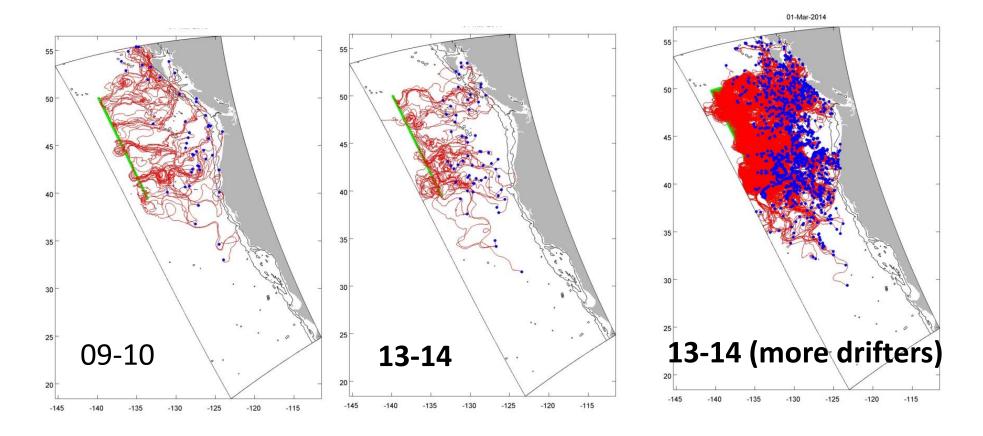
Onshore surface transport = geostrophic + Ekman



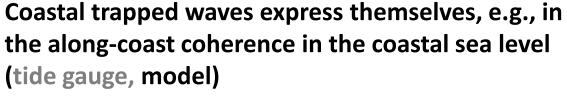
Winter w strong winds

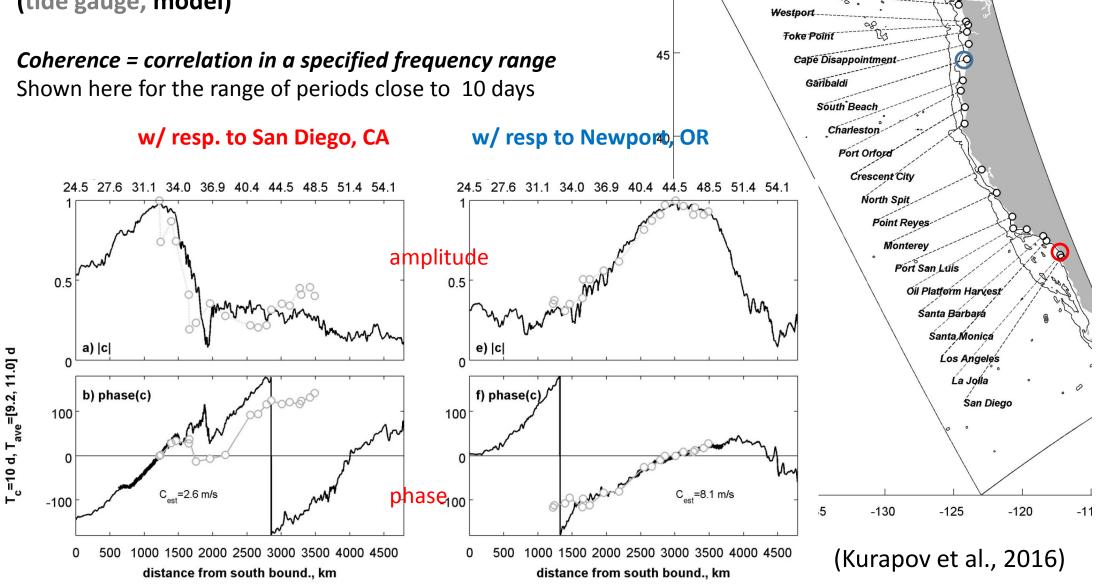
Winter w weak winds

Surface Lagrangian transport (1 Oct to 1 Mar):



Release: on the path of N. Pac. Current In 2013-14, these surface drifters avoid OR-WA (weaker onshore Ekman transport)

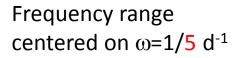


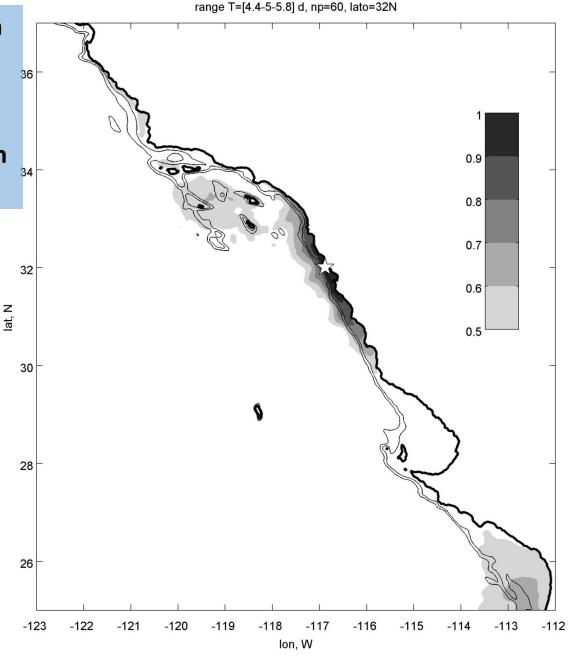


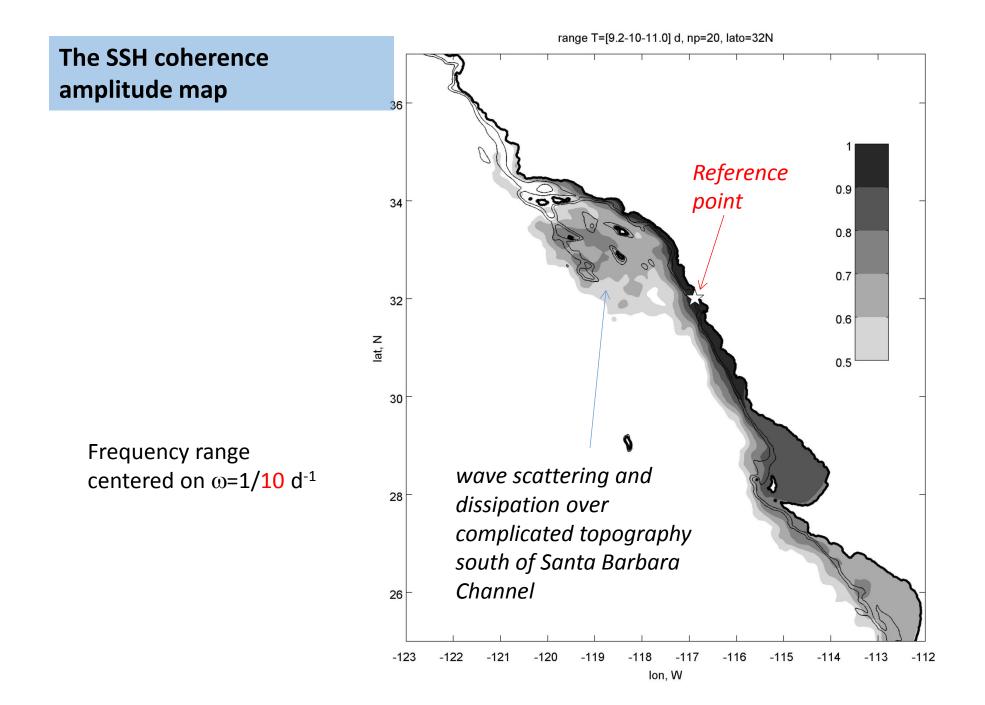
Neah Bay

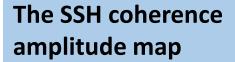
50

Using the 6-year model solution, we can compute 2-dimenstional SSH coherence amplitude maps and learn about spatial structure of the long coastal trapped waves (conduits of the signal from south to north)





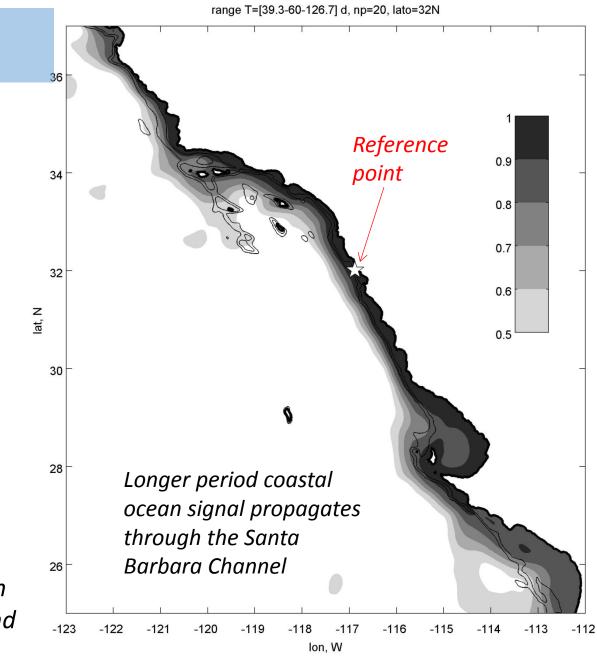




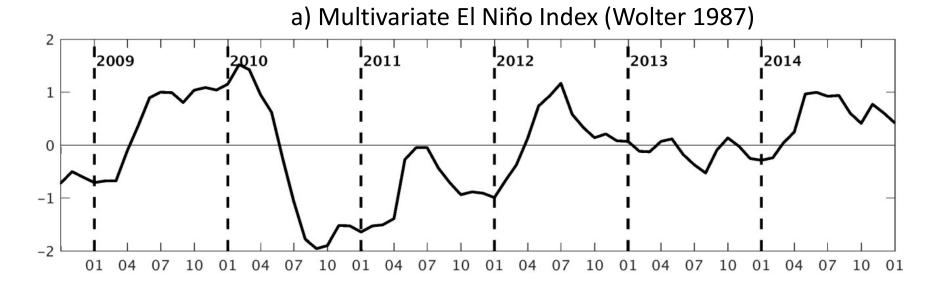
Frequency range centered on ω =1/60 d⁻¹

(40-126 days)

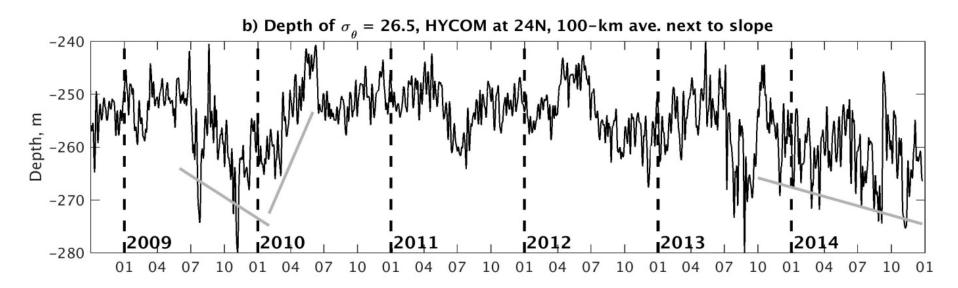
Low frequency signals from the southern boundary (24N) can influence coastal and slope variability in CA, OR, WA, BC



What kind of signal is found in the southern boundary conditions?



Depth of σ_{θ} =26.5 HYCOM at 24N (averaged between 0-100 km from cont. slope)



WCOFS: T and S are nudged to the global HYCOM solution in a 100-km wide band around the WCOFS open boundary

Properties on the isopycnal surface 26.5

In the entire domain, for each day, at each grid point we obtain

- depth of the isopycnal surface, z(x,y,t)
- T, S, u, v on this surface, PV

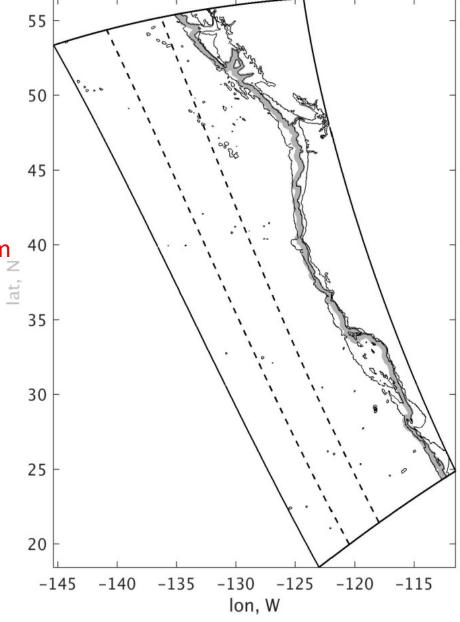
Average properties in cross-shore direction between 0-30-km offshore of the 200-m isobath (see gray shade in the figure)

 \Rightarrow e.g., z(s,t), T(s,t) where s = distance from southern boundary

T(s,t)= time_ave_T(s) + seasonal_T(s,t)+anomaly_T(s,t)

seasonal_T = harmonic fit (annual + semi-annual)

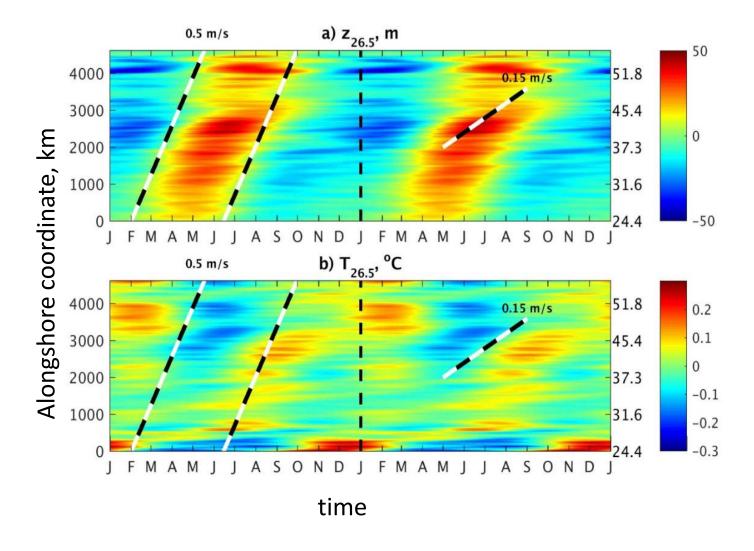
Effect of CTW? Effect of undercurrent? Climatology, anomalies?



Seasonal patterns (2 identical years shown):

Isopycnal depth: ±30 m, propagating pattern (0.5 m/s)

Temperature: subarctic waters are displaced by undercurrent waters (0.15 m/s), apparent in CA-OR-WA



Anomalies in the properties on the 26.5 isopycnal over the continental slope:

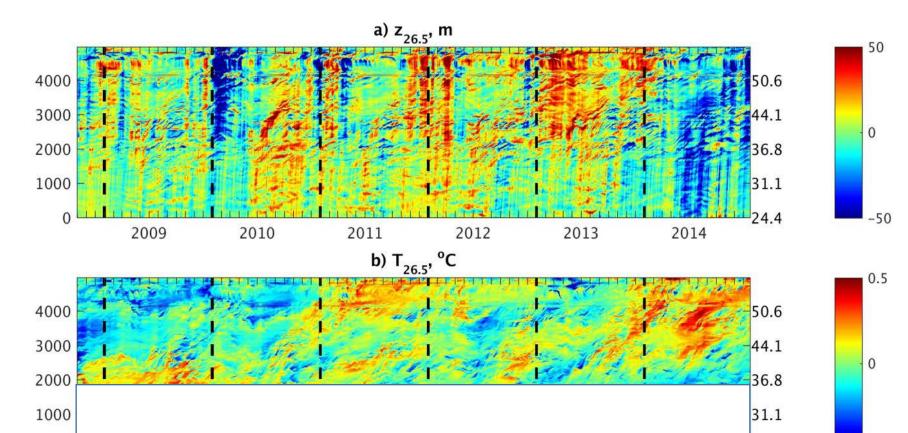
Z:

- fast propagating CTW
- El Niño effect in winter 2010 (atm. teleconnection)
- El Niño effect in summer 2014: propagation from the southern boundary

2009

2010

2011



2012

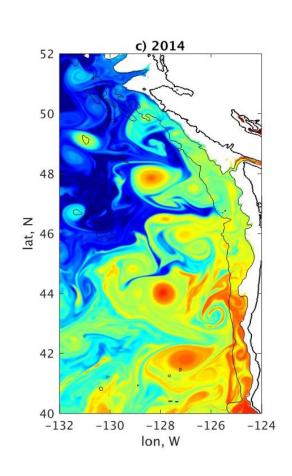
2013

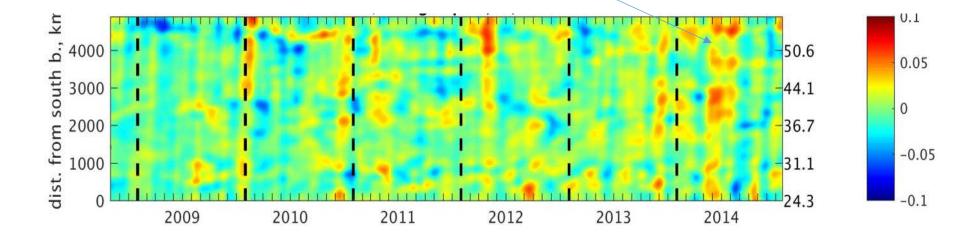
24.4

2014

T: warmer in OR-WA-BC

Anomaly from the seasonal cycle in alongslope current on 26.5: average across 0-100 km from 200 m isobath (i.e. average across the width of eddies separated from the undercurrent): the 2014 anomaly in v on the isopycnal is seen better





(A snapshot of T on the isopycnal surface, 1 Jun 2014: Coastal undercurrent eddies "cuddies" are apparent over the slope and in the adjacent ocean interior on the 26.5 isopycnal surface)

SUMMARY:

- A 2-km resolution regional ocean model of the West Coast from Mexico to BC has been developed as the base component of the NOAA West Coast Ocean Forecast System (WCOFS)
- Warm blob anomaly in NEP in Jan 2016 is not associated with anomaly in the atmospheric heat flux
- The wide-spread warming along the US West Coast in spring-summer 2014 is associated with the anomalous SWRAD
- Warm anomaly over the OR slope in summer-fall 2014 can be associated with the deepening of subsurface isopycnal surfaces over the continental slope in Mexico (24N), the CTW propagation, isopycnal surface depression in the north and the resulting stronger undercurrent
- The seasonal cycle in the depth of the 26.5 isopycnal shows a south-to-north propagating pattern, with the speed of 0.5 m/s. Temperature and PV show patterns characteristic of advection by the undercurrent (0.1-0.15 m/s)