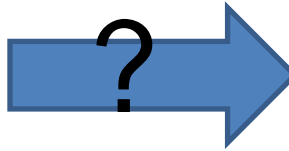
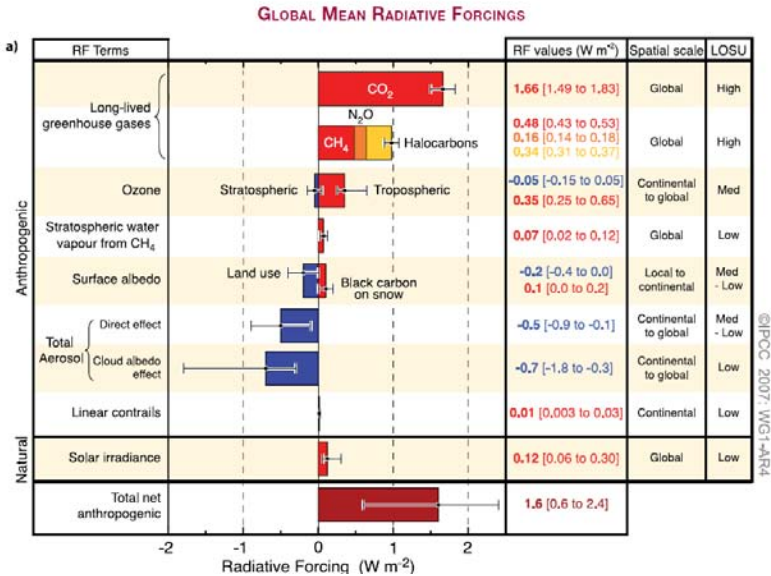


Climate change impacts on shelf and coastal marine ecosystems: contrasting ocean-shelf exchange, stratification and temperature effects on the NW European shelf

Jason Holt¹, Momme Butenschön², Sarah Wakelin¹, Yuri Artioli², Icarus Allen², James Harle¹, Jason Lowe³, Jonathan Tinker³

Climate impacts on shelf sea ecosystems

Changes in global radiative forcing

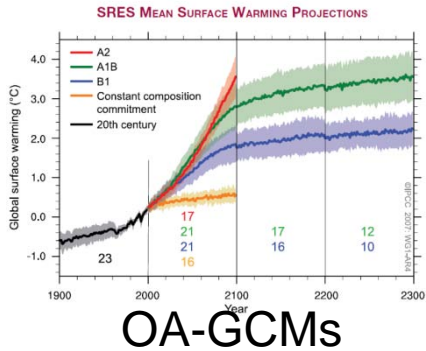


Changes in phytoplankton growth



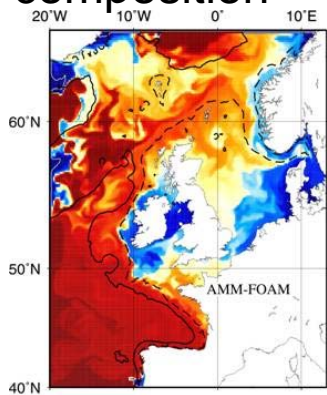
An extreme downscaling problem

Views of climate impacts on phytoplankton growth



The hypothesis:
Three key physical drivers

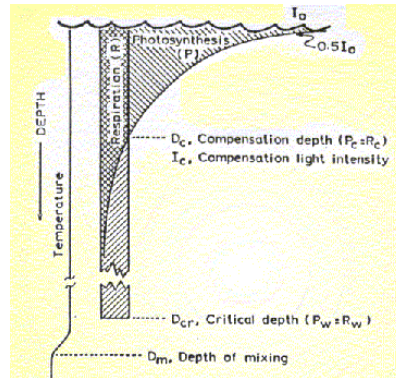
Biogeochemical composition



Huthnance et al (2011);
Wakelin et al (2009)

Ocean-shelf exchange

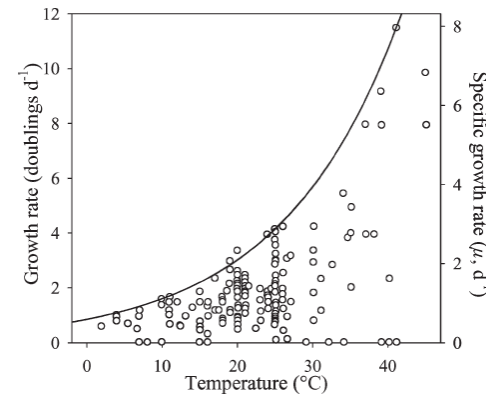
Phytoplankton blooms



Sverdrup (1953)

Turbulence-stratification-mixing
Interplay

Physiological response

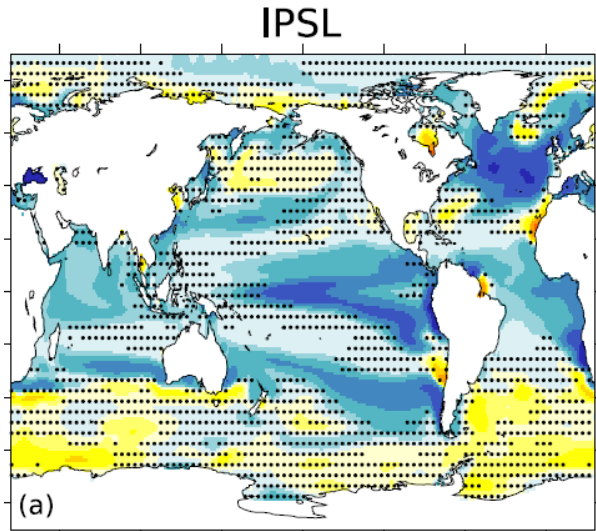


Eppley (1972)

Growth rate response to
temperature: autotrophic and
heterotrophic

Downscaling

Atmospheric conditions:
Temp., Precip., Wind, Radiation

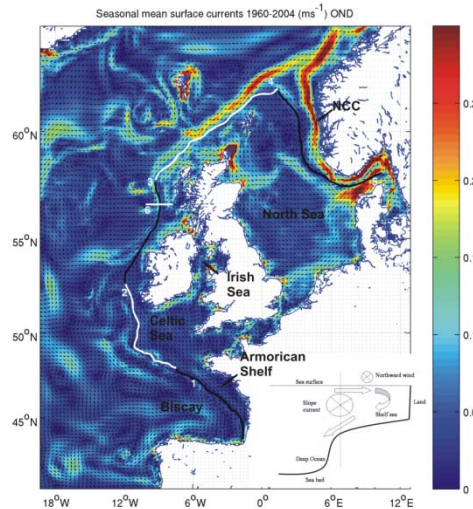


Global Earth System Model
Ocean-Atmosphere-Ice-
Ecosystem



Oceanic: T, S, N

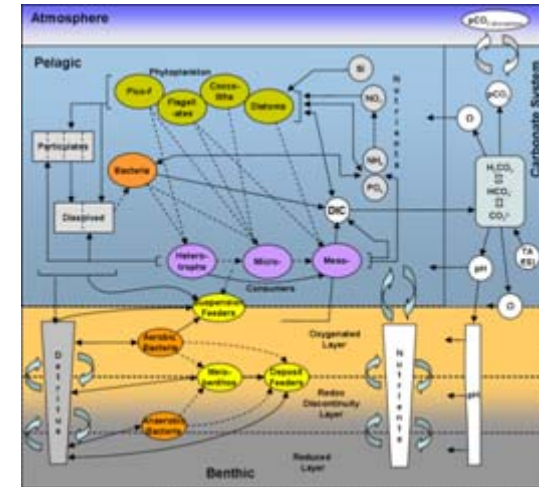
POLCOMS



Fine resolution regional
hydrodynamics model



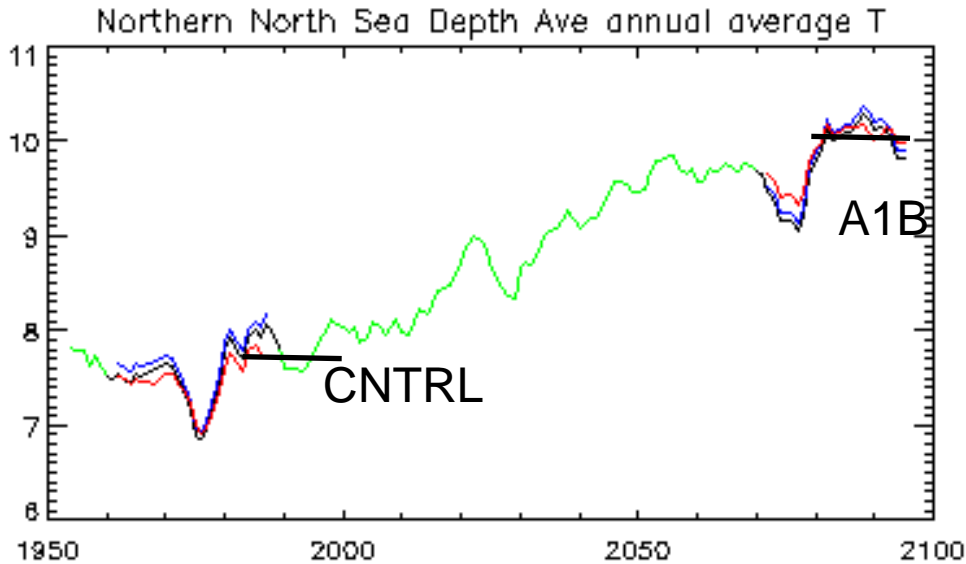
ERSEM



Regional ecosystem
model

Forced by IPSL-CM4 inc PISCES
Same models as used in QUESTFISH

The Timeslice Experiments



Consider A1B minus CNTRL

Assumes:

Timeslices are long enough to:

1. Adjust to new conditions ('spin-up')
2. Average-out internal variability

Both marginal here

A1B = 2082-2099

CNTRL = 1983-2000

18years + 5 year 'spin-up'

Reference:

ERA40 = 1960-2004

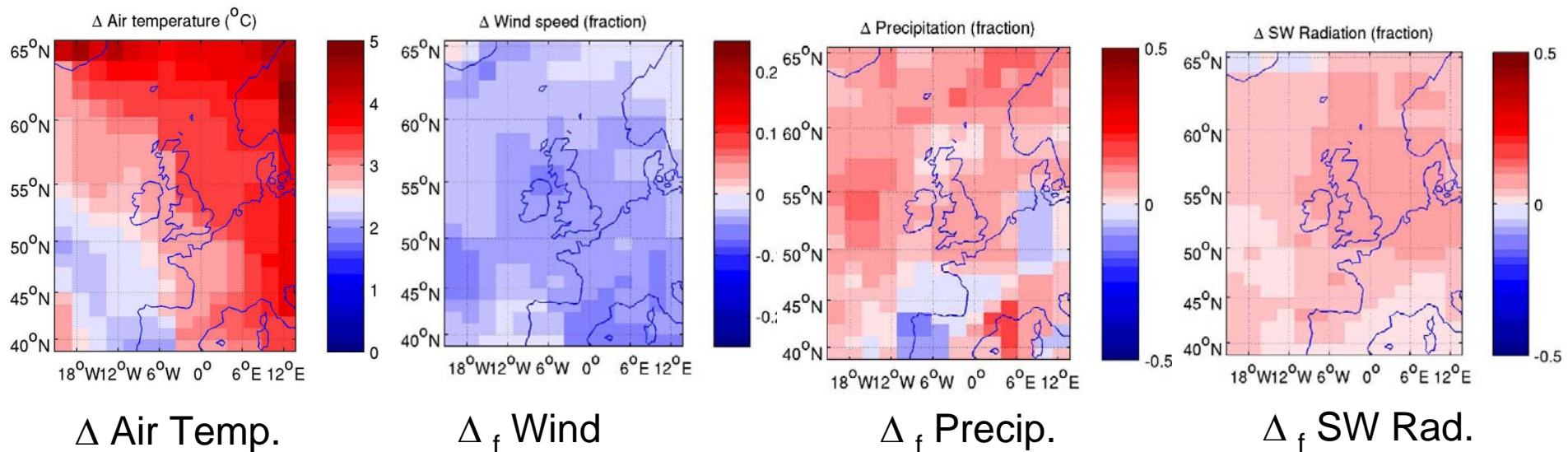
For details, validation and many results see:

Holt, et al , 2012. Oceanic controls on the primary production of the northwest European continental shelf: model experiments under recent past conditions and a potential future scenario.

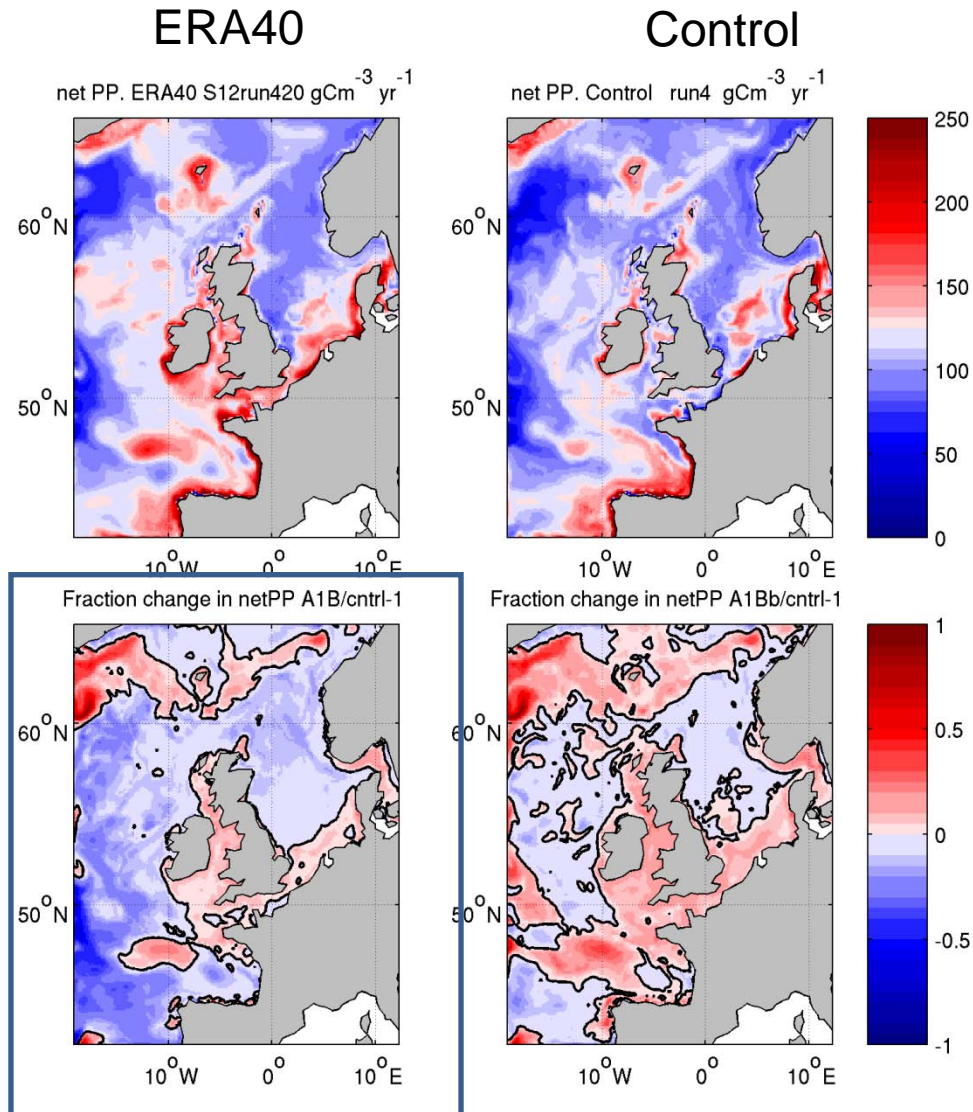
Biogeosciences 9, 97-117.

These experiments.....

- ARE NOT PREDICTIONS and NOT PROJECTIONS
- We shake the system and see what happens
- We choose for these SENSITIVITY experiments a perturbation that is DYNAMICALLY CONSISTENT and arises from a PLAUSABLE future radiative forcing scenario (A1B)
- This is NOT UNIQUE, but the diagnostics of the response gives important information on SYSTEMS BEHAVIOUR.



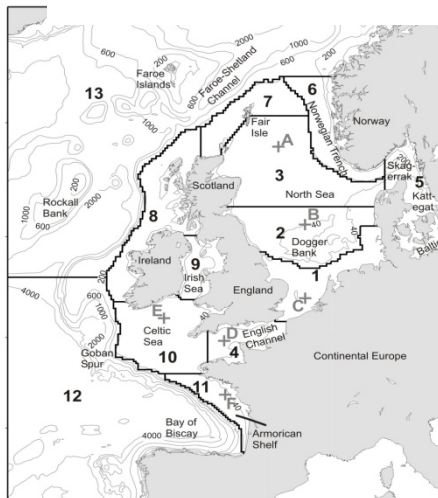
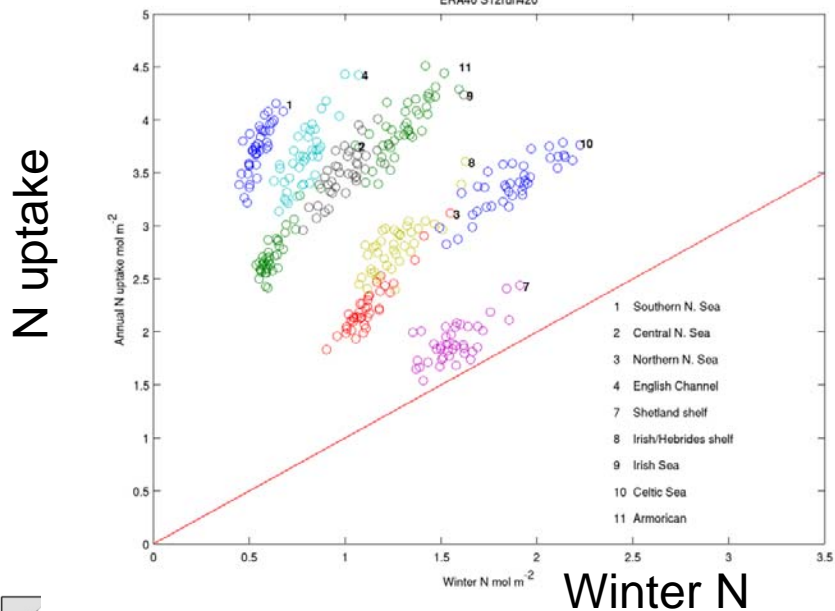
Change in net Primary Production



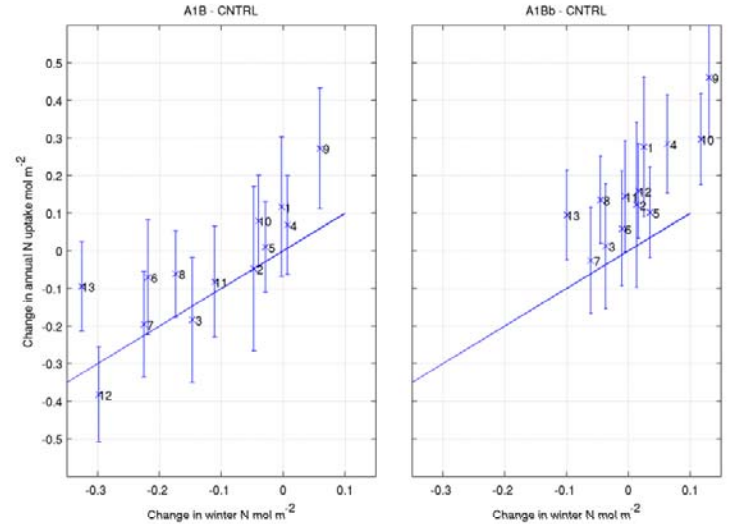
Fractional Change: A1B

Regional winter N v's N uptake following year

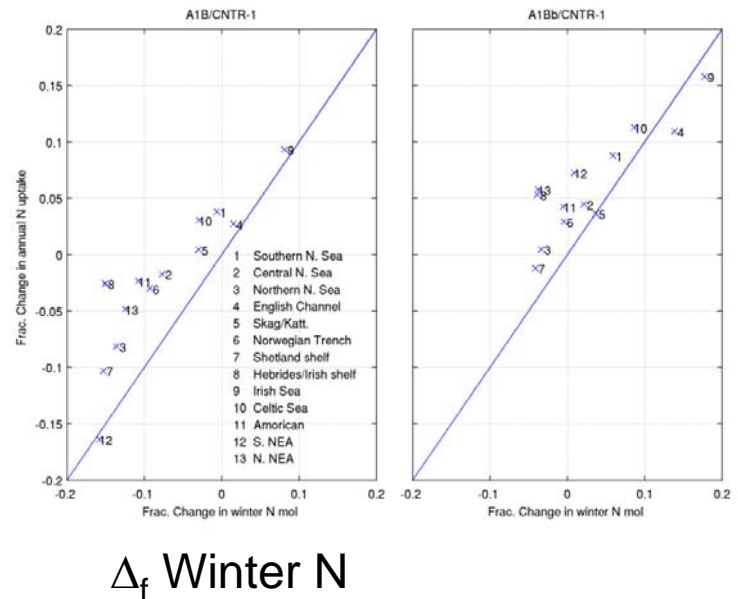
ER40 Reference



A1B - CNTRL



Δ_f N uptake



Transport across 200m isobath

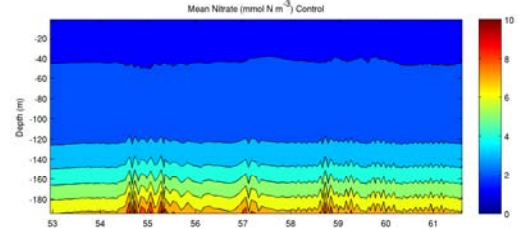
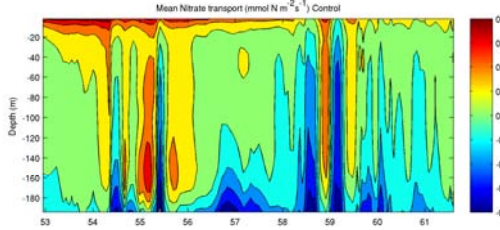
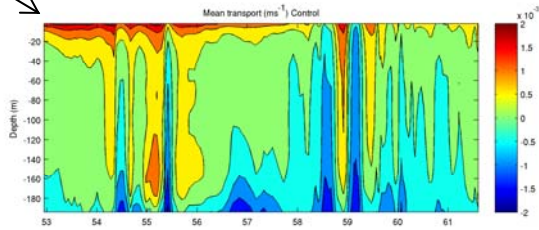
Volume Flux

DIN Flux

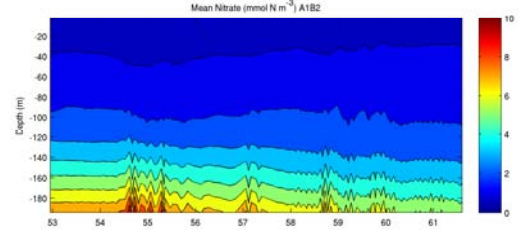
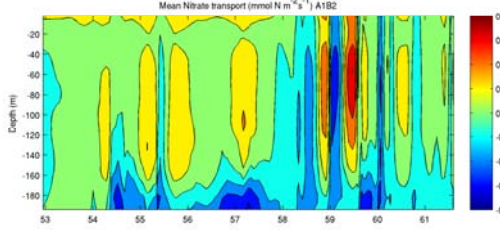
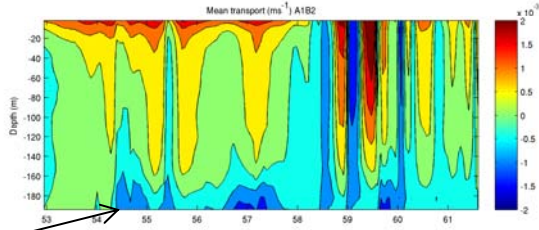
Mean DIN

On-shelf

CNTRL

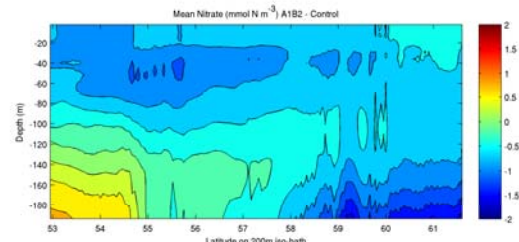
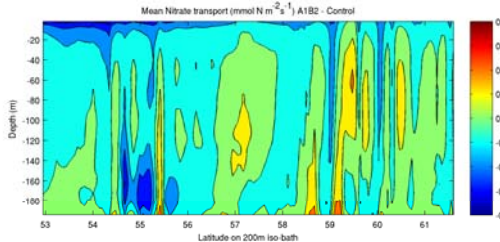
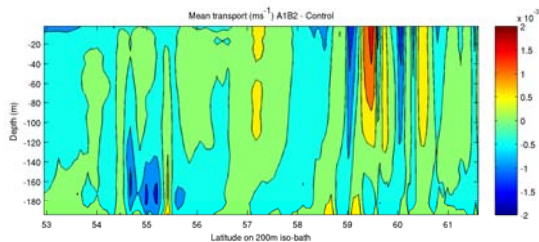


A1B



Off-shelf

A1B2
minus
CNTRL



53°N

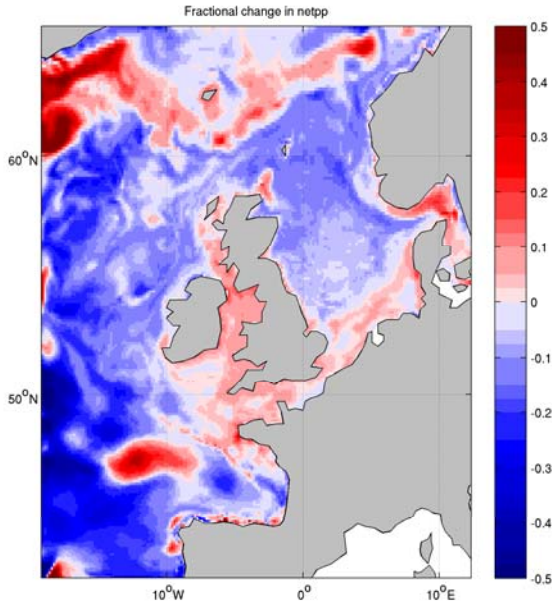
61°N

Small reduction in volume flux

See Holt et al *GRL* 2009

Substantial reduction in DIN concentrations

Process attribution experiments

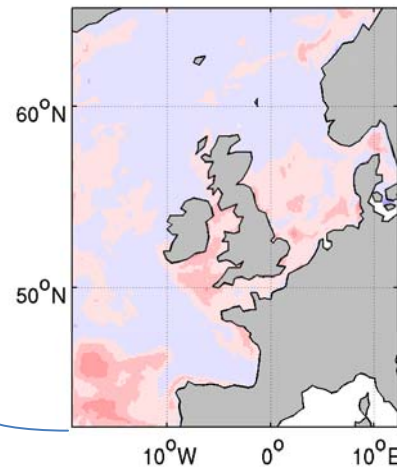
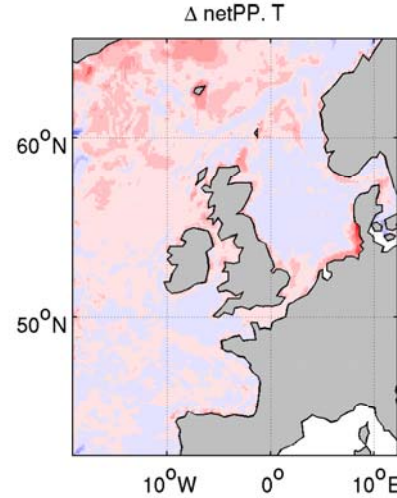


$$\Delta V_p = \Delta V - \Delta V_{p'}$$

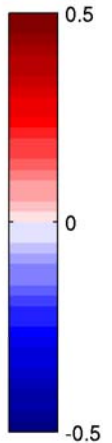
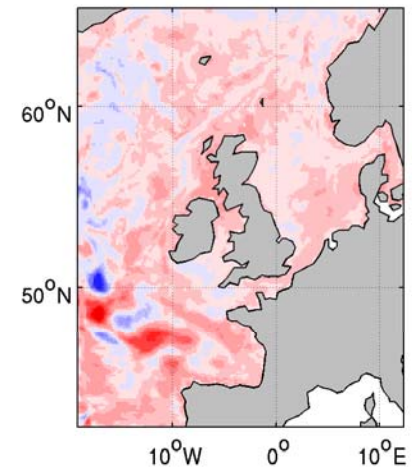
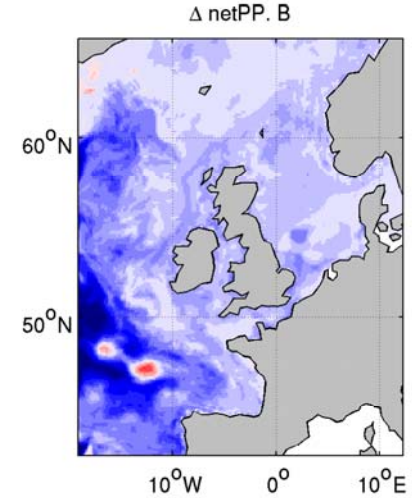
↑
Effect on change in V due to P, including all non-linearities

↑
Pair of time slices with process P absent

Temperature



Boundary N

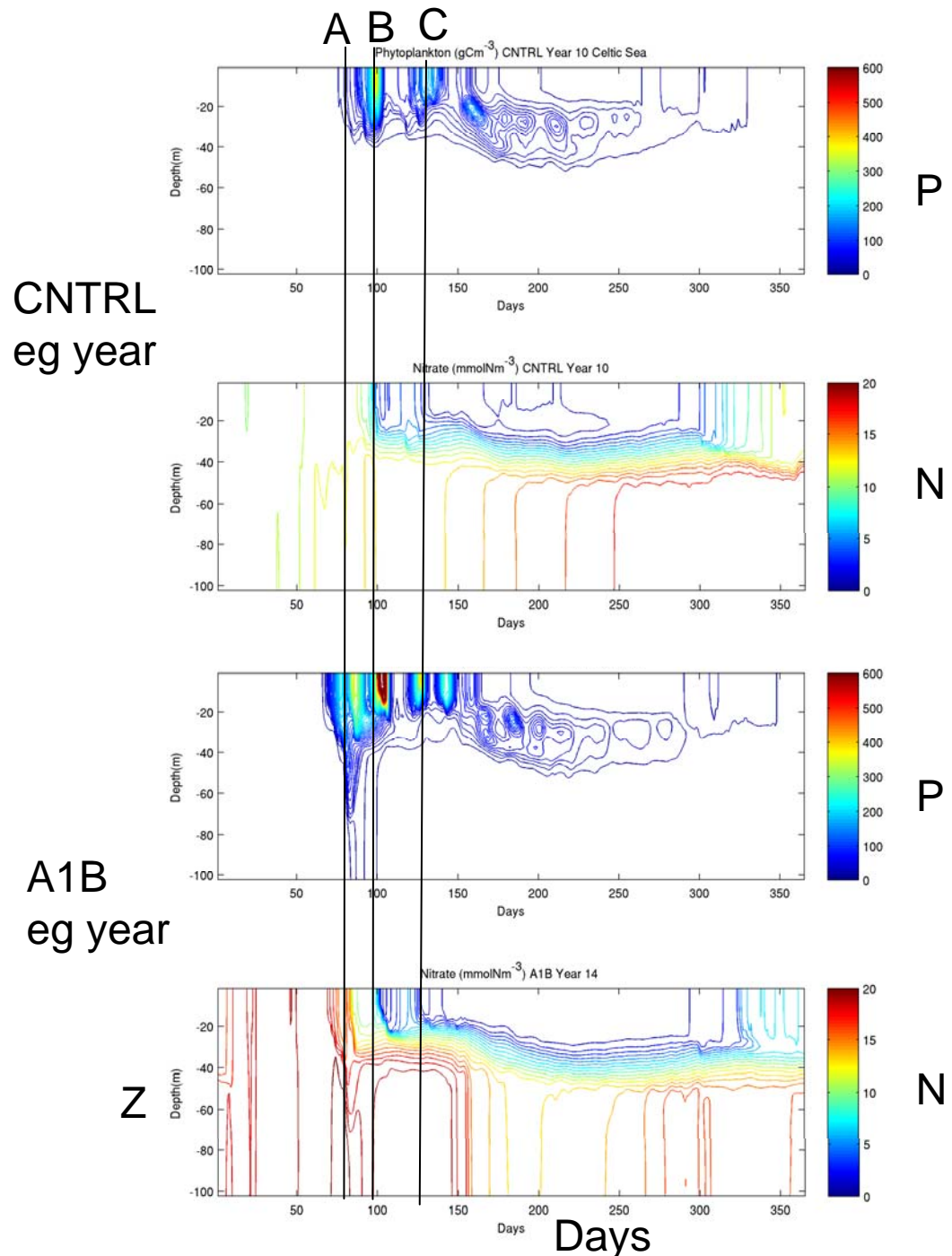


Seasonal Profiles: Celtic Sea

In A1B

- Earlier commence of growth (A)
- Stratification starts about same time (B)
- Surface layer depleted of N at about same time (C)
- Longer pre-stratification bloom*
- More efficient use of winter Nutrients
- Can be counteracted by reduced diapycnal mixing

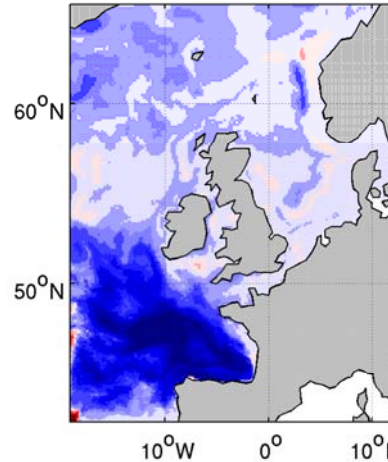
*Phytoplankton respond to reduced mixing but full depth nutrient flux still active.
Huisman et al *L&O* (1999)
'critical turbulence'



Changes in key times in seasonal cycle

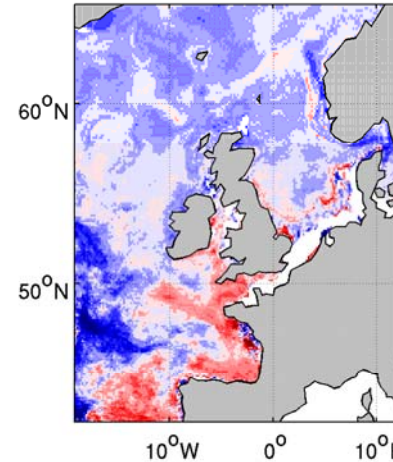
Growth Start

Δ Growth start time (days) A1B - CNTRL

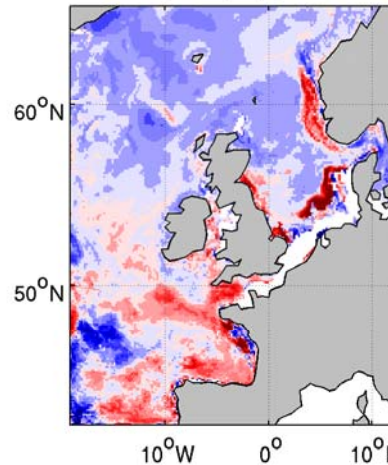


Strat. Start

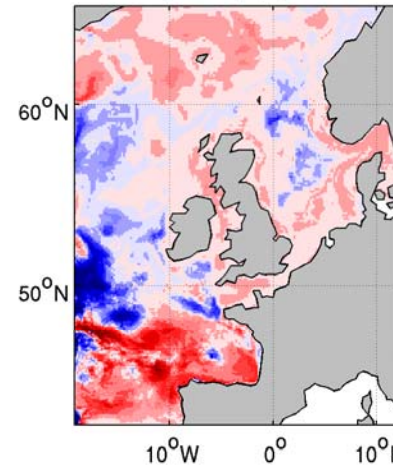
Δ Strat. start time (days) A1B - CNTRL



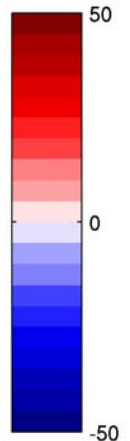
Δ Bloom stop time (days) A1B - CNTRL



Δ Growth stop time (days) A1B - CNTRL



Days



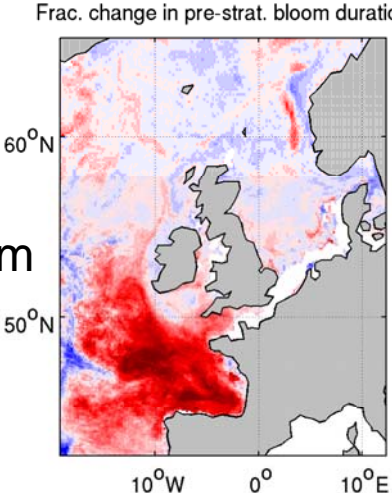
Bloom stop

Growth stop

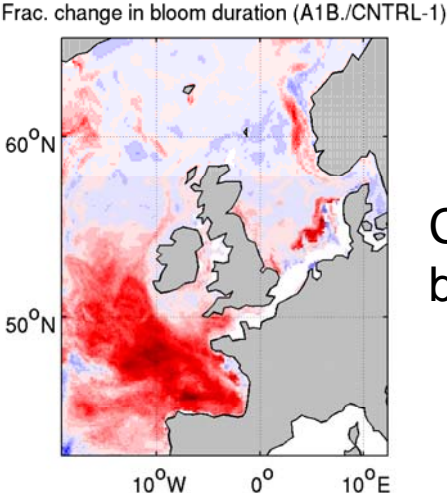
- Growth Start netpp $> 0.1 \text{gCm}^{-2}\text{d}^{-1}$
- Strat. Start = max dN/dt
- Bloom stop $N > 20\%$ wint N
- Growth Stop netpp $< 0.1 \text{gCm}^{-2}\text{d}^{-1}$

Changes in bloom production

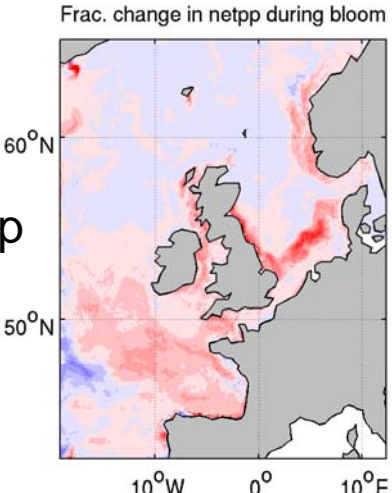
Change in pre-strat. bloom duration



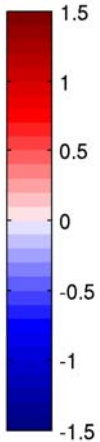
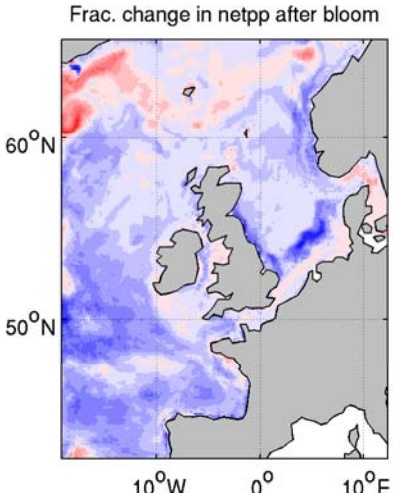
Change in bloom duration



Change in netpp during bloom



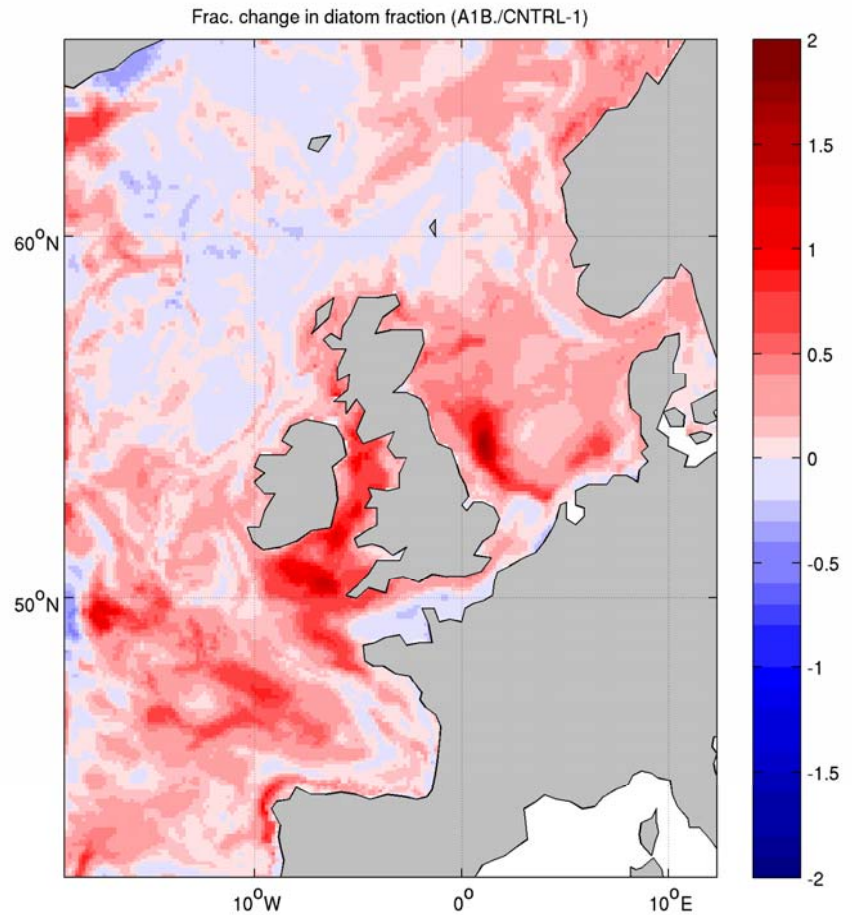
Change in netpp after Bloom.
(diapycnal mixing)



Other views of the system

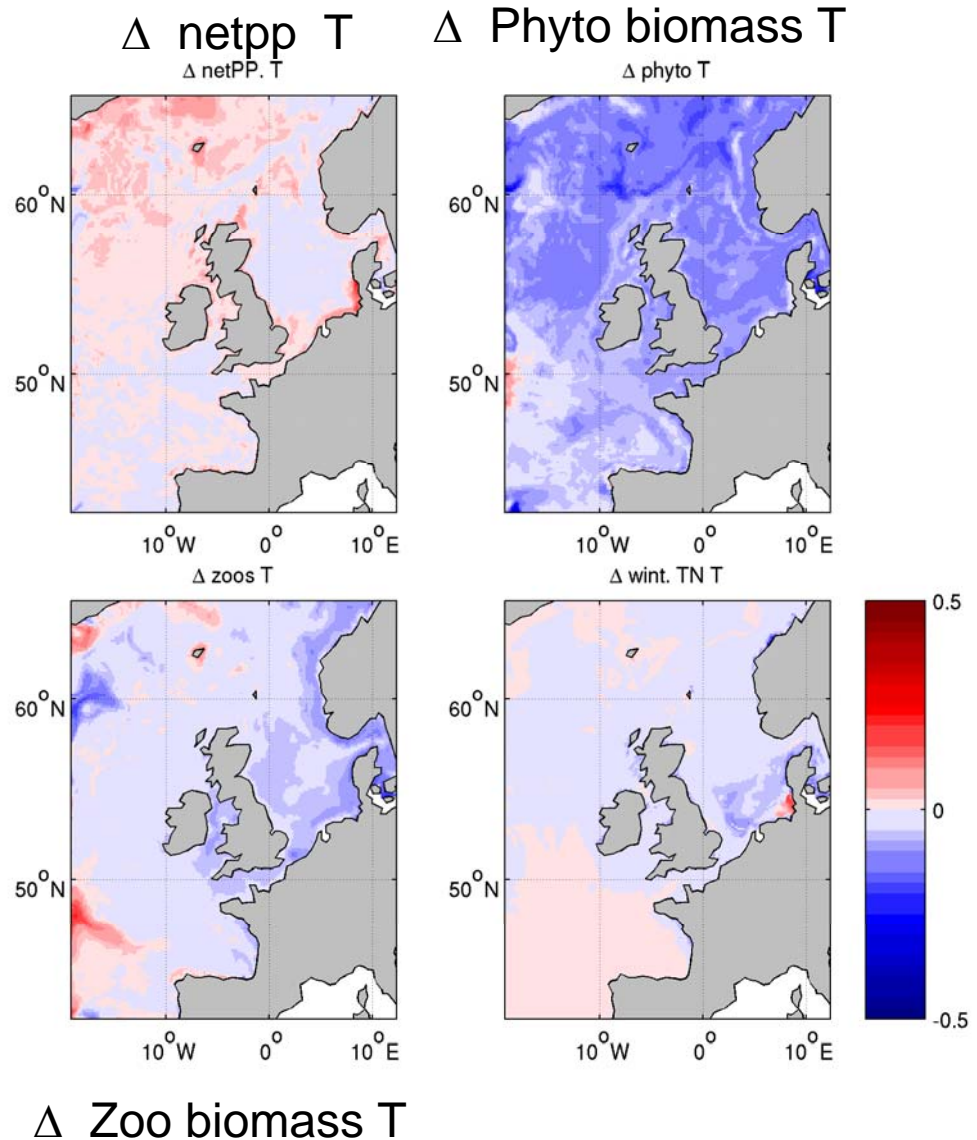
Change in community composition:
diatom fraction

Earlier blooms favour diatoms
and more efficient winter Si use



Temperature effects

- Temperature dependence is much more apparent on plankton biomass than netPP
- Heterotrophs and autotrophs have same q10 parameterisation



This is not the future.....

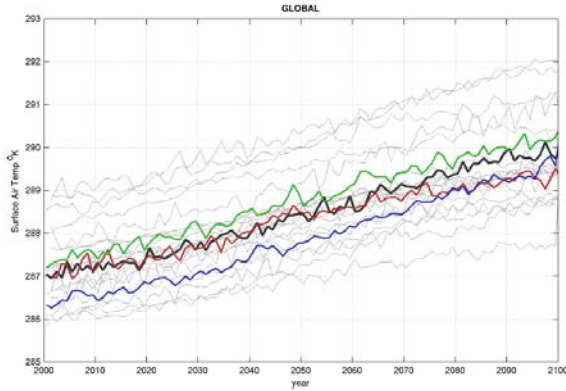
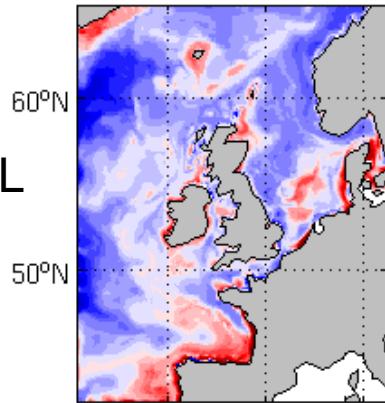


Figure 2: Global mean two metre air temperature from 23 IPCC AR4 climate simulations under the SRESA1B scenario. Highlighted are the IPSL (black), GFDL (red), HadCM3 (green) and HadGEM1 (blue) two metre air temperatures.

CNTRL

IPSL CM4

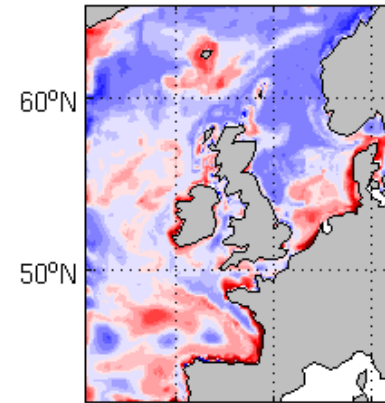
netPP IPSL cntrl $\text{gCm}^{-2} \text{yr}^{-1}$



Frac. change IPSL A1B/cntrl-1

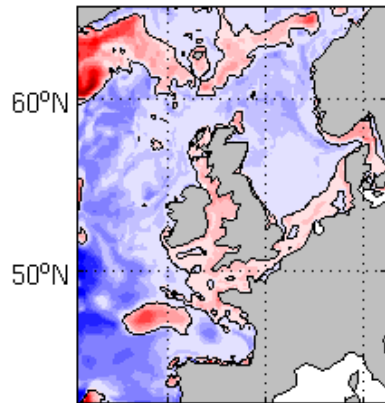
HadCM3

netPP HadCM3 cntrl $\text{gCm}^{-2} \text{yr}^{-1}$

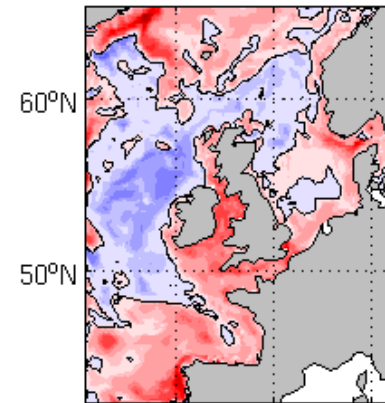


Frac. change HadCM3 A1B/cntrl-1

A1B



10°W 0° 10°E



10°W 0° 10°E

Conclusions

- Climate change impacts in shelf seas are highly nuanced, with multiple competing drivers
- This is a robust analysis of the response of the system, but need a description of likelihood to make projections
- Oceanic nutrients and stratification are first order controls of netPP changes, whereas temperature is a secondary effect in this model
- Subtle changes in mixing due to changes in light and wind mixing conditions have substantial effects on:
 - Bloom timing
 - Community structure
 - Total netPP
- But light (cloud cover) is an uncertain element in climate models
- Mixing in weakly/intermittently stratified conditions is weakness of the present generation of turbulence models