

Introduction

The North Pacific thermocline was one of the first regions where large variations in oceanic O₂ concentrations were observed and linked to ventilation changes (Andreev and Kusakabe, 2001; Ono et al., 2001; Watanabe et al., 2001; Emerson et al., 2004). Several mechanisms responsible for the observed variability in thermocline O₂ content have been suggested (Whitney et al., 2007; Sasano et al., 2018). Here, we investigate the hypothesis that changes in outcrop location/area of the densest outcropping isopycnals (including cessation of outcropping) may be causing O₂ variations in the thermocline downstream (Emerson et al., 2004; Mecking et al., 2008; Kwon et al., 2016).

Figure 4. Time series of O₂ on σ_{θ} = 26.6 kg m⁻³ at OSP (solid blue line) and of annual maximum outcrop area of σ_{θ} = 26.4–26.7 kg m⁻³, used as a proxy for ventilation, based on the EN4-OISST dataset (solid red lines). Also shown are the data filtered with a 5-year running mean (solid cyan and magenta lines, respectively) and the linear trends in each time series (dashed blue and red lines, respectively). Noticeably, a strong minimum in O_2 in 2005–2007 lags the minimum in annual maximum outcrop area in 1995–1997 by about a decade. Such a lag is expected, assuming it takes about 10 years for surface waters from the outcrop area in the northwestern North Pacific to reach OSP on $\sigma_{\rm A}$ = 26.6 kg m⁻³.





Figure 1. Examples of maxima (a, c) and minima (b, d) of March outcrop areas of the σ_{θ} = 26.4–26.7 kg m⁻³ isopycnal range: (a) 1983, (b) 1996, (c) 2001, and (d) 2015. Dataset used is EN4-OISST. Color shading and color bar show surface potential density anomalies (σ_{θ}), with anomalies in the σ_{θ} = 26.4–26.7 kg m⁻³ range indicated by red markers. Contour lines also show σ_{θ} outcrops as labeled in (c). Location of OSP (50°N, 145°W) is marked by an open black circle. Area used for northwestern North Pacific SST, SSS, and surface density averaging (Figs. 6&7) is outlined by a black box in (a).

Data



← Figure 5. Lagged correlations between detrended time series of O_2 on σ_{θ} = 26.6 kg m⁻³ at OSP and of annual maximum outcrop area of σ_{θ} = 26.4–26.7 kg m⁻³ from EN4-OISST: using unfiltered time series data (blue line) and time series data filtered with a 5-year running mean (red line). Maximum correlation values (r) for filtered data are 0.50 at a +10-year lag and 0.59 at a -10-year lag in agreement with ~10 year travel time (assuming bi-decadal cycles).



Figure 6. Annual maximum surface σ_{θ} averaged over 40–50°N, 150– 170°E (see box in Fig. 1a) and annually averaged PDO index. Time series track each other e.g. 1999–2005, but show opposite variations e.g. 1995–1999 & 2013–2018, resulting in a correlation close to 0.



Sea surface temperature (SST), salinity (SSS), and density

- EN4: 1°x1°, similar to WOD, use T & S at 5m, 1900–2020 (Good et al., 2013) - OISST: 1/4°x1/4°, SST from satellite & in situ, 1982–2020 (Reynolds et al., 2007) – EN4-OISST: created from en4 SSS (interpolated to1/4°x1/4°) & OISST, 1982–2020

Oxygen (O₂) measurements in ocean interior

- Ocean Station P (OSP) timeseries: 50°N, 145°W, 1956–2020 (Whitney et al., 2007) – GO-SHIP repeat hydrography at 152°W (P16), 47°N (P1), for comparison (not shown here)

Pacific Decadal Oscillation (PDO) Index (Mantua et al., 1997)

Figure 2. Time series of the surface outcrop area of the σ_{θ} = 26.4–26.7 kg m⁻³ isopycnal range in the North Pacific from 1982–2020. The blue line is based on surface density from the 1/4° EN4-OISST dataset, and the red line is based on surface density from the 1° EN4 dataset. The peaks in the outcrop area occur just after January (shown on the x axis) of each year, usually in March. There are distinct March minima in 1995–1997, 2011, and 2014–2015 compared to the March outcrop area in the years before and after with a distinct March maximum in 2001 in particular. Despite the 2001 maximum, the EN4 and EN4-OISST datasets both show a declining trend in the σ_{θ} = 26.4–26.7 kg m⁻³ outcrop area from 1982–2020.



Figure 7. (a) SST (bold blue line) and SSS (bold red line) at annual maximum surface σ_{θ} and (b) annual maximum surface σ_{θ} (bold black line) and contributions from SST (dashed blue line; with SSS fixed at mean annual cycle during density calculation) and SSS (dashed red line; with SST fixed) in the northwestern North Pacific (averaged over 40–50°N, 150–170°E; see Fig. 1a for area). The dashed red line (SST fixed) and dashed blue line (SSS fixed) in (b) explain 85% and 21% of variance of the full record, respectively, indicating that SSS has a much larger contribution to surface density variability. In terms of the long-term linear density decrease (solid black line), the SSS-based trend (red solid line) and the SST-based trend (blue solid line) contribute about equally. Dataset used is the EN4-OISST dataset.

Summary & Conclusions

- The size of the annual maximum $\sigma_{\theta} = 26.4 - 26.7$ kg m⁻³ outcrop area is used as an indicator of how much ventilation is occurring at the bottom of the ventilated thermocline. This is compared to O₂ on σ_{θ} = 26.6 kg m⁻³ where large O₂ variations occur at OSP. – A correlation is found with a lag that approximately matches the travel time of water from the northwestern North Pacific to the northeast (10 years). A correlation between surface density variations and the PDO (defined by SST variability) is not as clear. -SSS appears to dominate over SST in causing decadal changes in surface density whereas SSS and SST contribute about equally to long-term declining density trends.



Figure 3. (a) Time series of O_2 at OSP (see location in Fig. 1) on σ_{θ} = 26.2–27.0 kg m⁻³. Linear trends of O_2 on each isopycnal are shown as dashed lines. (b) Profiles of linear O₂ trends (blue) and of standard deviation of the O_2 variations (red) on isopycnals at OSP, with maxima near σ_{θ} = 26.6 kg m⁻³ as also found using repeat hydrography (Mecking et al. 2008). Data used are objectively mapped O_2 data at OSP. Maxima in declining O_2 trends at OSP have also been reported near this density by Whitney et al. (2007; σ_{θ} = 26.5 kg m⁻³) and Crawford and Peña (2016; σ_{θ} = 26.7 kg m⁻³), each using slightly different methodologies. This confirms that large decadal variations and declining trends in O₂ occur on isopycnals near the bottom of the ventilated thermocline.

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