

2nd ESSAS 23 May 2011, Seattle, WA, USA

Toward a simulation of iron circulation from the Okhotsk Sea to the Pacific

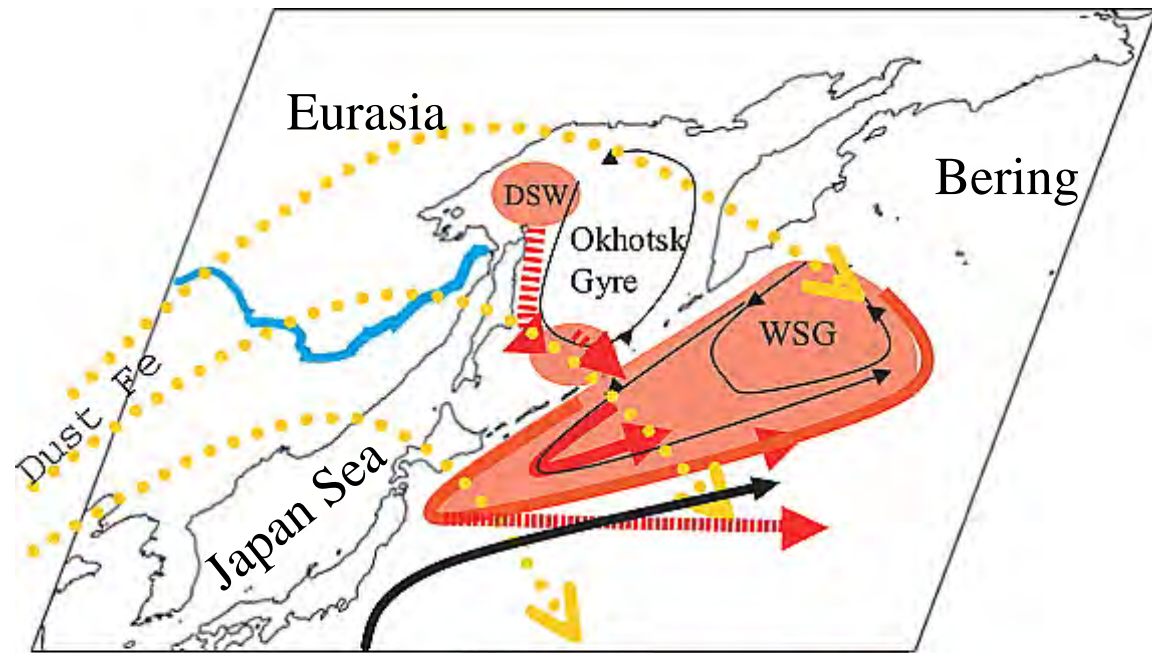
Keisuke Uchimoto¹

Tomohiro Nakamura¹, Jun Nishioka¹, Humio Mitsudera¹,
Kazuhiro Misumi² and Daisuke Tsumune²

1: Institute of Low Temperature Science, Hokkaido University, JAPAN

2: Central Research Institute of Electric Power Industry, JAPAN

Toward a simulation of iron circulation from the Okhotsk Sea to the Pacific



Fe transport →
Fe rich intermediate water ○

Nishioka et al. (2007, JGR)

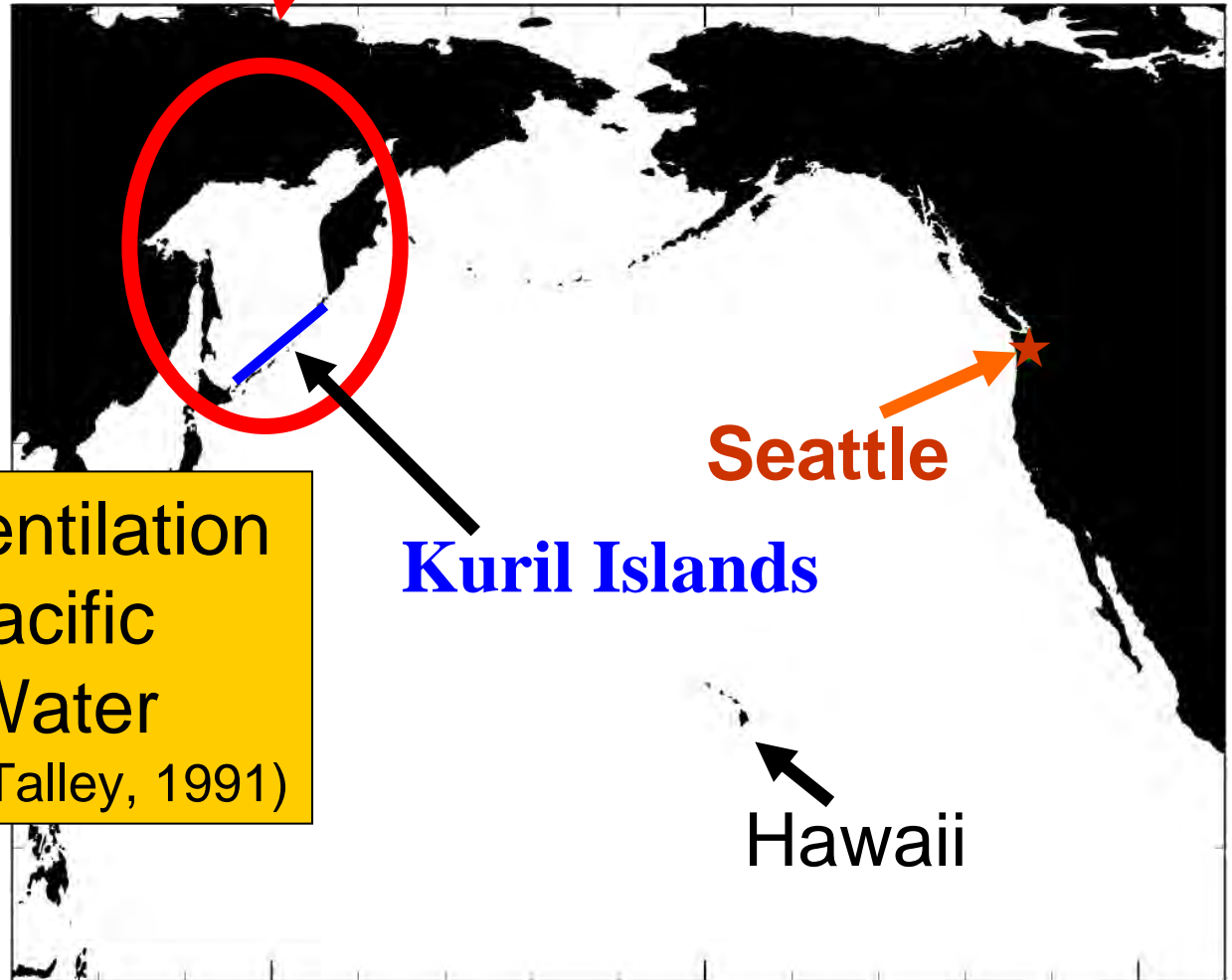
Presentation Outline

- **Introduction:** Importance of intermediate layer in Okhotsk Sea
- **Ocean model description**
- **Tracer experiment:** DSW on the shelf
- **Simulation of chlorofluorocarbons (CFCs):** intermediate layer ventilation in Okhotsk Sea
- **Iron model description:** Parekh's model
- **Simulations of iron distribution:** results of preliminary runs
- **Summary**

Introduction

Ventilation of the intermediate layer
in the Sea of Okhotsk

Sea of Okhotsk



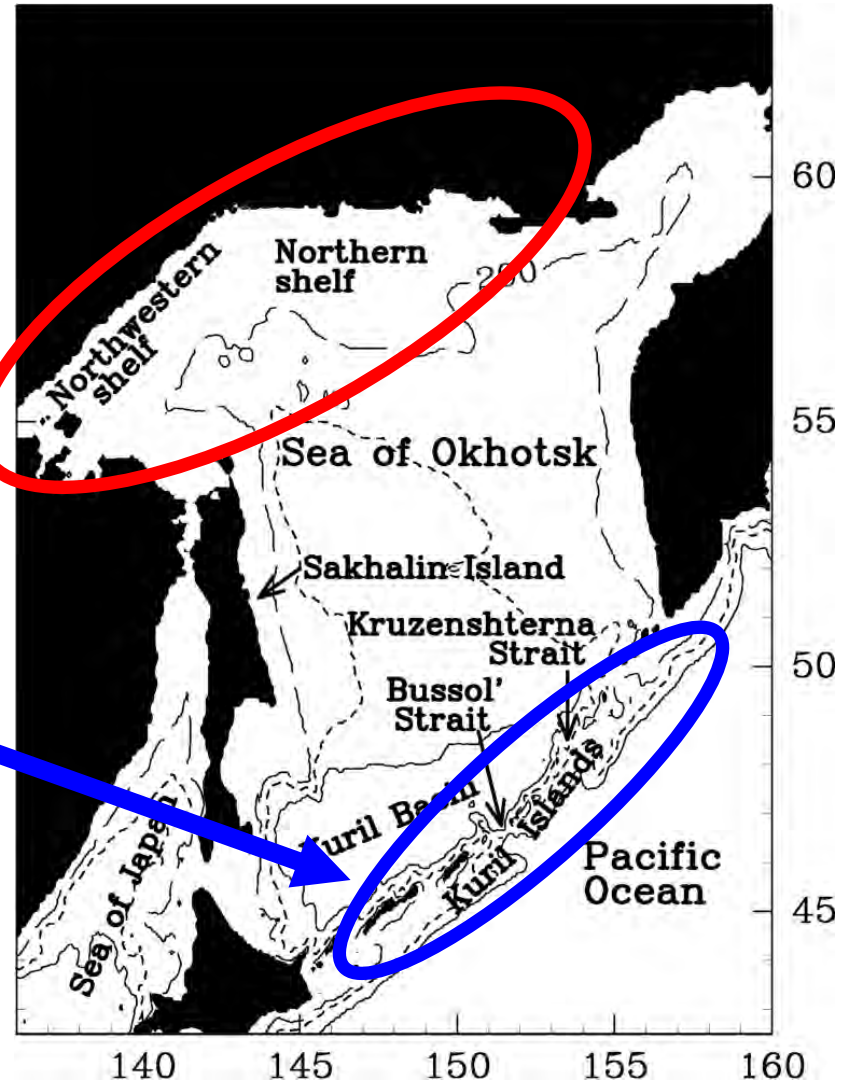
Location for ventilation
of the North Pacific
Intermediate Water
(e.g. Talley, 1991)

Sea of Okhotsk

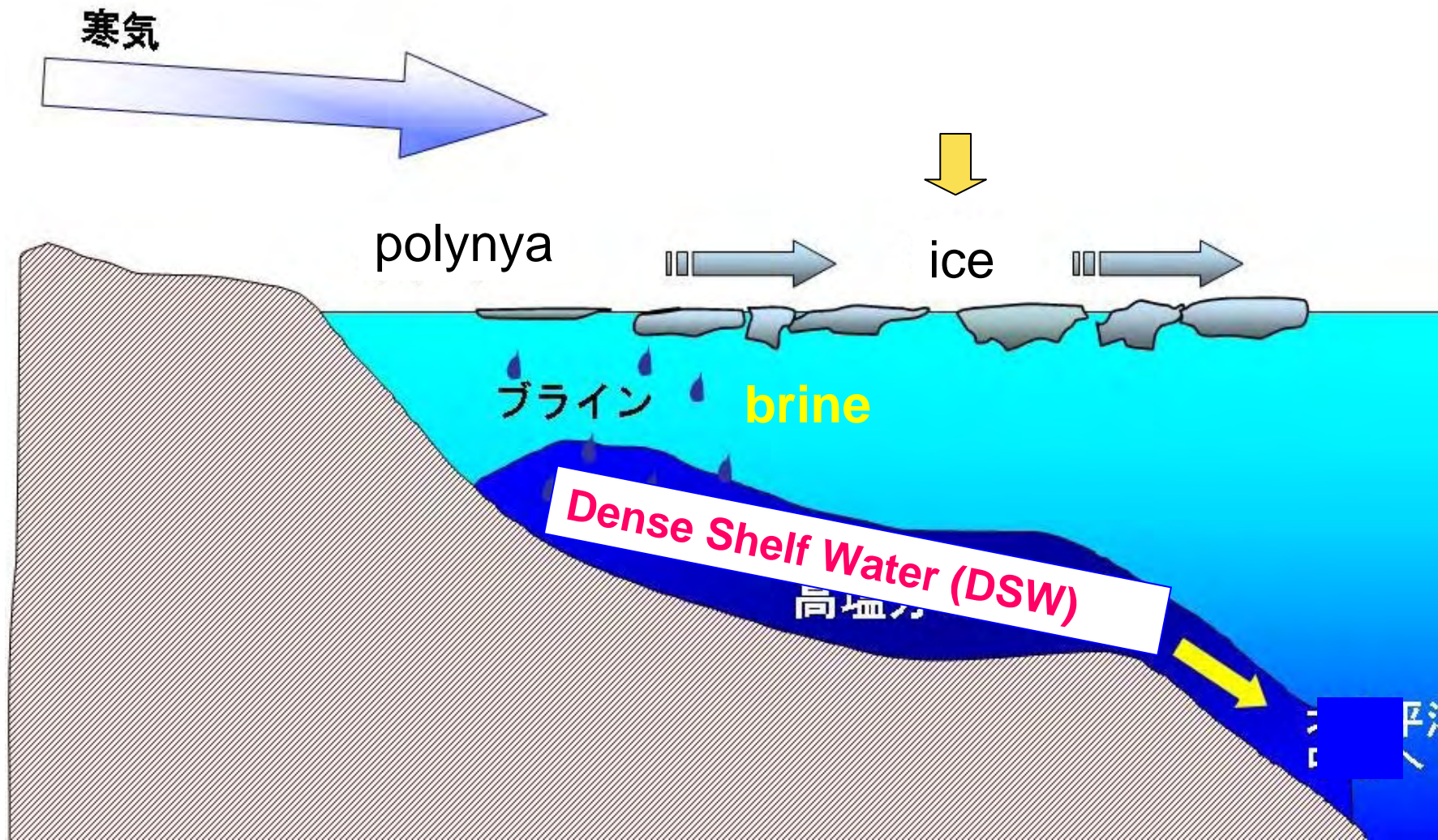
- Two ventilation processes for the intermediate layer
 - brine rejection during sea ice formation
 - tidal mixing along the Kuril Islands

Ventilation Processes Control Fe Transport

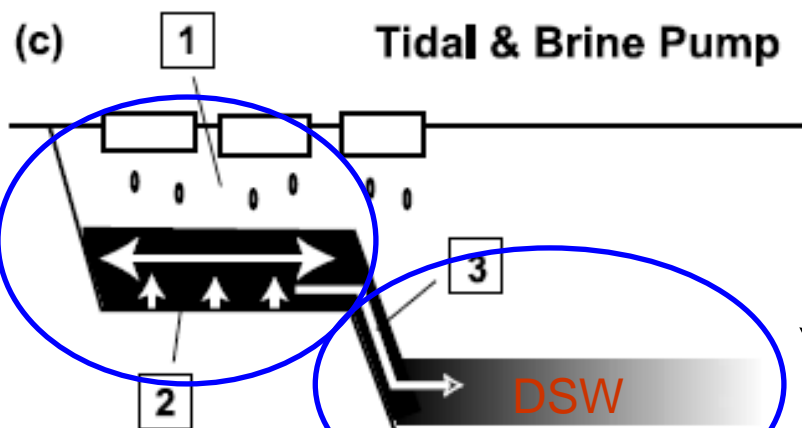
suggested by Nishioka et al. (2007, JGR)



brine: high salinity water



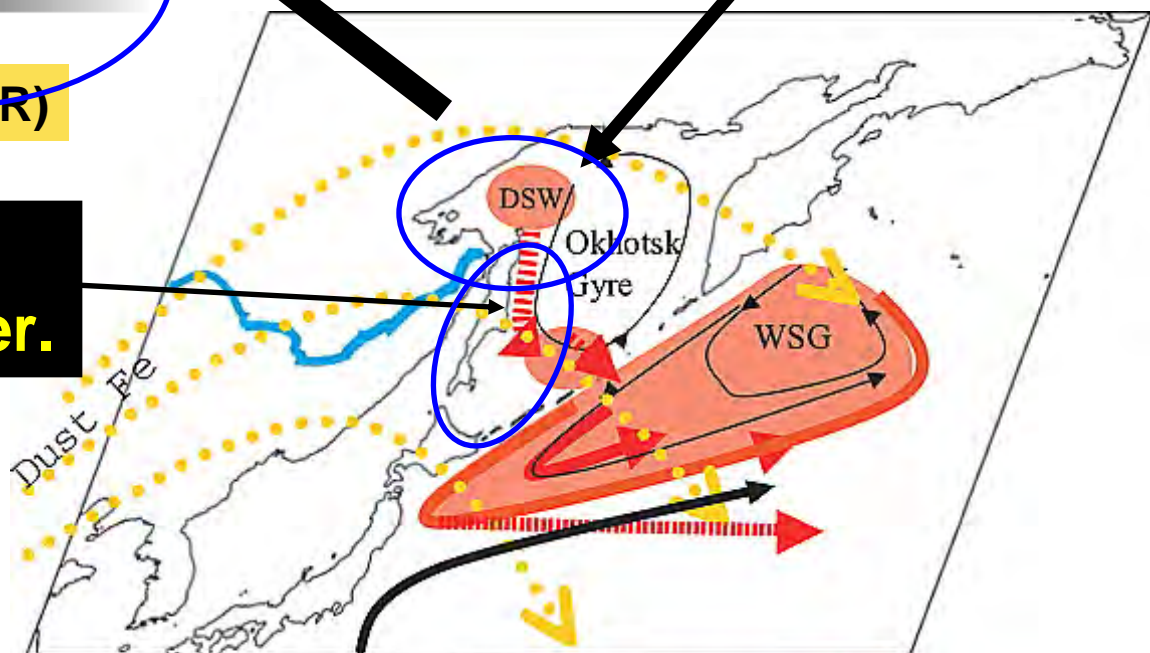
How brine rejection controls Fe transport?



Dense shelf water (DSW) is formed in the northern part of the Sea of Okhotsk

Nakatsuka et al. (2004, JGR)

Fe is transported in the intermediate layer.



Fe transport

→ Nishioka et al. (2007, JGR)

Fe rich intermediate water

○ (2007, JGR)

Misumi et al. (accepted, JGR) well
simulate the Fe distribution in the North Pacific.

But their model *poorly represents ventilation*
processes in the Sea of Okhotsk.

An ocean model

that can represent the ventilation of the intermediate layer in the Sea of Okhotsk

is essential for simulating iron distribution.



- **Tracer experiment:** behavior of DSW
- **CFC simulation:** ventilation of intermediate layer

MODEL

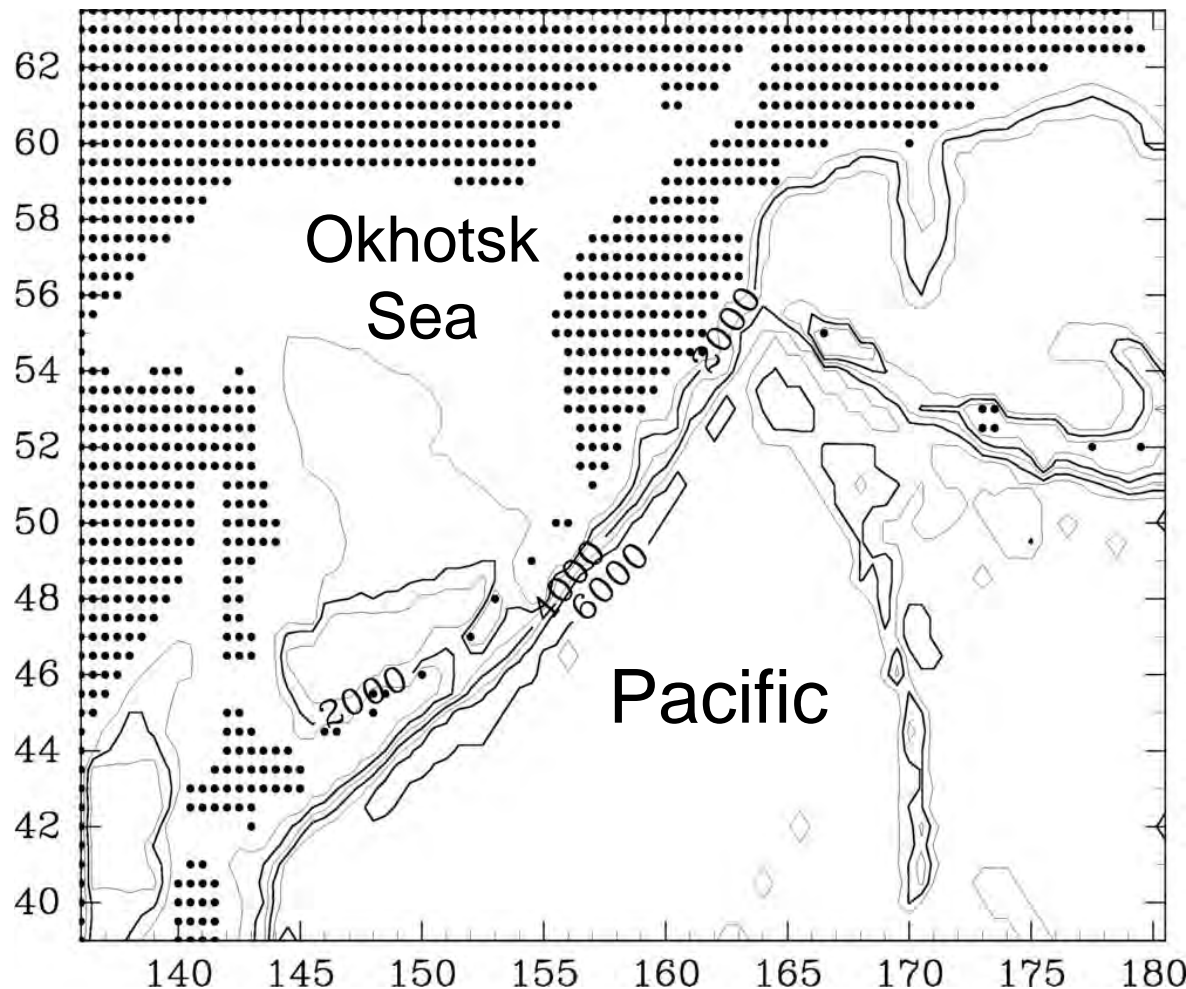
OGCM: Sea Ice-Ocean Model (COCO);
primitive equation model

Resolution:

horizontally 0.5°

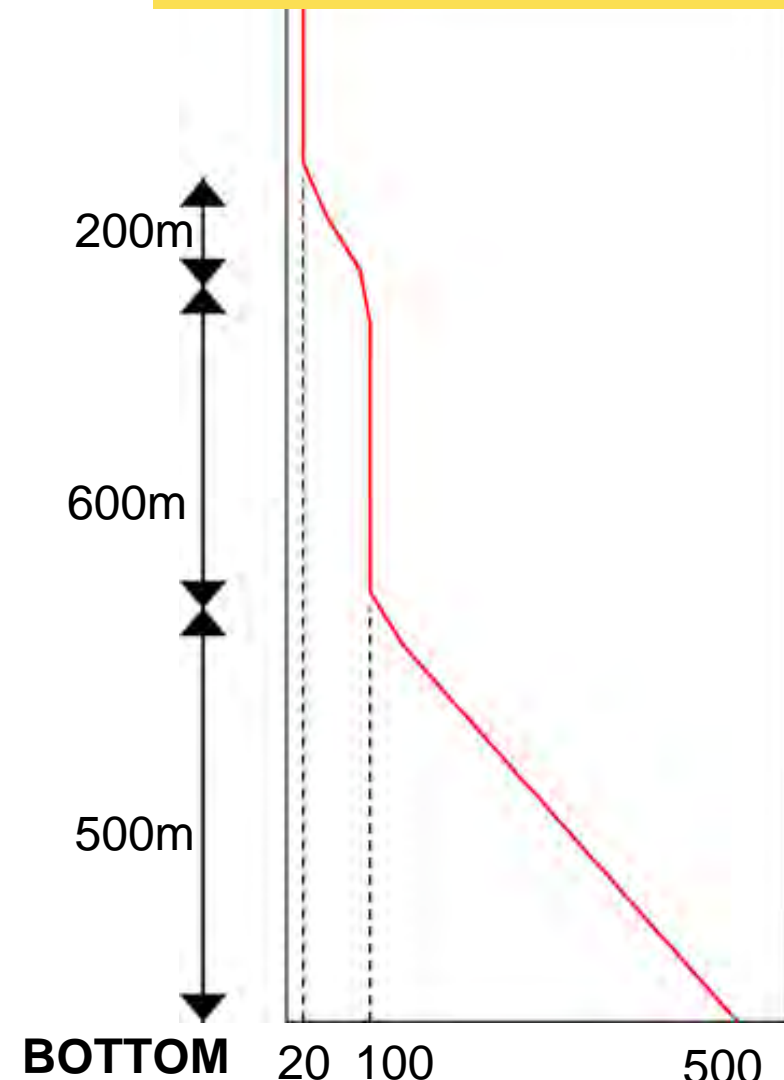
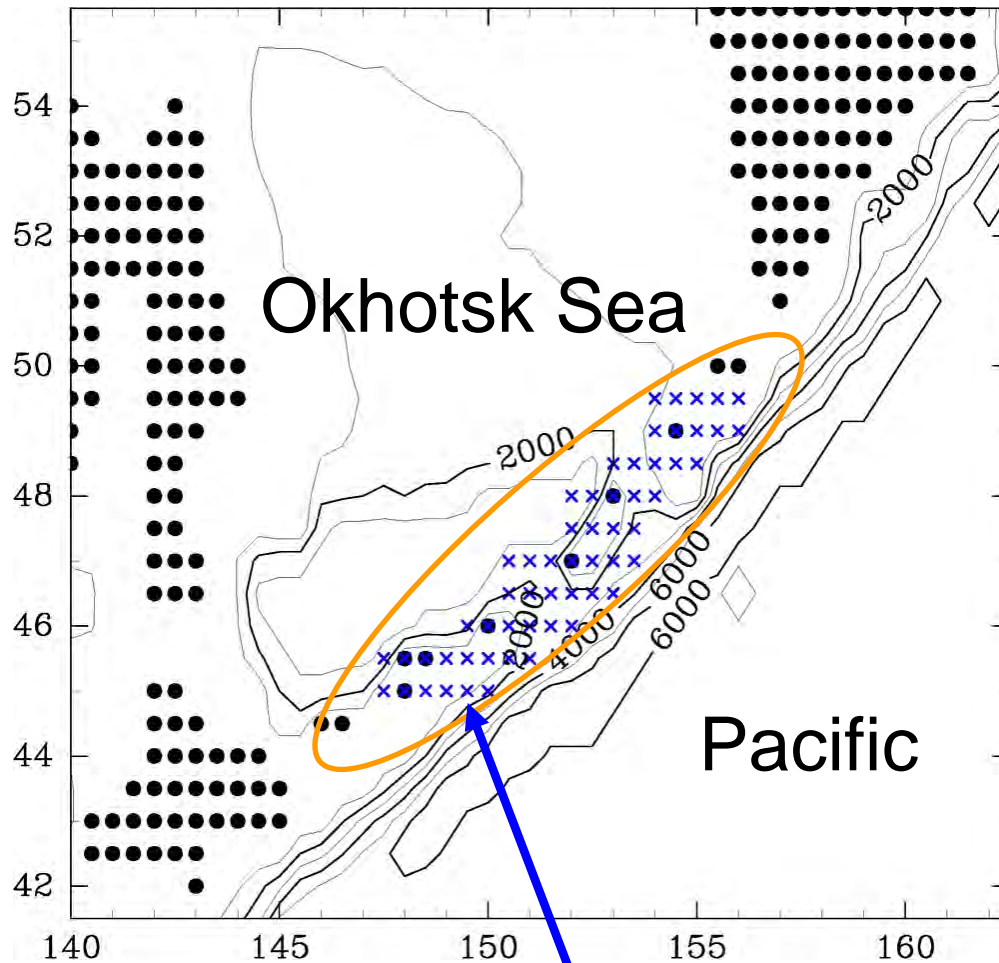
vertically 51 levels

Forcing: daily climatological atmospheric data (OMIP data)



Tidal mixing parameterization along the Kuril Islands

Vertical diffusion coefficients as tidal mixing effects



CROSSES

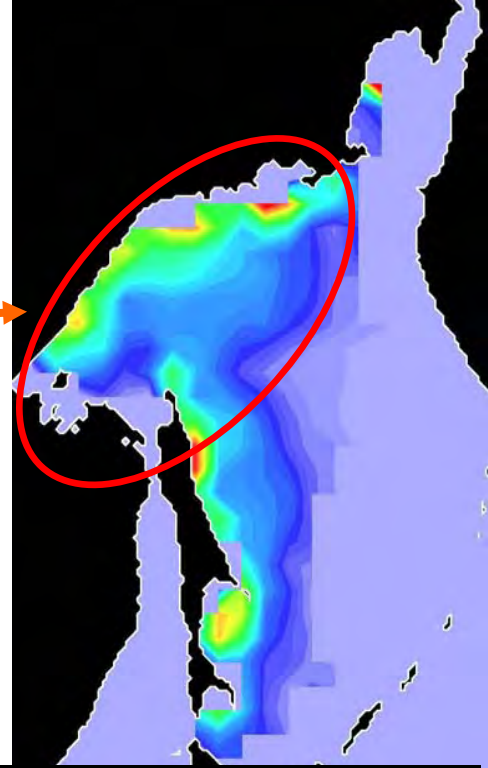
grids with the increased vertical diffusion as tidal mixing effects

cm^2 / s

Tracer experiment

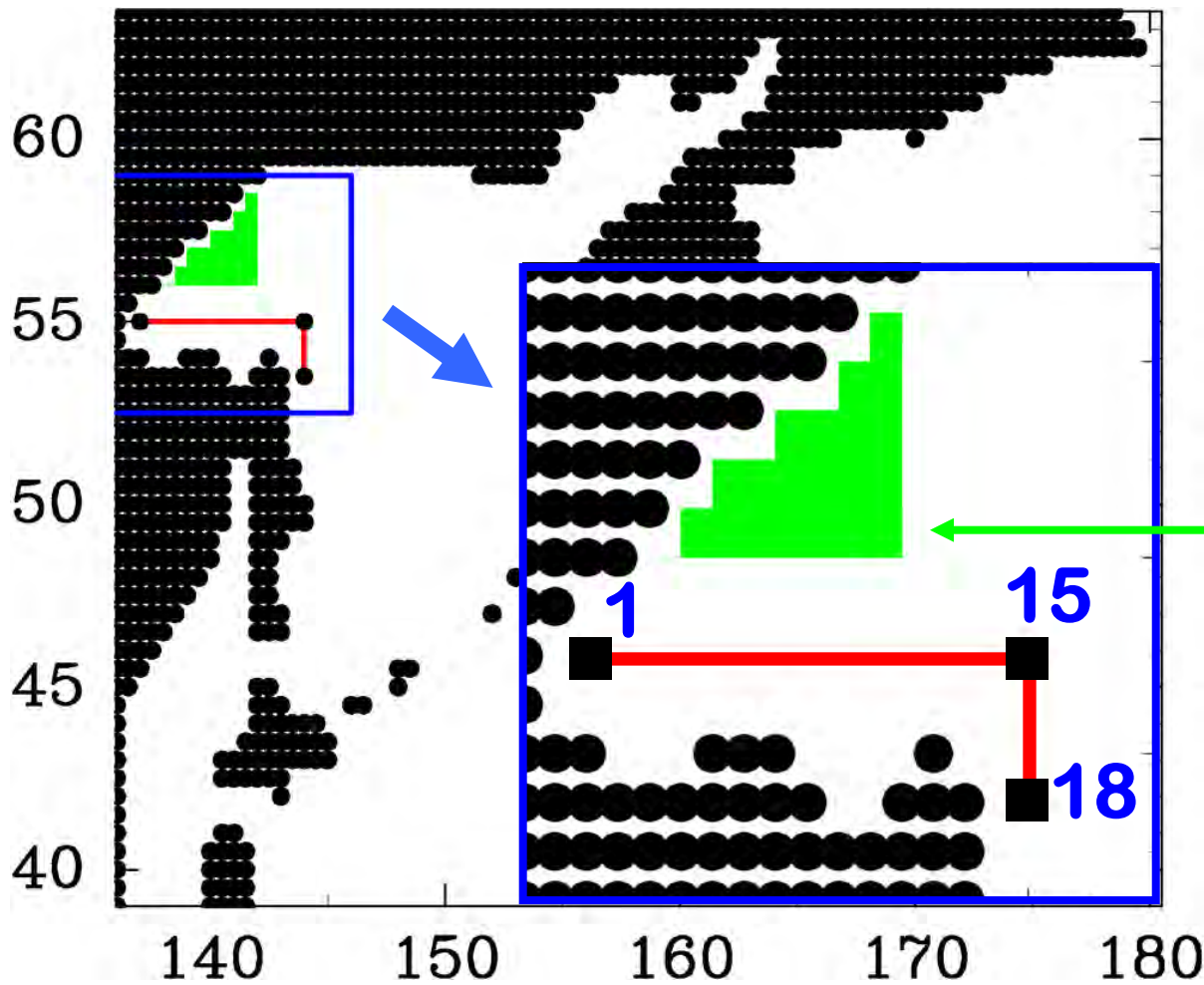
Uchimoto et al. (2011, Hydrological Research Letters)

Highest Sea Ice production area

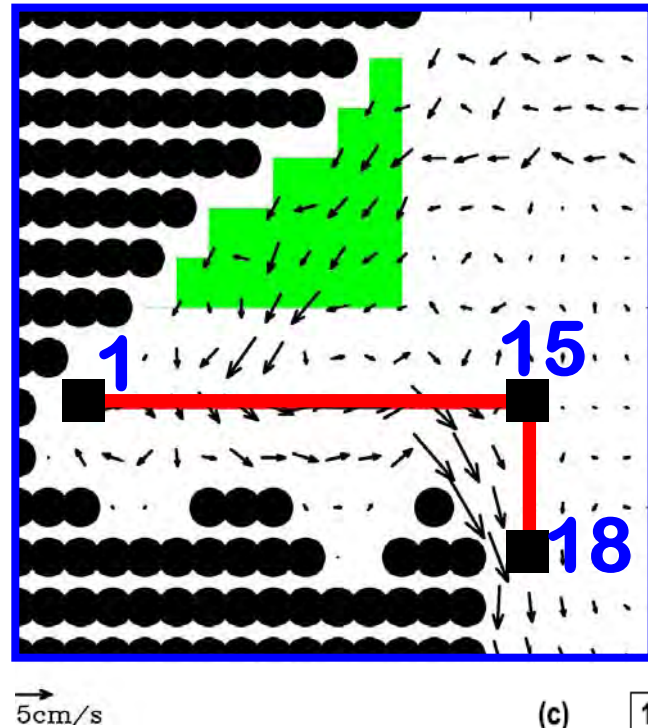
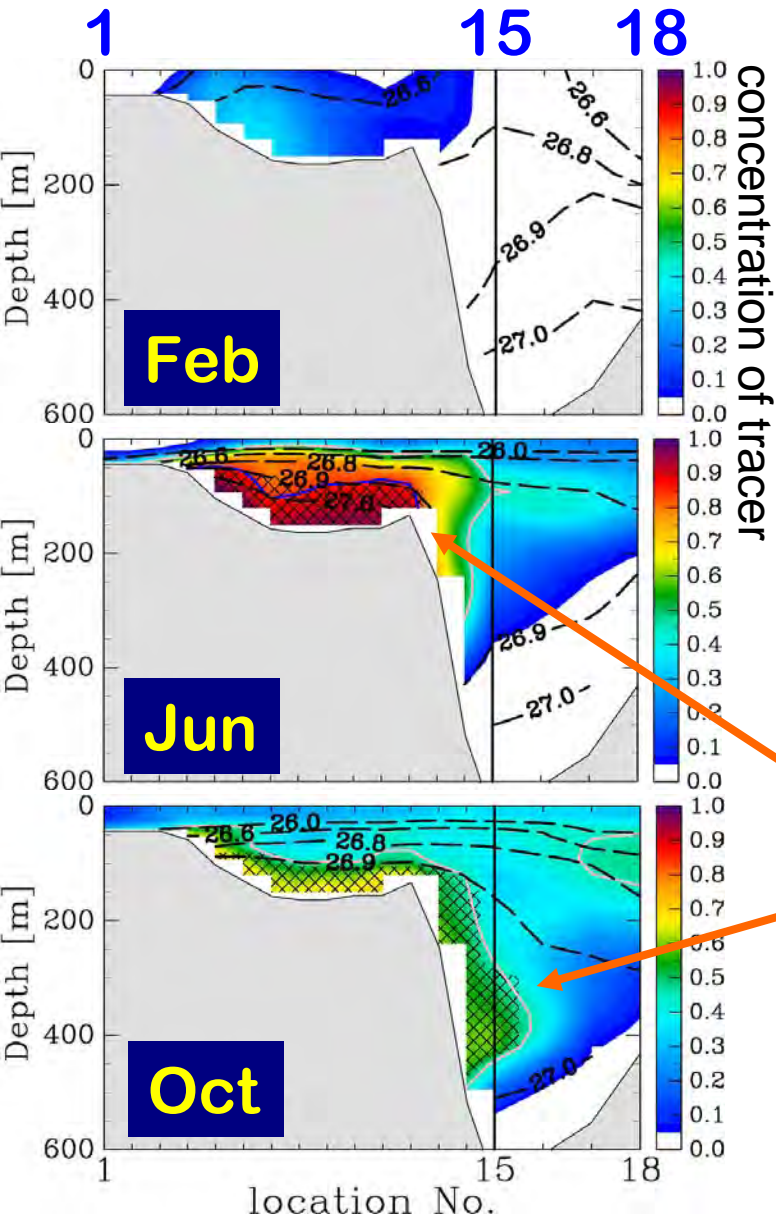


Ice production
Kimura and Wakatsuchi (1999)

Tracer is injected
from Jan. to Apr.
at the sea surface
(restored to 1.0).

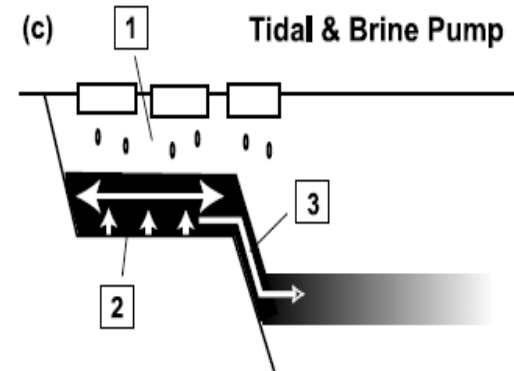


Tracer experiments



Vectors:
annual mean
velocity

DSW ($\sigma_{\theta} > 26.9$, $T < -1^{\circ}\text{C}$)
cross-hatching



Nakatsuka et al. (2004, JGR)

CFC simulation

Uchimoto et al. (2011, JGR)

performed according to OCMIP-2 protocols

Ventilation of intermediate layer in the model

- CFC on $26.8 \sigma_{\theta}$ and $27.4 \sigma_{\theta}$ surfaces
- Δp CFC along path of DSW

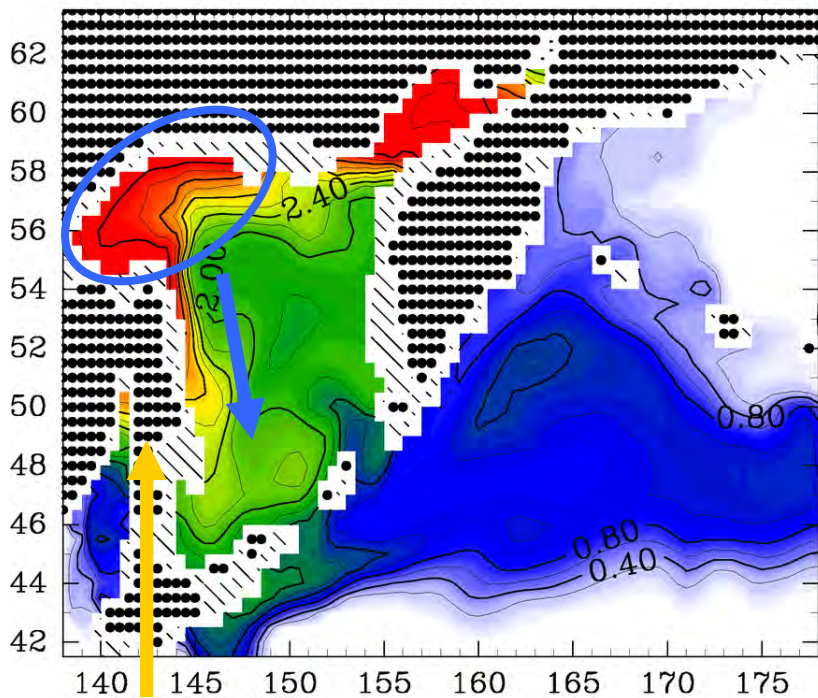
CFC12 on $26.8 \sigma_{\theta}$, $27.4 \sigma_{\theta}$

Ventilation owing to dense shelf water (DSW)

Ventilation owing to tidal mixing along Kuril

CFC12 on 26.80

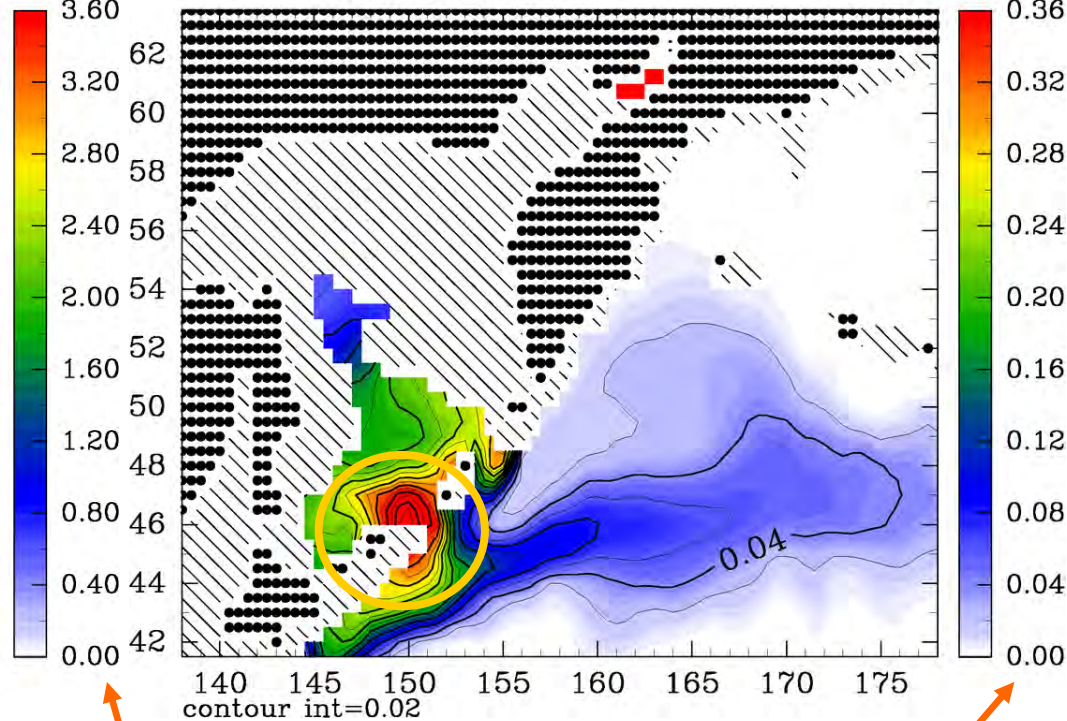
SEP



Sakhalin Island

CFC12 on 27.40

SEP



Color scales are different by a factor of 10.

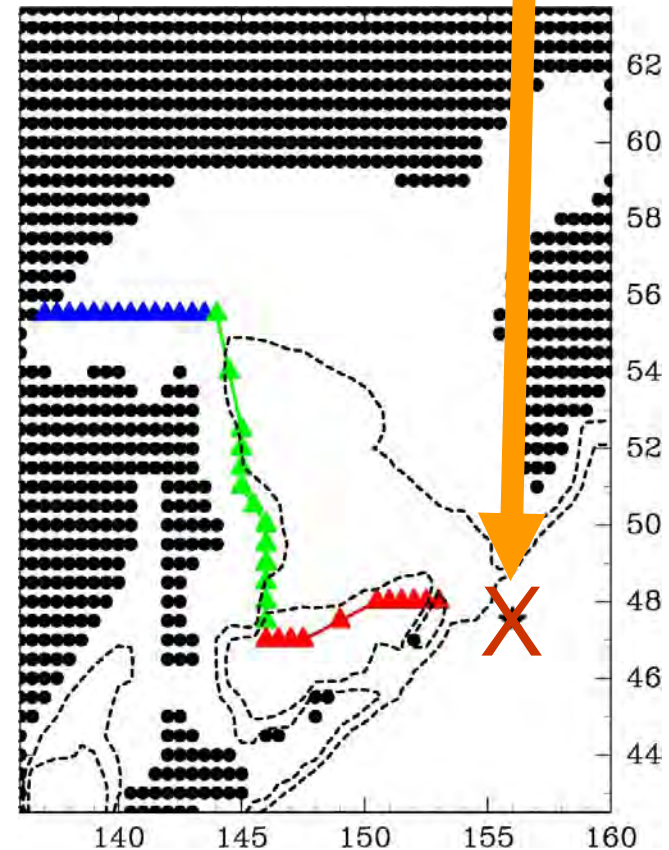
$\Delta p\text{CFC}$ (Δ partial pressure of CFC)

$$\Delta p\text{CFC}(\vec{x}, \rho) = p\text{CFC}(\vec{x}, \rho) - p\text{CFC}(\rho)$$

at the ref. point



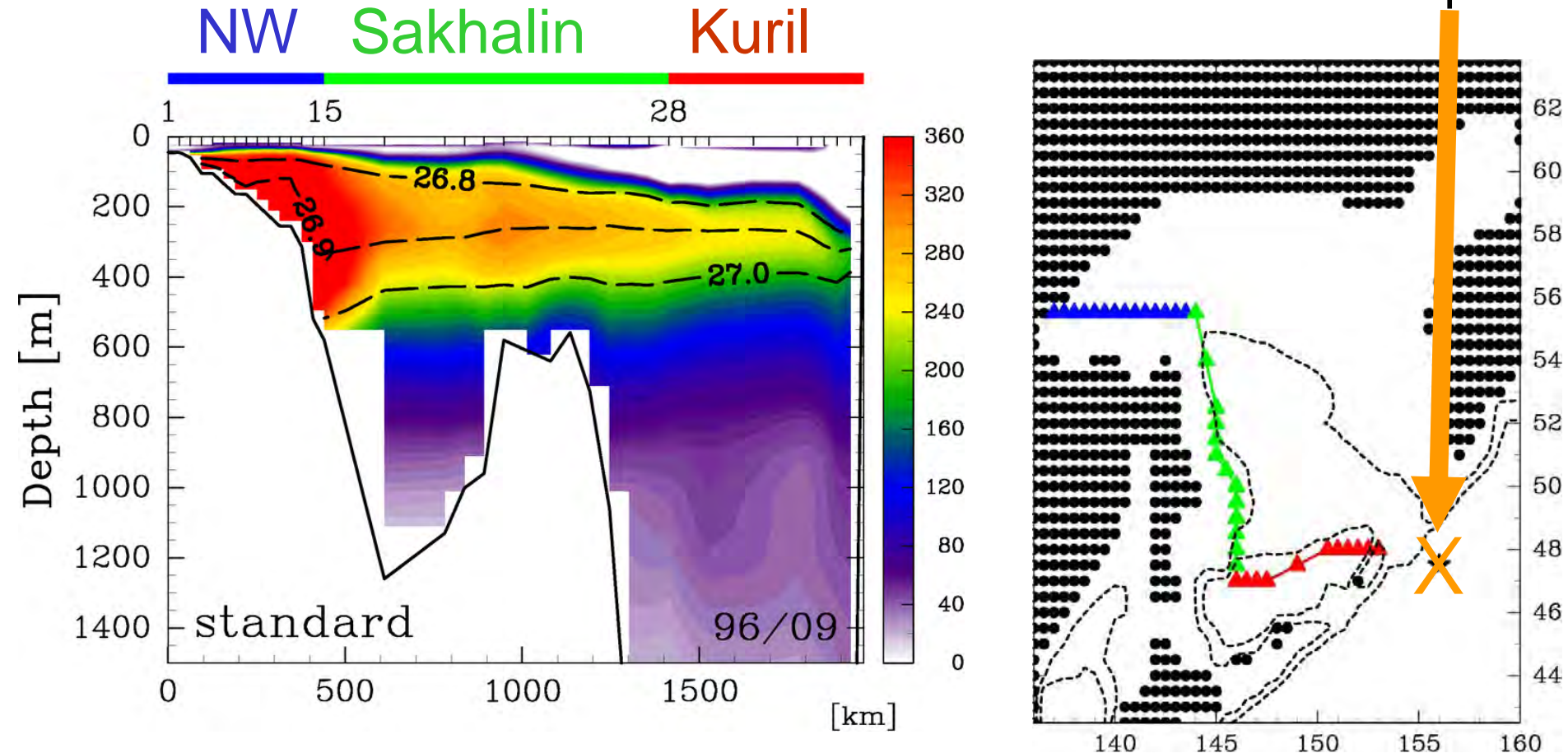
Index of the water ventilated
within the Sea of Okhotsk



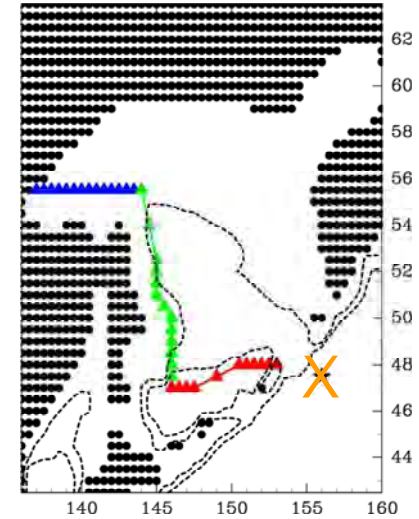
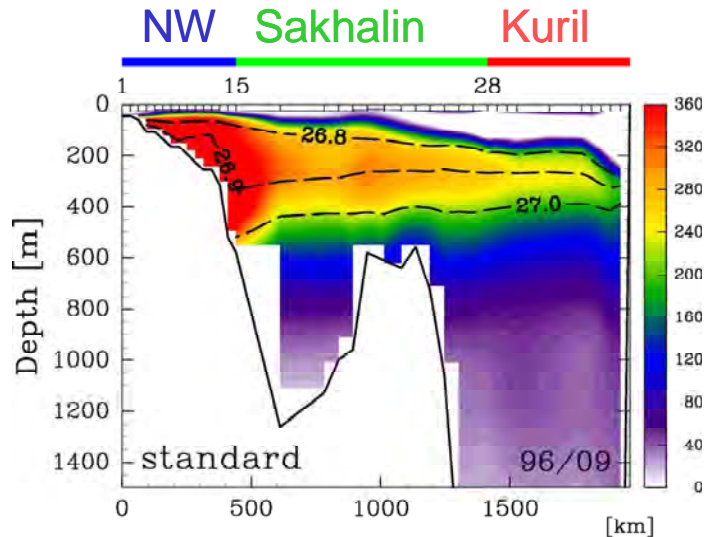
$\Delta p\text{CFC}$ (Δ partial pressure of CFC)

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at the ref. point



$\Delta pCFC$ (Δ partial pressure of CFC)

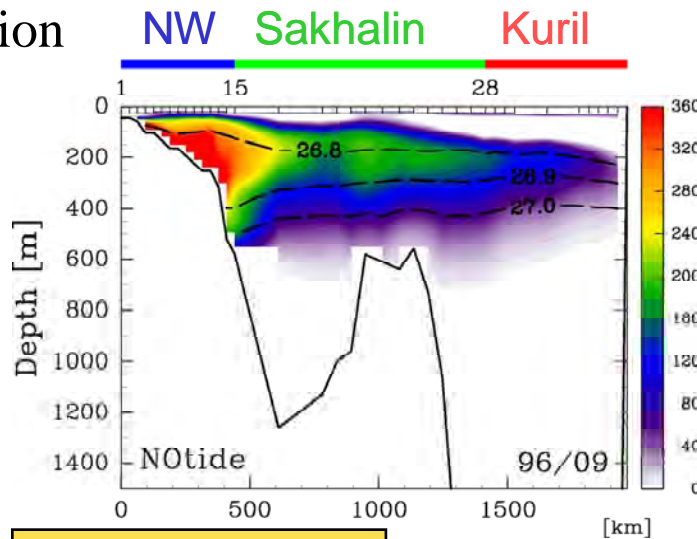


Two experiments

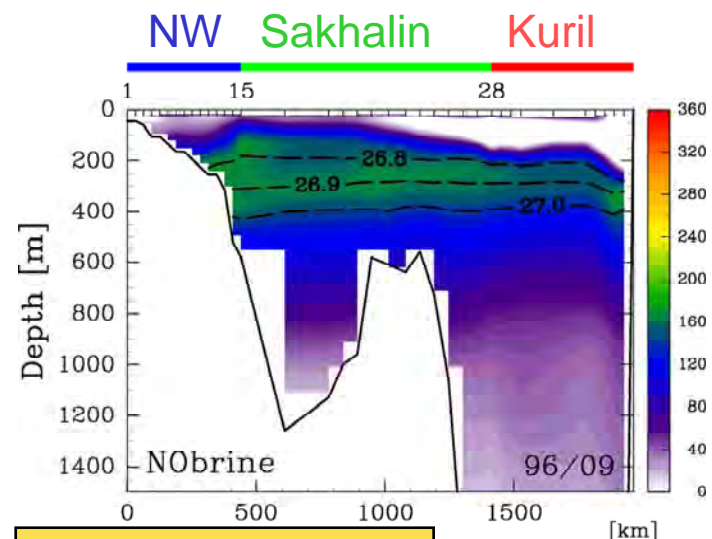
- without tidal mixing along Kuril Islands
- without brine rejection

Ventilation through brine rejection

Ventilation through tidal mixing



without tidal mixing



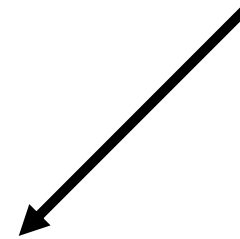
without brine rejection

simulation of iron circulation

Iron (biogeochemical) Model description

Model developed by Parekh et al. (2005, GBC)

PO₄, DOP (dissolved organic P), and Fe (total dissolved iron)



$$\text{Fe} = \text{Fe}' + \text{FeL}$$

Fe': free

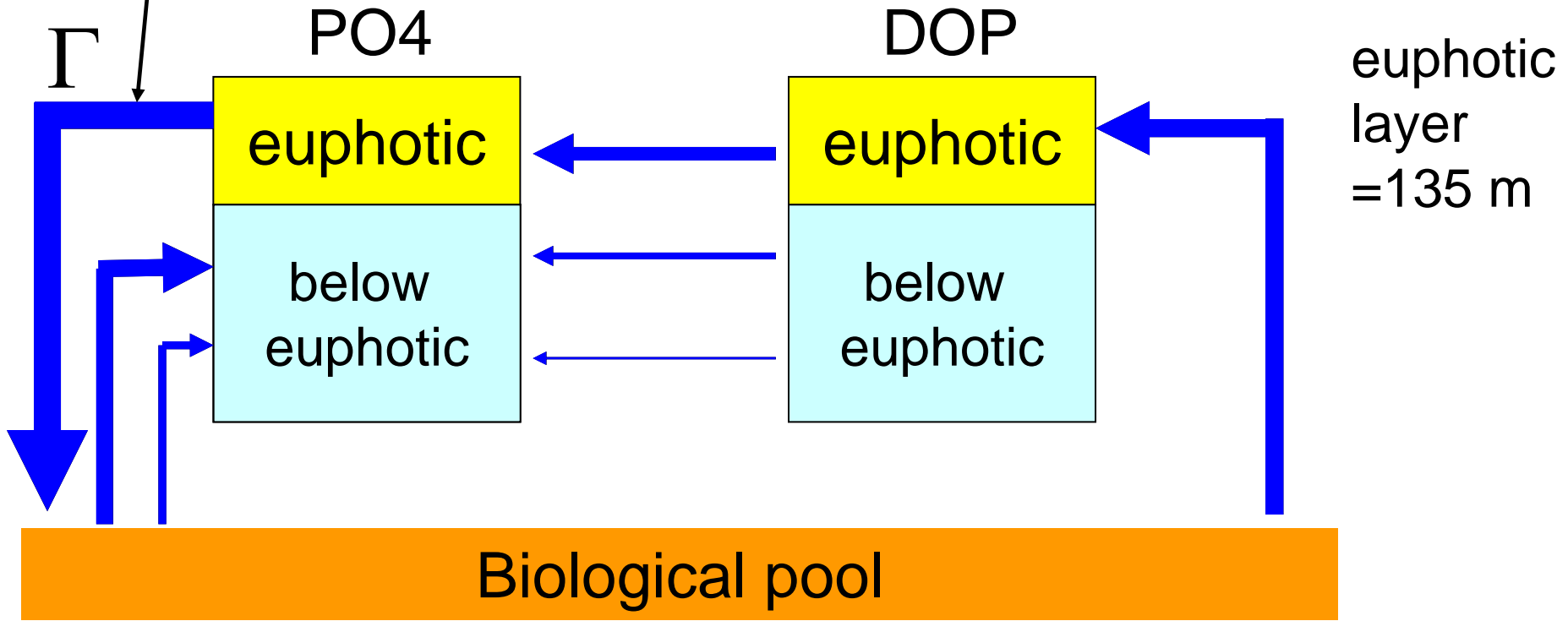
FeL : complexed (L : ligand)

Phosphorus cycle

$$\Gamma = \alpha \frac{PO_4}{PO_4 + K_{PO_4}} \frac{Fe}{Fe + K_{Fe}} \frac{I}{I + K_I} \quad (I : \text{light})$$

Biological uptake

remineralization DOP→PO4



Iron Cycle

Fe (dissolved)

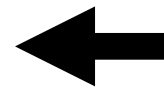
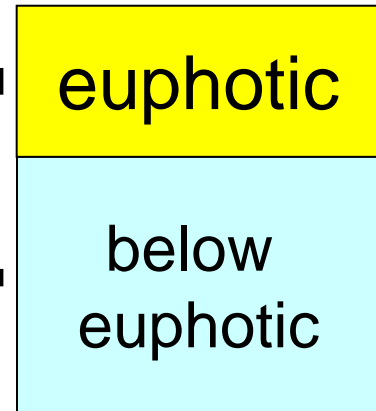
•Fe is 3.5% of dust
•solubility 1%

•Biological uptake

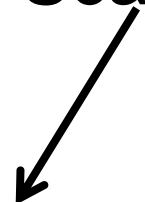
•Remineralization

→proportional to those of PO₄.

aeolian dust (Mahowald et al. 2005, GBC)



Scavenging

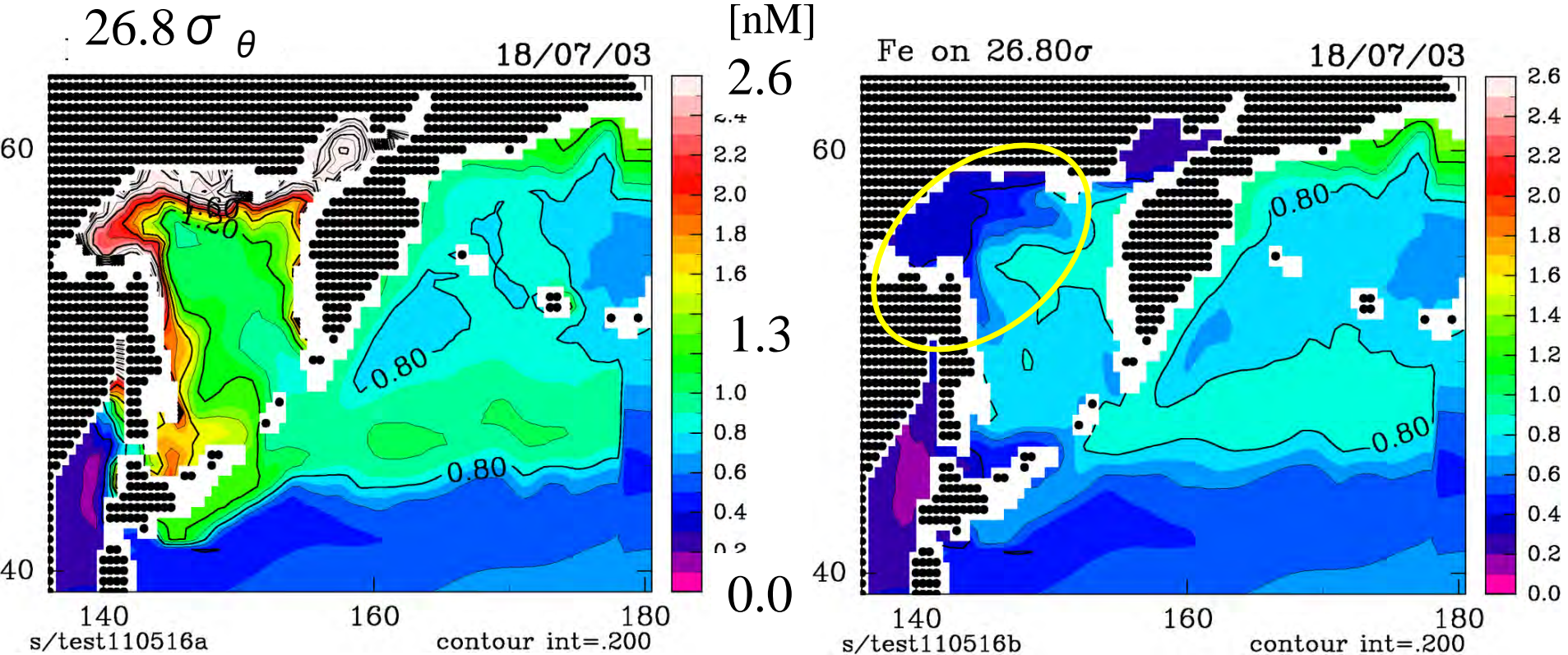


proportional to Fe'
(Fe=Fe'+FeL)
total ligand conc.=1.2 nM

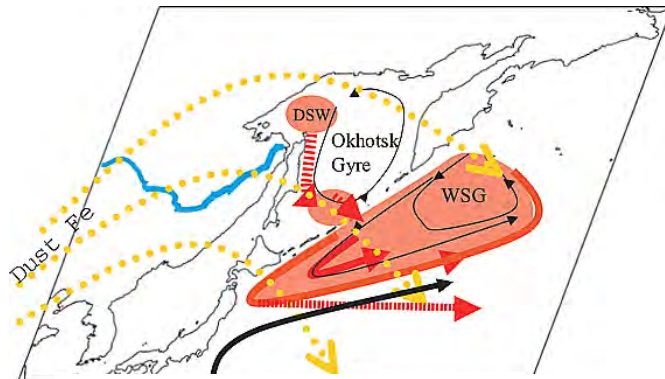


sediments (shelf shallower than 300 m)
1 μmol/m²/d

Fe distribution in the model

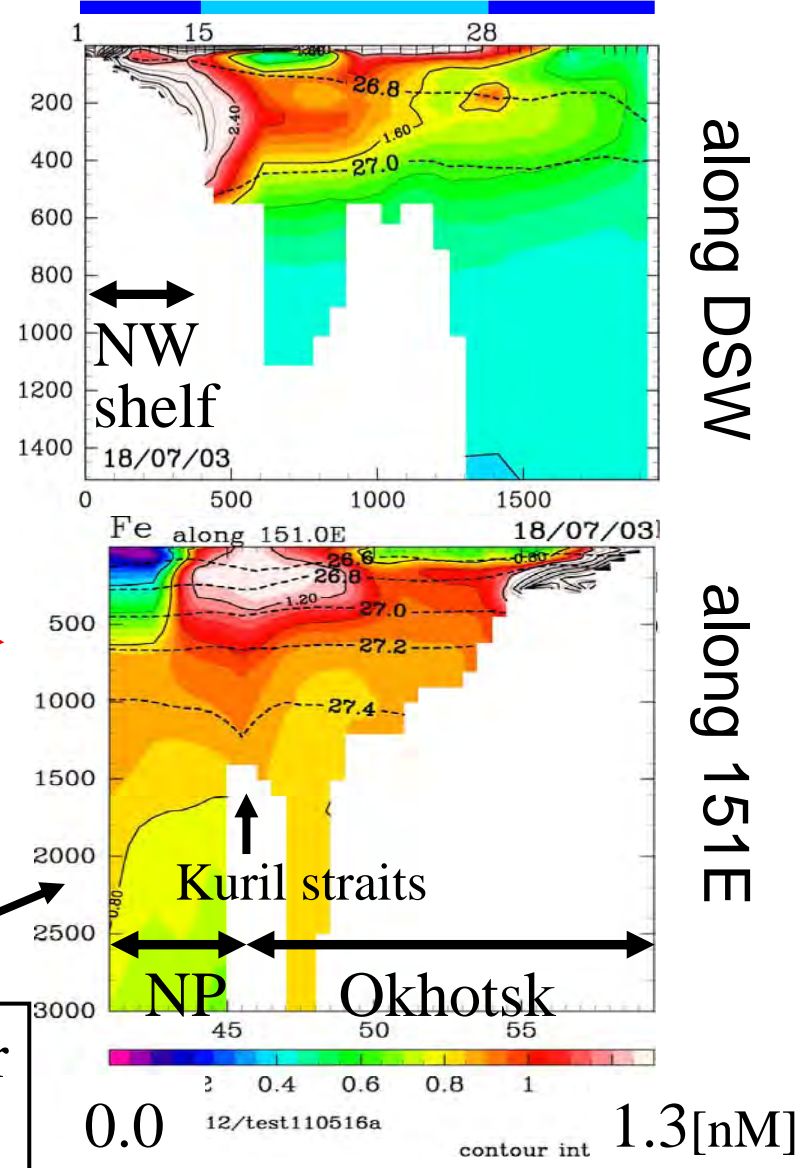
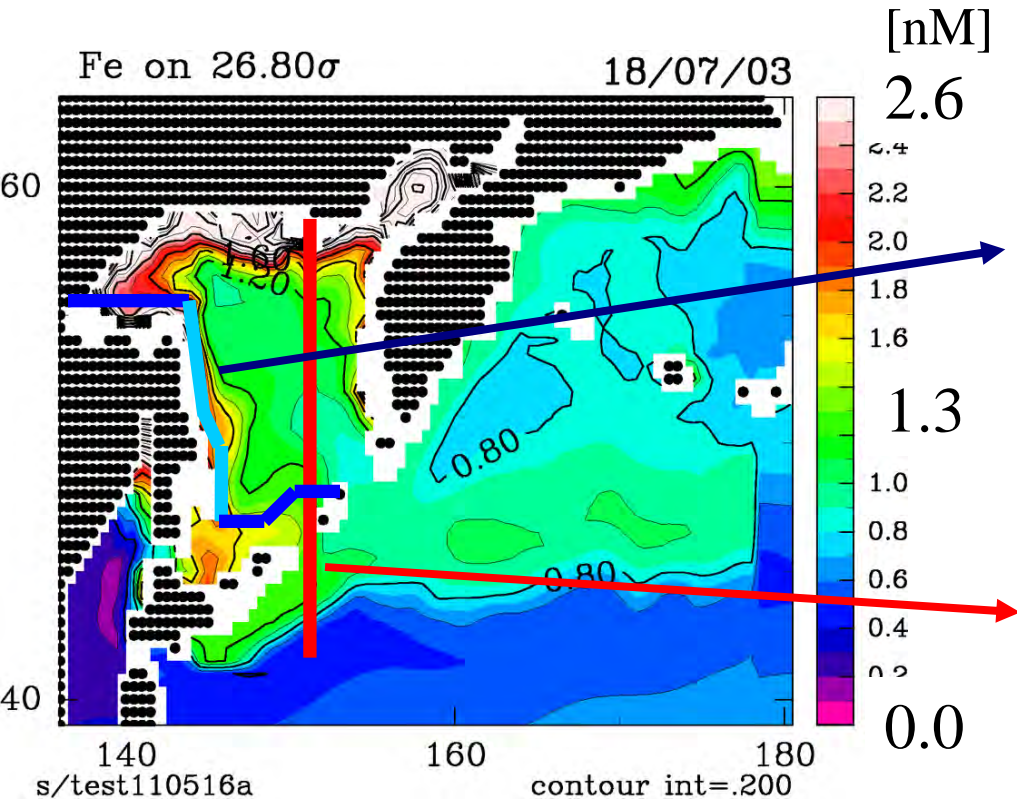


experiment without
sedimentary Fe



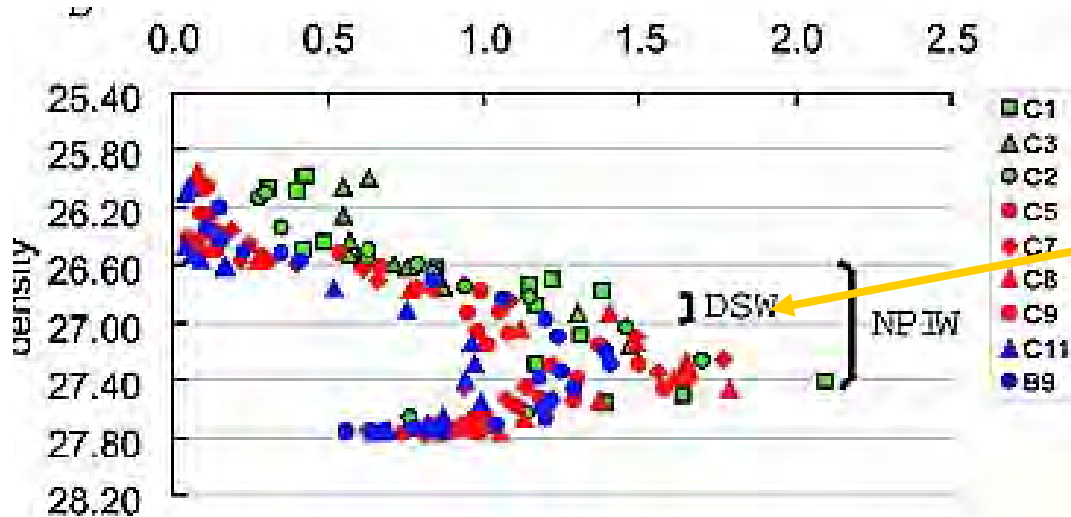
Nishioka et al. (2007)

Fe distribution in the model



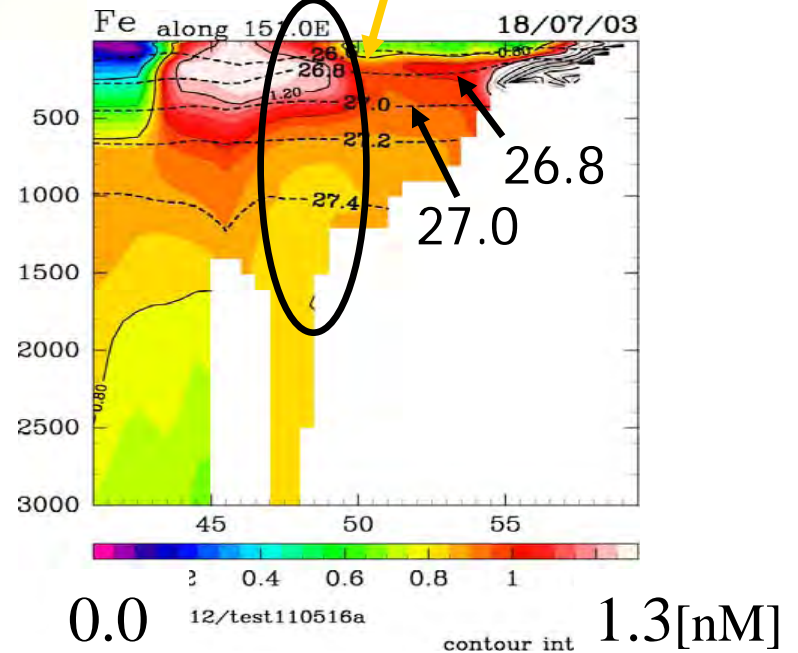
color scale is different by a factor of 2 only in this figure

Fe distribution in the model



greens: observed in Okhotsk
Nishioka et al. (2007, JGR)

Max. concentration
around $27 \sigma_{\theta}$



Summary

- We have constructed an ocean model that can represent well the circulation and ventilation in the intermediate layer.
 - tracer experiment, CFC simulation
- Combined with Parekh's model, the model represents iron flows in the intermediate layer, *although still in a preliminary stage.*

Thank you for your attention!

acknowledgement:

- This work has been supported by the New Energy and Industrial Technology Development Organization (NEDO).
- Aeolian dust data were provided by Dr. Mahowald.